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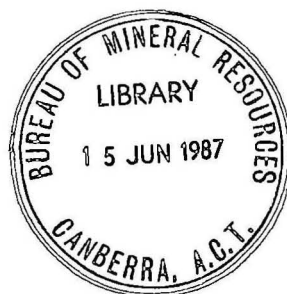
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DEPARTMENT OF NATIONAL DEVELOPMENT
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GEOLOGY AND GEOPHYSICS

RECORDS 1956, N^o. 136

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SEISMIC REFRACTION SURVEY OF
No. 1 POWER STATION SITE,
KIEWA, VICTORIA



by

D. F. DYSON and M. J. O'CONNOR

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A B S T R A C T

Details and results are given of a seismic refraction survey made in response to an application by the State Electricity Commission of Victoria, to investigate the area proposed for the No. 1 power station site at Kiewa, Victoria.

The purposes of the survey were to determine the thickness of the weathered layer and to supply a contour plan of the depth to unweathered rock from which sites worthy of more detailed investigation could be suggested.

The thickness of soil over the area ranges from zero to 43 feet. The weathered layer of granodiorite, which varies considerably in character (as shown by the velocity range from 3,900 to 6,200 feet/sec.), has a maximum depth of 213 feet. The seismic velocity of the unweathered granodiorite bedrock ranges from 11,000 to 18,000 feet/sec., the lower velocity being correlated to shearing in the granodiorite.

Comparison between seismic results and drilling data along and near the traverses indicates that depths determined by the seismic method have a standard error of approximately 9 feet.

Locations are suggested for further drilling investigations for both a surface and an underground power station.

1. INTRODUCTION

The planned Kiewa No. 1 Power Station is part of the Victorian State Electricity Commission's Kiewa No. 1 Power Development Scheme, which consists of the Rocky Valley Dam, the Pretty Valley diversion tunnel, penstock line and a power station.

The seismic refraction survey of the power station site was made by the Bureau in response to an application by the State Electricity Commission, during October and November 1955. A preliminary report on the survey (Dyson, 1956) was submitted to the Commission in July, 1956. The co-ordinates of the approximate centre of the area, on the grid of the 4-mile military maps, are 1356500 North and 1560500 East.

The purposes of the survey were to determine the depths to unweathered bedrock in the surveyed area and to supply a contour plan of depths to unweathered rock to assist in the design of an underground power station.

The geophysical party consisted of D. F. Dyson (party leader) and M. J. O'Connor, geophysicists, and J. P. Piggott, field assistant. The Commission supplied additional assistants and carried out the necessary topographical survey of the traverses.

2. GEOLOGY

The country rock consists of granodiorite which weathers at and near the surface to a completely decomposed rock. Parts of the hill slopes are covered with sandy material (quartz from the granodiorite) washed down hill.

The shear and fault pattern in the area is relatively complicated. One system of almost parallel faults strikes N50E to N80E, and another system with basic dykes associated with it strikes about N60W.

A deeply weathered zone (disclosed by excavation and diamond drilling) which crosses the north-western part of the area is known as the "Pretty Valley Fault Zone". Some features of the fault pattern are shown on Plate 1.

From results obtained in other areas, it may be assumed that the granodiorite in, or close to, the fault and shear zones is usually weathered to greater depths than in places where faulting and shearing are absent. This was confirmed by surface mapping, excavations and diamond drilling.

3. METHODS

The seismic refraction method depends on the degree of contrast between the elastic properties of the different zones of rock. In foundation investigations the objects are the determination of the depths to elastic discontinuities and an indication of the rocks concerned. The physical laws involved resemble those of optical phenomena in that energy propagated through the ground will undergo reflection and refraction at discontinuities.

An explosive charge detonated at or near the ground surface produces a train of seismic waves. In the normal refraction method the charge is in line with a series of detectors (geophones), known as a spread, and the first arrival times of the seismic waves are recorded. A graph of first arrival times against distances from the shot point is called a time-distance curve.

A technique known as the Method of Differences (Heiland, 1946, p.548) was used. If A is a shotpoint at one end of the spread then the time taken for the energy to reach a point B within the spread is TBA and the time to a point C beyond B is TCA. Similarly, shooting from C the time for the energy to travel from C to B is TBC and from C to A the time taken is TAC. The times TCA and TAC are the reciprocal times, and should be equal.

The depth is computed from the vertical travel time TD which is defined as $\frac{1}{2}(TBA + TBC - TCA)$. The depth, D, becomes

$$\sqrt{\frac{V_2^2}{V_2^2 - V_1^2}} \cdot TD V_1, \text{ where } V_1 \text{ is the velocity of the first recorded wave through the layer above the discontinuity, and } V_2 \text{ is the velocity through the layer below. For } V_2 > 3V_1,$$

$$\sqrt{\frac{V_2^2}{V_2^2 - V_1^2}} \text{ approaches the value 1, and the depth becomes } V_1 TD.$$

The velocities of the successive refracting layers are determined from the time-distance curves. Seismic velocities give a fair indication of the type of rock encountered but a certain amount of geological control is necessary before a more definite interpretation can be made. A sandstone for example, may have the same velocity as a fractured granodiorite.

Two types of geophone spread were used :-

- (a) Normal spreads, in which the geophones were placed at 50-foot intervals, with shot distances of 50 and 300 feet.
- (b) Weathering spreads, with geophone intervals of 10 feet and shot distances of 10, 50, 150 and 300 feet.

The vertical travel times and the velocities of the near-surface layers were computed from data obtained from these spreads.

A 12-channel portable "Century" Seismograph, Model 506, with Technical Instrument Company geophones (natural frequency 19 c.p.s.), was used.

Operational Difficulties.

The following difficulties were encountered during the survey :-

- (i) Deep soil over most of the area (up to 43 feet thick) necessitated the use of large charges (a maximum of 20 lbs.) of gelignite and shot holes as deep as 12 feet.
- (ii) Construction work in or near the area resulted in high noise level on the records, the thick soil layer acting as a channel for low frequency waves.
- (iii) Power line interference, both directly and by transmission through the ground, particularly when wet.
- (iv) The energy transmitted through the refractors was dissipated by faults and shear zones.
- (v) The presence of buildings restricted the selection of shot points and the amount of the charge.

To improve the reliability of the records the spreads were often shot a second time during periods of comparative quietness and using different shot points.

4. RESULTS.

The results of the survey are shown by the cross-sections on Plates 2 to 4 and the contour map of the surface of the "high velocity rock" on Plate 1. Table 1 shows the correlation between velocity and rock type identified by surface observations and diamond drilling data.

TABLE 1

Rock Type	Seismic Velocity (feet per second)
Soil	1,000 \pm 100
Weathered granodiorite	3,500 \pm 1,000
Fractured and partly weathered granodiorite	7,000 \pm 2,000
Fractured or sheared granodiorite (unweathered)	12,500 \pm 1,500
Unweathered granodiorite	17,500 \pm 500

The soil ranges in thickness from zero to approximately 40 feet. Soil velocities in the area range from 900 to 1,100 feet/sec., and a uniform value of 1,000 feet/sec. was used.

Average velocities in weathered rock underlying the soil range from 3,900 feet/sec. to 6,200 feet/sec. The lower velocities in this range are considered to represent completely weathered granodiorite and the higher velocities fractured and partly weathered granodiorite.

The term "high velocity rock" is used for rocks with velocities ranging from 11,000 feet/sec. to 18,000 feet/sec. Velocities of 11,000 to 15,000 feet/sec. were recorded on traverses A, B, C, and D where they intersect the Pretty Valley Fault Complex. The velocity of the refractor along traverse H, which is within, and parallel to, the boundaries of the Pretty Valley Fault Complex, ranges from 16,000 to 17,000 feet/sec. This indicates that the major fracture pattern is parallel to the strike (N60°E) of the Pretty Valley Fault Complex. High velocities of 17,000 to 18,000 feet/sec. indicate unweathered, unfractured granodiorite.

The large range of velocities in the "high velocity rock" (from 11,000 feet/sec. on traverse B to 18,000 feet/sec. on traverse G) is due to the presence of shear zones, faults and dykes. The local influence of these features cannot be distinguished clearly on the time-distance curves, and it is therefore impossible to determine their position precisely.

In the preliminary report on the survey (Dyson, 1956), a contour map of the "high-velocity rock" was drawn from the sections on Plates 2 to 4, using geophysical data only. As

these data can strictly be used only for points along the traverses, a certain amount of interpolation is necessary between traverses in order to complete the contour map. The uncertainty of such interpolation is increased in the present survey by the presence of structural complexities such as shears with intersecting faults. This is particularly so in the area bounded by traverses A, E, F and part of C. Drill hole data indicate that the "high-velocity rock" is at a greater depth in this area than that shown in the contour map contained in the preliminary report. That contour map has now been replaced, therefore, by the map shown on Plate 1 of this report which has been compiled from both geophysical and drilling data.

The map shows a trough striking east-north-east in the northern part of the area. It coincides with the Pretty Valley Fault Complex and suggests that weathering penetrates downwards along fault or shear structures. There are indications in the geophysical data of other smaller structures, but the information available is insufficient to delineate these structures. An example is the trough on traverse F between stations F10 and F11. This may correspond to the geologically identified Fault B1, but for precise positioning of this fault at the level of the "high velocity rock", diamond drilling investigations would be required.

5. ACCURACY

Errors in depth estimates to the "high velocity rock" are due mainly to errors in the calculated velocity in the weathered layer and to errors in the vertical travel times.

It appears from the data the the velocity in the weathered layer can be determined with a standard error of 715 feet/sec. (number of data used, 15). This means that 95 per cent of the velocity determinations are within twice the standard error, i.e. within ± 1430 feet/sec. With a vertical travel time of t msec., this represents an accuracy of within $\pm 1.4 \times t$ feet.

Errors in vertical travel time depend mainly on the accuracy with which the records can be read and the ability to make reliable corrections for soil thicknesses ("soil correction"). If the errors in reading the records are within ± 1 msec., then the accuracy in the vertical travel time is about $\pm \frac{1}{2}\sqrt{3}$, i.e. within about ± 1 msec.

The maximum error in the estimates of soil thickness is about ± 5 feet. With a weathered layer velocity of 5,000 feet/sec., and a soil velocity of 1,000 feet/sec., the resultant errors in soil correction are within $-\frac{5}{1} + \frac{5}{5} = \pm 4$ msec.

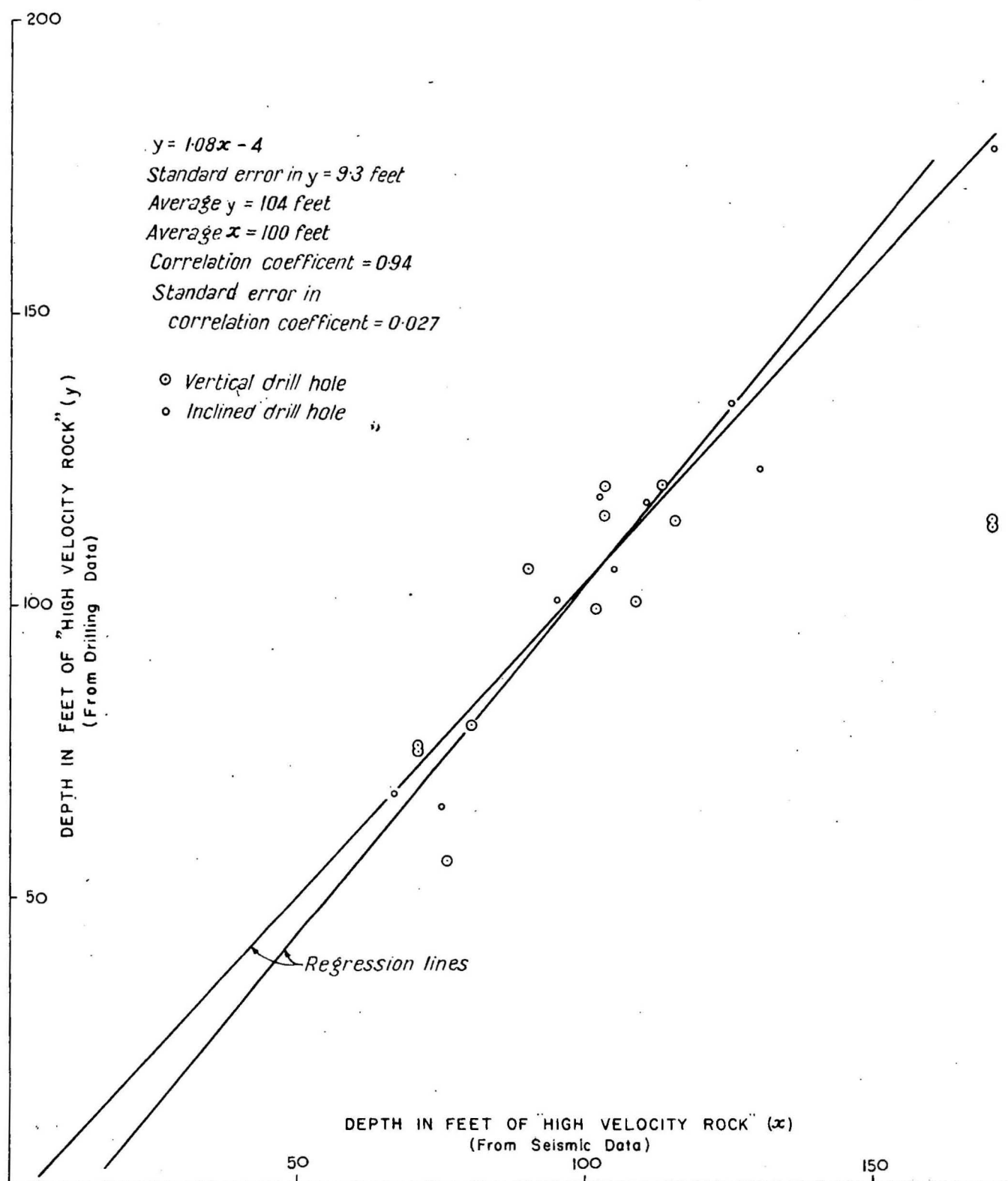
Based on the above estimates, the errors in vertical travel time are within $\sqrt{12 + 42}$, i.e. within about ± 4 msec. With weathered layer velocities of about 5,000 feet/sec., this results in errors of within ± 20 feet.

The total error in depth estimates to the "high velocity rock" is then within $\pm (20 + 1.4t)$. For an average velocity in the weathered layer of 5,000 feet/sec., the maximum possible percentage errors are as shown in Table 2.

The depths of the "high velocity" refractor obtained by drilling and by the seismic method are compared in Table 3.

Drill holes omitted from Table 3 include :-

FIGURE 1



SEISMIC REFRACTION SURVEY
 OF No. 1 POWER STATION SITE, KIEWA, VICTORIA
 STATISTICAL ANALYSIS
 OF DEPTH DETERMINATIONS

5.

- (i) Those outside the area surveyed.
- (ii) Inclined holes drilled to investigate vertical structures crossing the area.

TABLE 2

Depth (ft.)	Maximum Error (ft.)	Maximum Percentage Error
25	27	108
50	34	68
100	48	48
150	62	41
200	76	38

TABLE 3

Drill Hole No.	Depth to high velocity rock (ft)		Difference	Difference as % of drilling data
	Drilling data	Seismic data		
1086 Vertical	62	76	+ 14	+ 23
1087 "	121	103	- 18	- 15
1088 "	100	102	+ 2	+ 2
x1089 "	107	90	- 17	- 16
x1090 "	76	71	- 5	- 7
1091 "	116	103	- 13	- 11
x1094 "	80	80	0	0
x1095 "	121	113	- 8	- 7
1097 "	101	108	+ 7	+ 7
1180 "	100	109	+ 9	+ 9
1184 "	75	71	- 4	- 5
1166 Inclined	66	75	+ 9	+ 14
1169 "	118	110	- 8	- 7
1170 "	124	130	+ 6	+ 5
1171 "	68	67	- 1	- 1
1174 "	102	95	- 7	- 7
1176 "	135	125	- 10	- 7
1178 "	179	170	- 9	- 5
1182 "	107	105	- 2	- 2
x1185 "	119	102	- 17	- 14

In the inclined holes the depth given is the vertical depth from the surface to the point at which the drill hole intersected the refractor.

In the drill holes marked with an asterisk the seismic refractor does not represent unweathered granodiorite, but fractured granodiorite.

In Figure 1 the depths calculated from the seismic data are plotted against the depths obtained from the drilling operations. The line which best fits the data is expressed by the equation :-

$$y = 1.08x - 4$$

Standard error in y is 9.3 feet

Average x = 100 feet

Average y = 104 feet

The correlation coefficient is 0.94 with a standard error of 0.027. As a close approximation the seismic data along the traverses indicate the depth of the "high velocity rock" fairly accurately, with a standard error of 9 feet in the range from 60 to 170 feet.

These results indicate that the accuracy is within the limits shown by Table 2;

6. CONCLUSIONS AND RECOMMENDATIONS

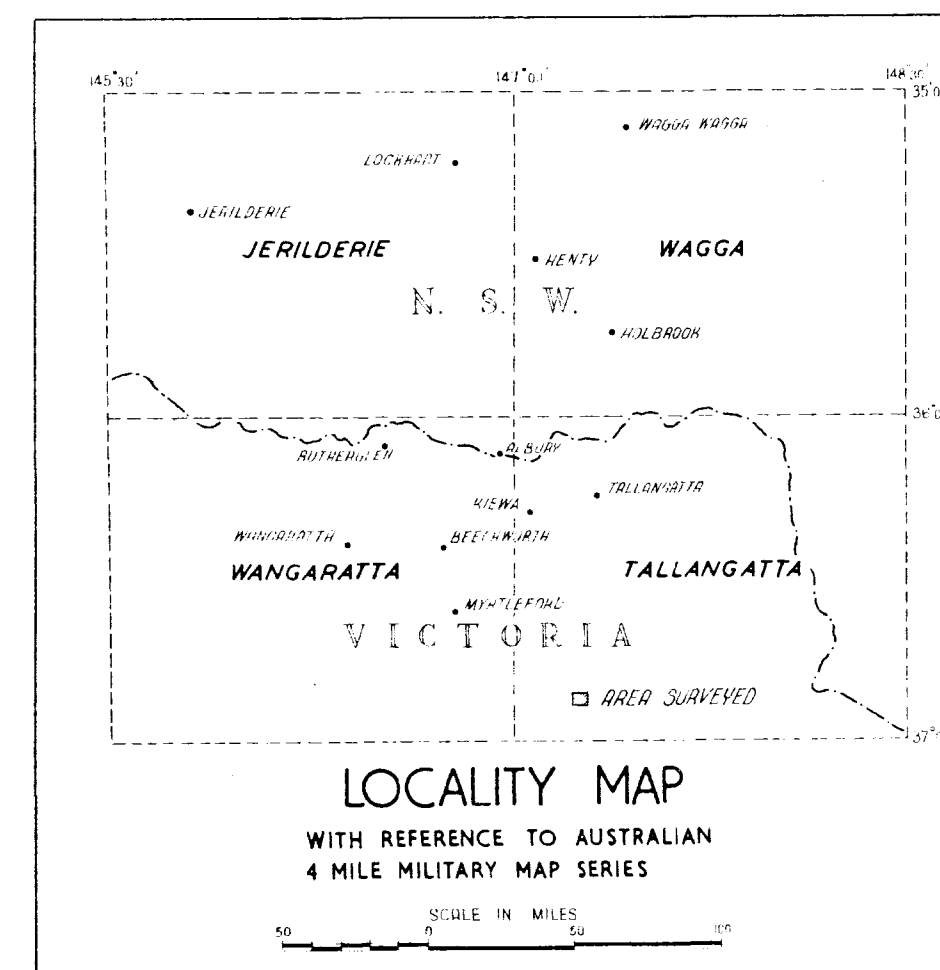
The survey has provided information for the selection of two possible areas worthy of detailed investigation, one for a surface power station and the other for an underground station.

It is considered that the area bounded to the south by traverse G and between stations G11 and G16 is the most suitable site for further investigations for a surface power station. The depth to suitable foundation rock should be of the order of 50 feet. Upon the removal of the weathered layer in this area precautions may be necessary to guard against slipping of the weathered material on the uphill side of the area. The most likely point at which slipping may take place is near stations C1 and C2.

Detailed investigations for an underground station should be made within the triangular area bounded by lines joining E1, F8 and G21 + 50 feet. This area should have sufficient cover of unweathered rock, with no major intersecting faults or shear zones. The tail race from this site would pass through Fault B and possibly other faults which would require investigation by diamond drilling.

7. REFERENCES

- | | | |
|---------------------|---|--|
| Beavis, F., 1956 | - | Report on the Geological Investigation for the No. 1 Underground Power Station, S.E.C. (Internal Report). |
| Dyson, D.F., 1956 | - | Preliminary Report on a Seismic Refraction Survey of No. 1 Power Station Site, Kiewa, Vic. Bur. Min. Res. Geol. & Geophys., Records 1956 No. 53. |
| Heiland, C.A., 1946 | - | GEOPHYSICAL EXPLORATION, Prentice-Hall, New York. |



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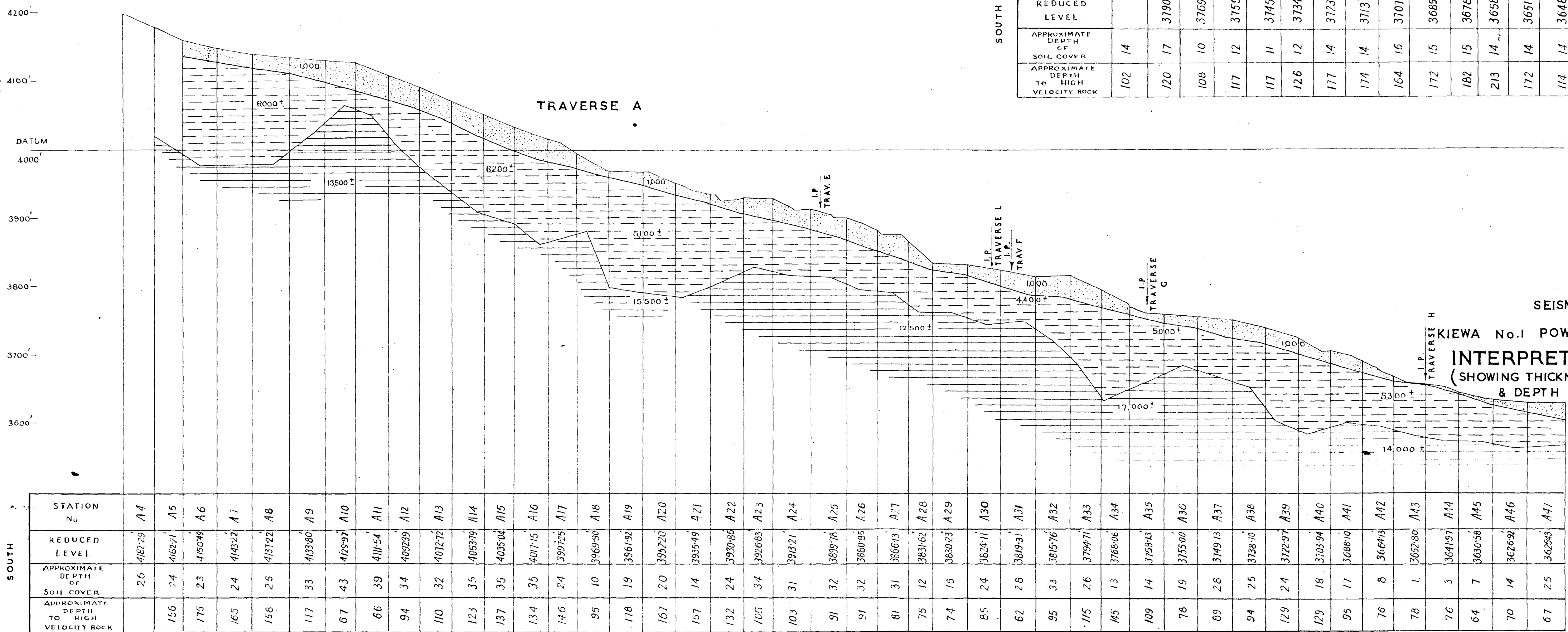
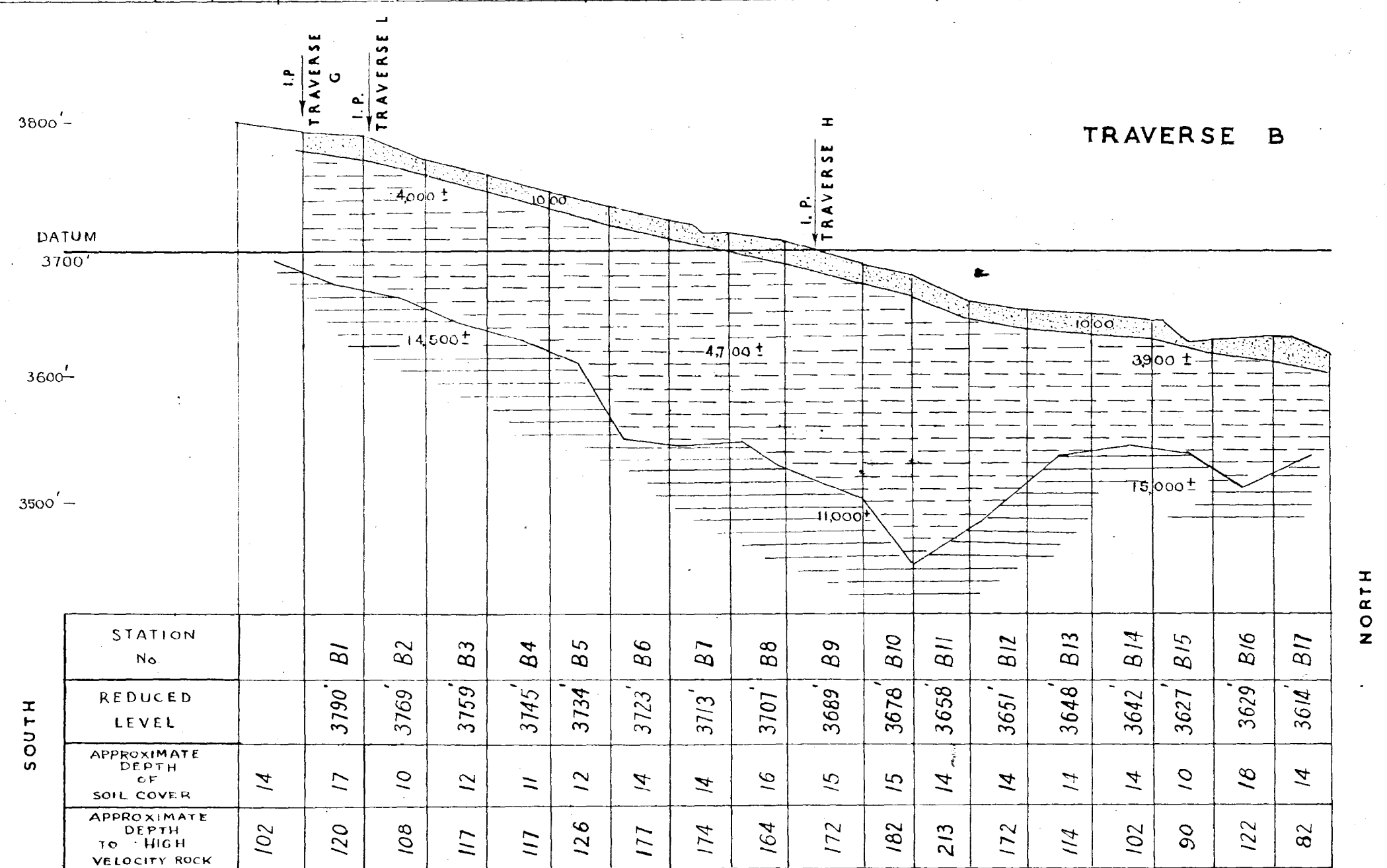
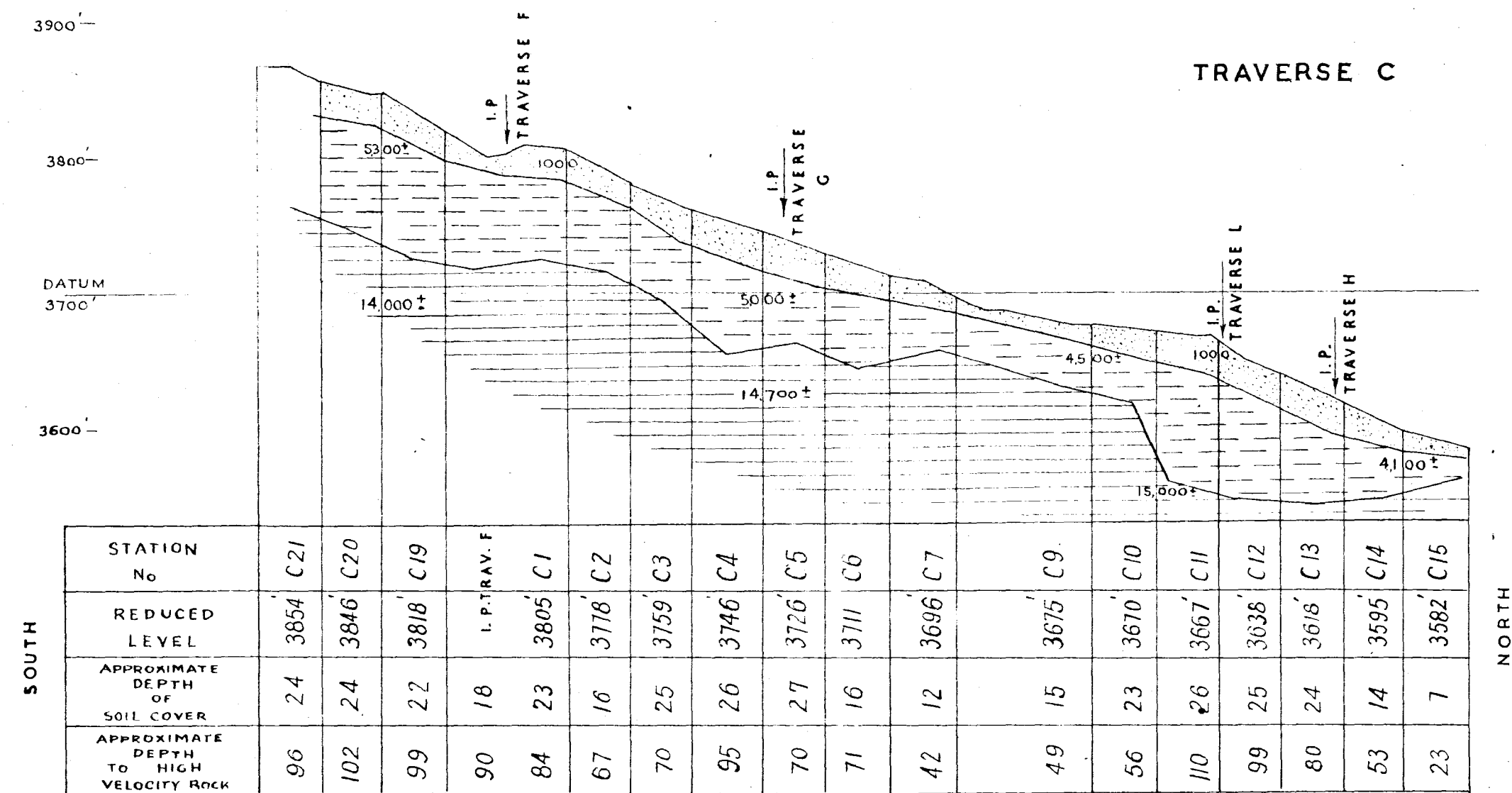
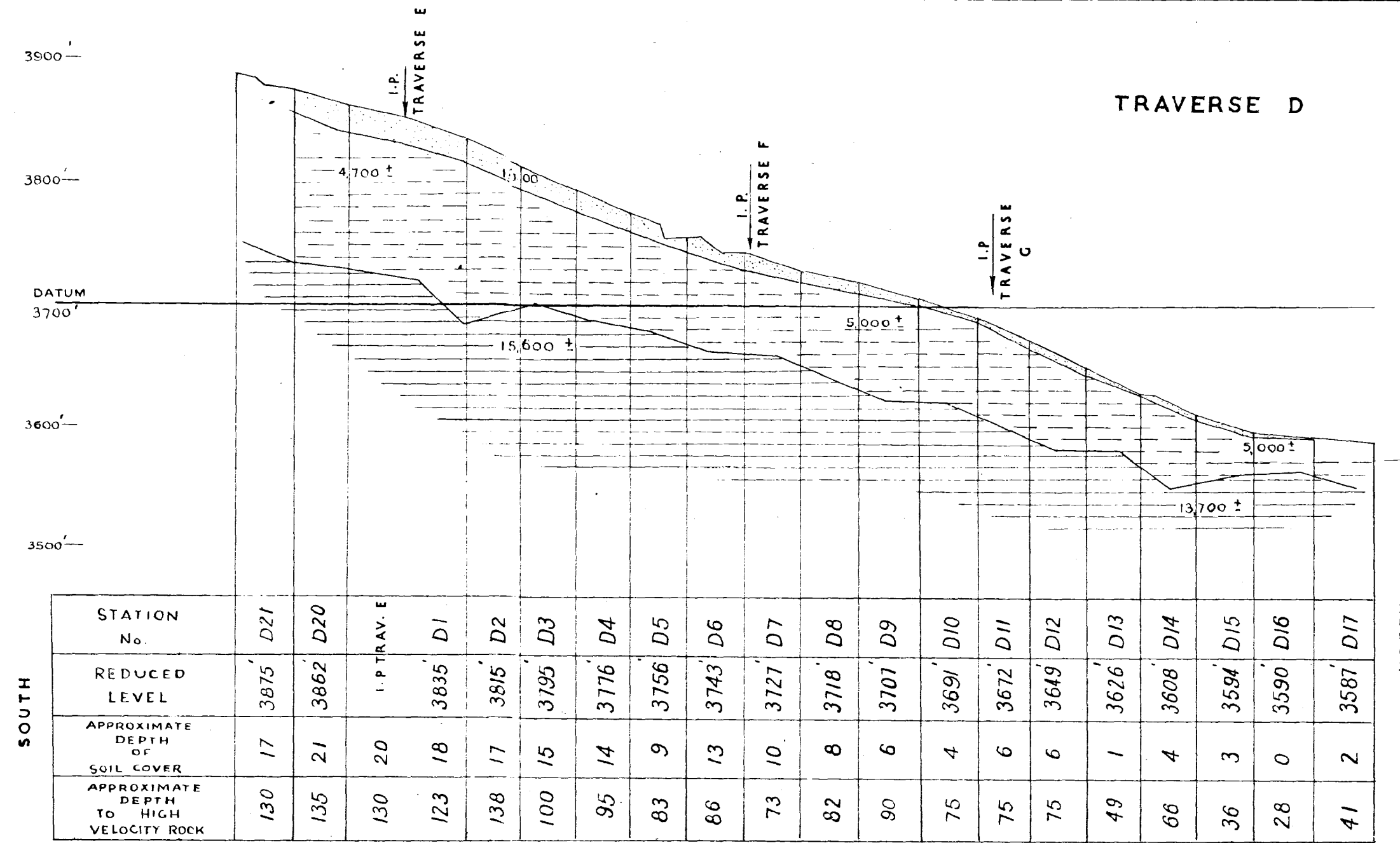
SEISMIC REFRACTION SURVEY, No. 1 POWER STATION SITE, KIEWA, VICTORIA.
REVISED CONTOURS (R.L) OF "HIGH VELOCITY ROCK"

SCALE IN FEET
0 100 200 300

LEGEND

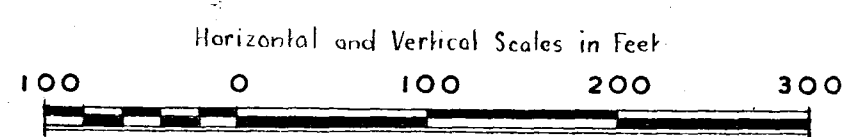
- Seismic traverses
- "High Velocity Rock" contour (Feet above S.E.C datum - 92' below M.S.L.)
- Vertical drill hole
- Inclined " "
- Fault Diorite Dyke Complexes } Beavis (1956)
- Fault

GEOPHYSICIST



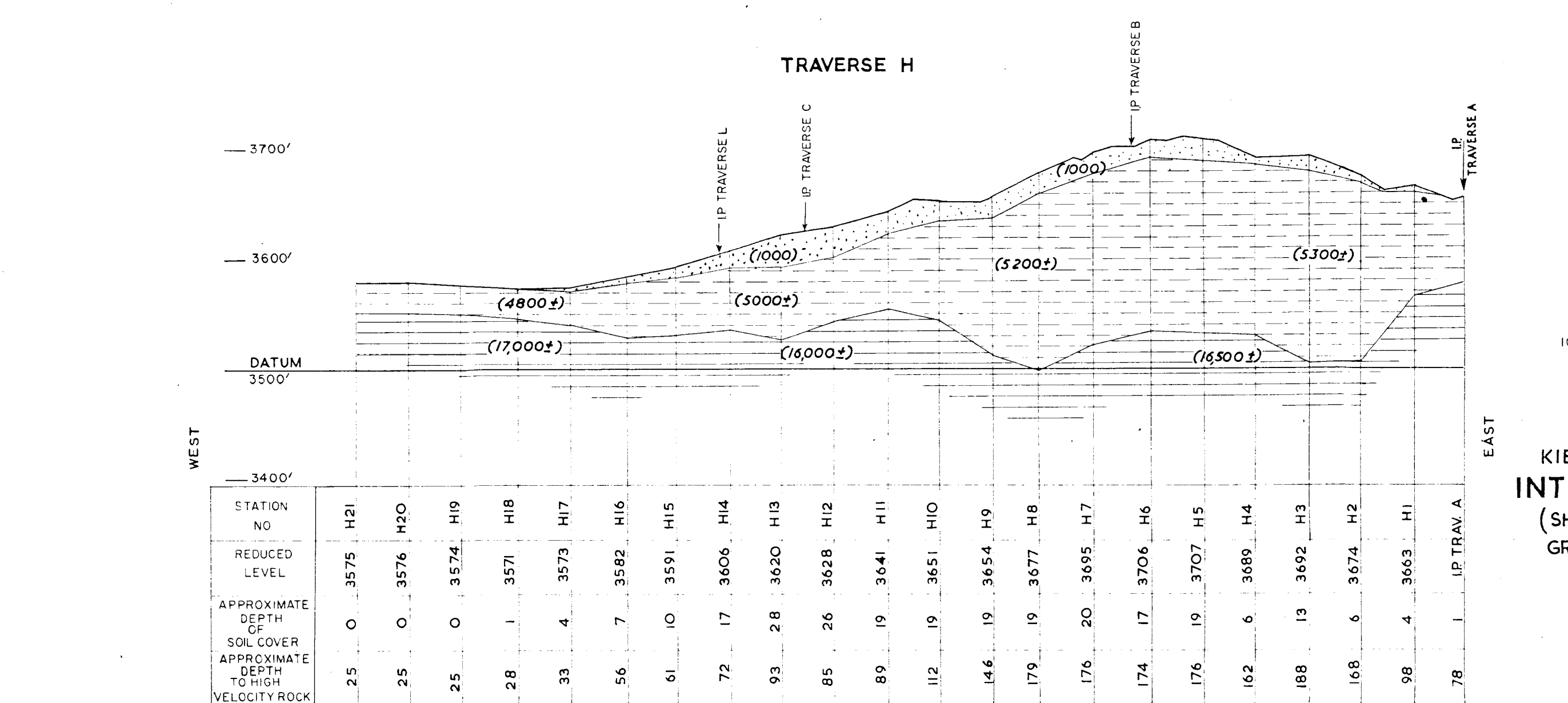
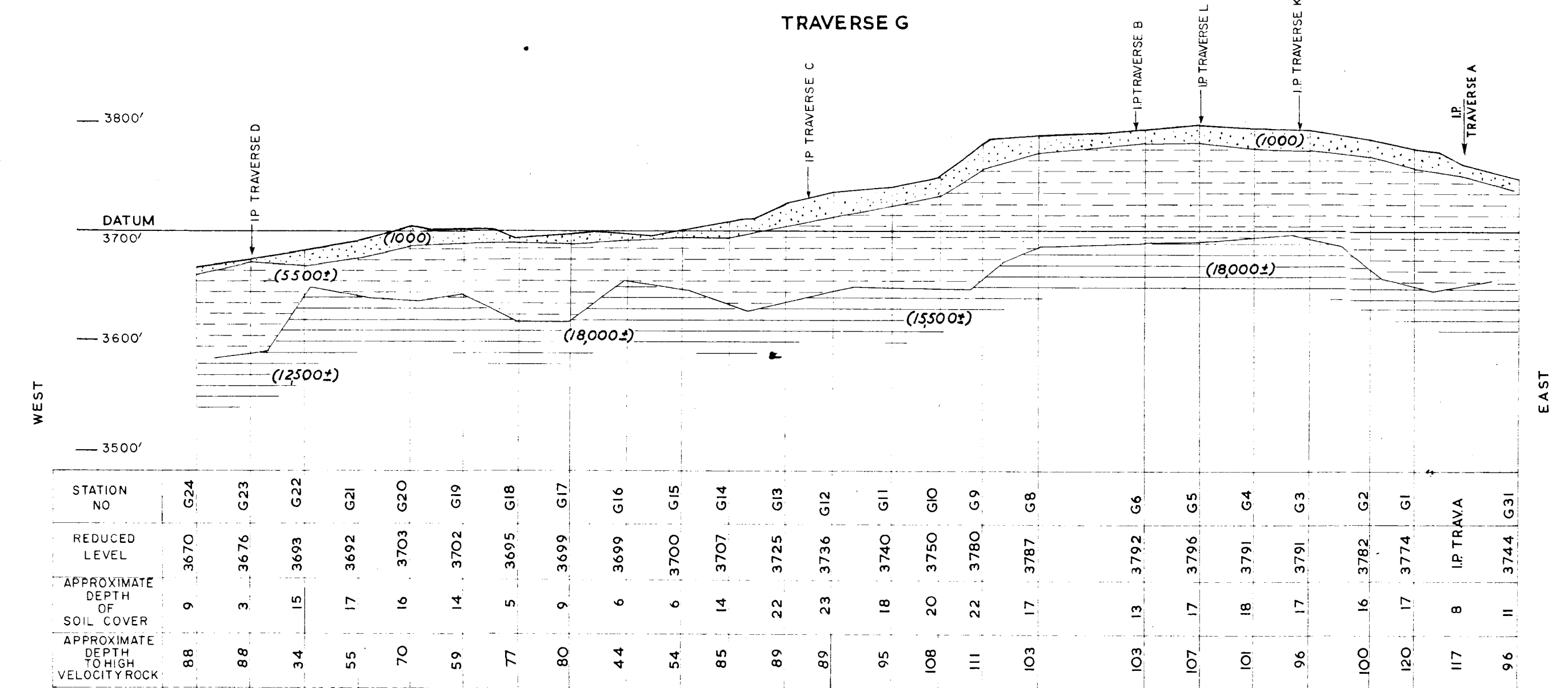
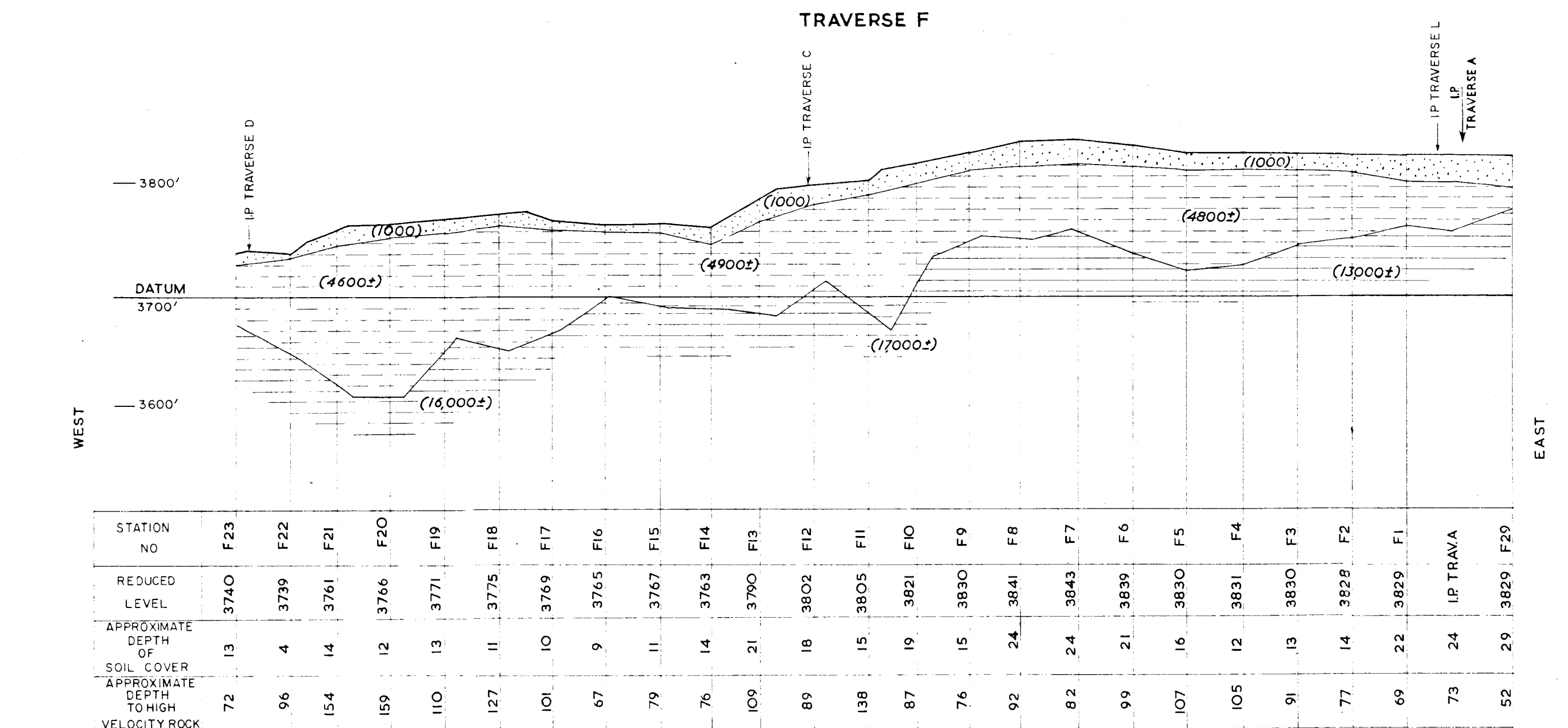
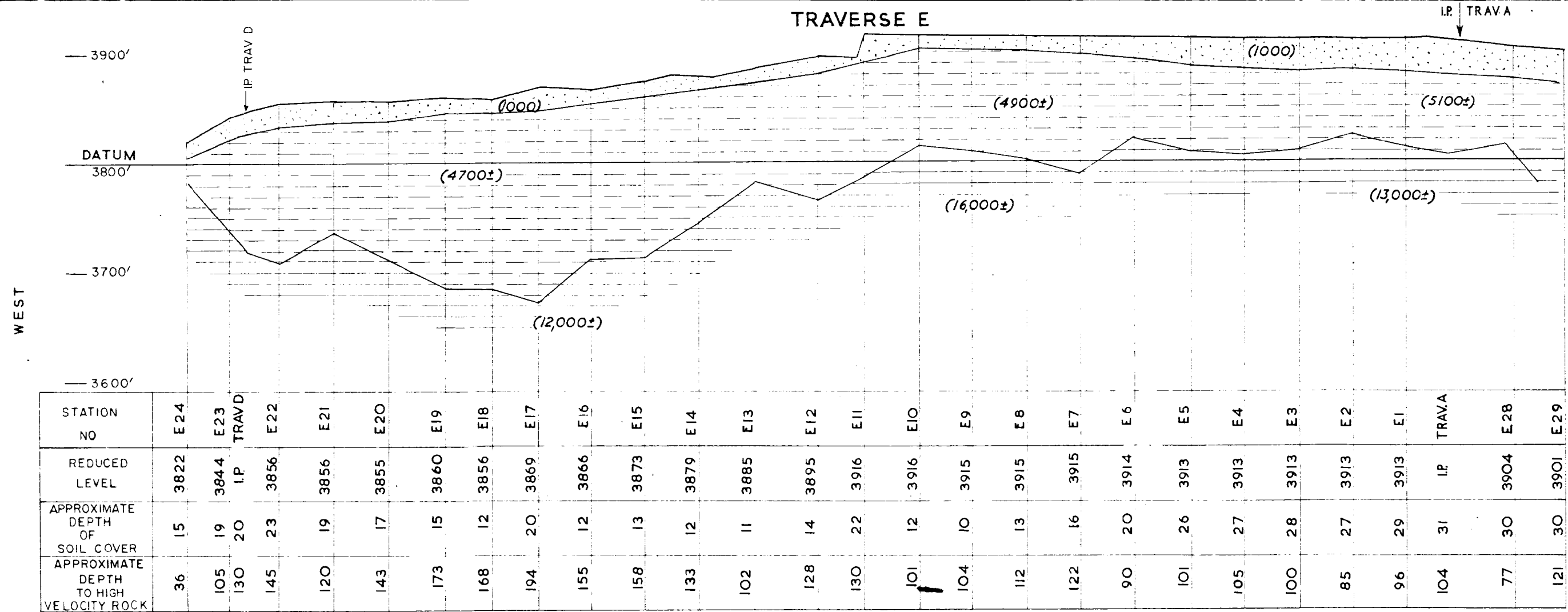
LEGEND

- Soil
- Weathered rock
- High velocity rock
- 4000 ± Approx. seismic velocity in ft/sec
- I.P. Intersection point



SEISMIC REFRACTION SURVEY
OF THE
KIEWA No.1 POWER STATION SITE, S.E.C., VICTORIA
INTERPRETATION OF SEISMIC DATA
(SHOWING THICKNESS OF SOIL & WEATHERED GRANODIORITE
& DEPTH TO UNWEATHERED GRANODIORITE)
TRAVERSES A - D

J. J. Dyson
GEOPHYSICIST



LEGEND

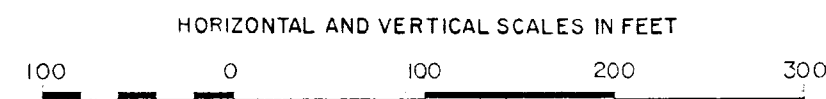
SOIL

WEATHERED ROCK

HIGH VELOCITY ROCK

(4600±) APPROXIMATE SEISMIC VELOCITY, FEET/SEC

IP INTERSECTION POINT

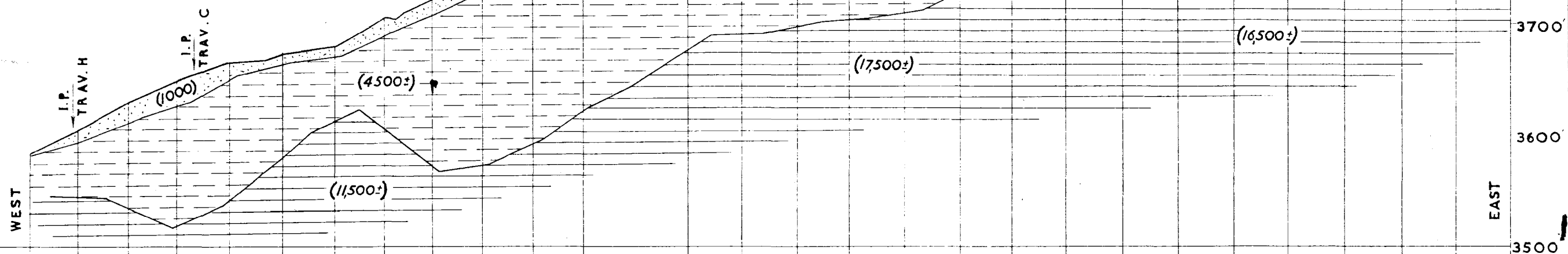


SEISMIC REFRACTION SURVEY OF THE
KIEWA NOI POWER STATION SITE, S.E.C., VICTORIA
INTERPRETATION OF SEISMIC DATA
(SHOWING THICKNESS OF SOIL & WEATHERED
GRANODIORITE & DEPTH TO UNWEATHERED
GRANODIORITE)

TRAVERSES E-F-G-H

J. F. Sykes
GEOPHYSICIST

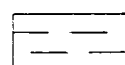
TRAVERSE L

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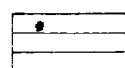
LEGEND



SOIL



WEATHERED ROCK



HIGH VELOCITY ROCK

(4600±) APPROX. SEISMIC VELOCITY IN FT/SEC.

I. P. INTERSECTION POINT



SEISMIC REFRACTION SURVEY OF THE
KIEWA No1 POWER STATION SITE, S.E.C., VICTORIA

INTERPRETATION OF SEISMIC DATA

(SHOWING THICKNESS OF SOIL & WEATHERED GRANODIORITE
& DEPTH TO UNWEATHERED GRANODIORITE)

TRAVERSES L-K

J. F. DeLoe
GEOPHYSICIST

TRAVERSE K

