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AUSTRALIAN ICE THICKNESS MEASUREMENTS
IN ANTARCTICA

BY SEISMIC AND GRAVITY METHODS

1957-59

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ABSTRACT.

Ice thickness measurements carried out by field parties based on Mawson during 1957-59 consisted of:-

- (i) regional traverses in the form of closed loops extending several hundred kilometres inland from Mawson;
- (ii) semi-detailed traverses in the vicinity of a line of ice flow stakes about 25 Km. from Mawson.

The regional traverses showed that, beyond about 175 Km. inland, the area surveyed is influenced strongly by the Lambert Glacier - Amery Ice Shelf system situated some 200 Km. to the East. Preliminary contour plans of the ice and rock surfaces show fairly close correspondence. A sub-glacial extension of a range of mountains outcropping through the ice 80 Km. to the East of the traverses was found.

Work along the semi-detailed traverses close to Mawson detected sub-glacial extensions of the out-cropping mountain ranges in the area. These extensions may explain the general direction of the coastline near Mawson.

1. INTRODUCTION.

During the period of the International Geophysical Year, 1957-58, the Australian National Antarctic Research Expeditions (A.N.A.R.E.) carried out a programme of ice thickness measurements on the Antarctic ice cap. Field parties for this work were based at Mawson station. One geophysicist of the Commonwealth Bureau of Mineral Resources was attached to A.N.A.R.E. during each of the two years covered by the I.G.Y. for the purpose of making these and other measurements.

The work carried out falls under three headings:-

- (i) Regional traverse, 1957-58 field season;
- (ii) Regional traverse, 1958-59 field season;
- (iii) Semi-detailed traverses near Mawson.

2. REGIONAL TRAVERSE, 1957-58.

A preliminary report on the work carried out during the field season 1957-58 was presented at the Moscow meeting of C.S.A.G.I. in August, 1958. Reduction of the results has since been extended and includes additional altitude data obtained during the second Regional Traverse, 1958-59.

The traverse carried out was planned as a single line running due South along meridian 62°E for as great a distance as could be attained, so as to form the basis of closed loops to be surveyed later in selected areas of particular interest found during the traverse. Southerly progress was blocked by heavy crevassing at a distance of 576 km. from Mawson and the party headed South-West from this point in the hope of clearing the outcropping nunataks and the crevassing associated with them, but was forced by impassable surfaces to turn back when a total distance of 642 km. from Mawson had been covered. After another short run of 18 km. from S.P.21 to S.P.23, the party returned to Mawson along the original route. Surplus fuel and explosives were left as a depot at a point 380 km. from Mawson. The traverse is shown on Plate 2. This survey occupied a total of 100 days.

The techniques used may be summarised thus:-

- (i) Seismic reflection stations were established at 32 km. (20 mile) intervals in general, with stations closer together in areas where the rock surface level appeared to be changing rapidly;
- (ii) Explosive charges averaging 450 gm.
 T.N.T. were detonated at the bottom of holes averaging 30 metres deep.
 These holes were drilled with a hydraulic flight auger;
- (iii) 12-channel, wide-pass-band seismic recording equipment was used,

transported in a thermally-insulated, electrically heated cab;

- (iv) Small scale refraction methods were used at most reflection stations to determine the near-surface velocity distribution. Full scale refraction methods (unreversed) were used at two locations to measure wave velocities at depth in the ice;
- (v) Altitudes were measured barometrically at each of the seismic stations and at points generally 8 Km. (5 miles) apart between them:
- (vi) Gravity measurements were made at each point where an altitude measurement was made, using a Worden Geodetic gravity meter.

3. REGIONAL TRAVERSE 1958-59.

The plan for the second regional traverse was to proceed as rapidly as possible to the depot established during the first survey and from there to run traverses in the form of loops approximately 160 km. square to the East and West. Both loops were intended to study sub-glacial extensions of the Prince Charles Mountains; the loop to the West was intended also to study ice surface contours which were believed to rise considerably in this direction, and to find if possible the Westerly limits of ice drainage into the Lambert Glacier system.

Soon after leaving Mawson the party ran into difficulties with a belt of crevasses which led to a delay of 16 days while tractors were recovered from crevasses and repairs made to equipment. At the same time, the ice drill was so seriously damaged that only two further holes could be drilled after this incident. For the remainder of the seismic work air shooting techniques were used. During travel to the 380 km. depot more fuel was consumed than had been anticipated and this fact, together with the loss of time experienced, led to the abandonment of the loop to the East.

Measurements commenced when the party set off West from the depot, and, as had been expected, altitude in this direction was found to increase rapidly. After only 18 km. had been covered the tractors, which were losing power due to the increase in altitude, became bogged down in soft snow. The party then turned North for 43 km. when a slightly improved surface was found and the party again turned West. After a further 26 km. progress the party turned North and ran parallel to the original traverse for 163 km. and some 50 km. West of it. Course was then altered to intersect the first traverse at a point 240 km. North of the depot camp. The return to Mawson was made along a different route from that followed by the first party, producing two further small loops.

During this second survey, air shooting was used at each seismic reflection station except the first two, where the drill could still be used. For air shooting a square pattern of nine charges with 9 metres (30 feet) separation was used.

The charges were each 450 gm. T.N.T., suspended on bamboo poles 1.5 to 1.8 metres above the surface. A single linear spread of 12 geophones was used with 30.5 metres (100 feet) separation between them, and the pattern of charges was generally offset 610 metres (2000 feet) from the centre of the spread. Results from this method of air shooting were reasonably good, but generally inferior to those from shot holes.

Apart from the enforced use of air shooting, the equipment and techniques used in the seismic work were the same as had been used during the 1957-58 field season. No refraction measurements were made.

Gravity measurements were made in the same way as previously but the interval between gravity stations was reduced from 8 Km. to 4.8 Km. (3 miles). Differential altitude measurements were also made at 4.8 Km. intervals, using Fuess aneroid barometers.

The traverse carried out is shown on Plate 2, together with the 1957-58 traverse. The 1958-59 survey occupied a total of 110 days.

4. SEMI-DETAILED TRAVERSES NEAR MAWSON:

This project was carried out during the Antumn of 1958. Referring to Plate 5, it will be seen that some 25 km. inland from Mawson the Frammes Mountains outcrop through the ice. There are four principle ranges, the Henderson, Masson, David and Casey Ranges, and, as they appear above the ice, they appear to be aligned roughly parallel to the presumed Northerly direction of ice flow.

Early in 1957 a line of stakes was established and surveyed to the North of these ranges with the purpose of measuring the surface rate of ice flow. The positions of the stakes were fixed relative to the rock outcrops (Mellor, 1959). Flow rate measurements are continuing.

To establish the sub-glacial rock profile across the line of stakes and to study its relation to the surface rate of flow, a seismic and gravimetric traverse was run along this line, supplemented by gravity measurements along two of the valleys, approximately perpendicular to the main traverse.

Shot holes were used for the seismic work on these traverses. Explosive charges ranging from 100 gm. to 450 gm. were shot at depths between 10m and 30m. Split linear geophone spreads were used in most cases, with the interval between geophones 17 metres and an offset of 61 metres from the shot hole to the nearest geophone on each side. Considerable interference from surface noise was experienced.

Levelling of seismic stations was carried out by the odolite, by observations of mountain peaks of known position and height. Gravity stations were levelled by the differential barometric method using the seismic stations as control points.

5. RESULTS FROM THE REGIONAL TRAVERSES.

Considerably more altitude information was obtained

by the second field party to supplement and improve the results previously given for the first traverse. In particular, during a six day halt at the depot camp for maintenance and repairs on the transport equipment, detailed records were made of atmospheric pressure variations. These were found to correlate well with records kept at Mawson over the same period and these results, together with differential levelling carried out during the traverses, have been used to derive a revised value for the altitude of the depot camp. Other altitudes have been adjusted to fit this value and to eliminate misclosures around the loops.

The altitude results are presented in the form of a "Preliminary Ice Surface Contour Map" which is reproduced as Plate 3. In mapping the contours, altitude values obtained during journeys to the Prince Charles Mountains in earlier years were also used.

It will be seen that, close to the coast, the direction of greatest surface slope is to the North. Proceeding inland, the contours swing round as the direction of greatest slope acquires a progressively greater Easterly component. At a distance of 300 km. from the coast the slope is to the East. The swing round towards the South continues approximately another 170 km. At this distance the direction of greatest slope is approximately South-South-West. Towards the most Southerly point reached the contours indicate the direction of slope as approximately East-North-East. The highest part of the area covered is at Latitude 70°07' South, Longitude 60°50' East, 295 km. from the coast and approximately 250 km. West of the edge of the Amery Ice Shelf. The altitude here was 2570 metres and was increasing to the West.

These results show that the depression of the ice surface associated with the Lambert Glacier-Amery Ice shelf system extends for at least 250 Km. and probably much more, to the West of the system.

The results of the ice thickness measurements made during the regional traverses are tabulated in Table 1. Reflection times have been corrected for delays due to low velocity layers near the surface. The velocities used are those derived from the refraction measurements viz.:-

SP4 - 3810 metres/sec.

SP5 - 3825 " SP6 - 3840 "

SP7 and all points to the South - 3860 metres/sec.

Depths derived from gravity measurements were determined using measured Bouguer anomalies to interpolate between pairs of seismic reflection stations and still are subject to minor correction on revision.

Using values for the altitude of the rock surface under the ice, obtained from the ice thickness and altitude results, a "preliminary Rock surface contour map" has been drawn (plate 4). The rock surface is considerably more irregular than the ice surface and the contours as drawn through the observed points can do no more than show the broad trends.

The most prominent feature of this contour map is the submerged range at SP12. Unfortunately reflection quality here was poor (as might be expected from the convex surface presented by a rock altitude maximum) and some doubt exists as to the exact height of the rock surface here. However, if the seismic result from this station is ignored and gravity interpolation only is used between SP11 and SP13 a somewhat lower rock "high" still results. This almost certainly represents a sub-glacial extension of the spur of the Prince Charles Mountains which appears through the ice 80 Km. to the East at the "Riddell Nunataks"; which have a maximum altitude of about 2300 metres.

It will be noted that over considerable areas of the map the rock surface appears below sea level; there is a tendency for the areas below sea level to occur predominantly on the Western side of the area traversed, but nevertheless it seems likely that to the East, in the area of the Lambert Glacier, most of the surface would fall below sea level to produce the "sink" into which the ice flows to form the Glacier. It is unfortunate that no measurements could be obtained in this area.

It will be noted that some of the trends exhibited by the ice surface contours appear to follow trends in the rock surface contours, although the rock surface contours are considerably more irregular. Both sets of contours run generally East-West near the coast. The ice surface contours commence to swing round towards a North-Vest to South-East direction some 90 km. inland, at which point the rock surface is very irregular, with some areas more than 500 metres below sea level. A very prominent belt of ice domes and depressions running East-West was crossed here. Between this region and SP12, a distance of about 175 km., both sets of contours have a pronounced North-West to South-East trend. South of SP12, as far as the depot camp at 380 km. there appears to be little correlation between the rock and ice surface contours. Further South still, the information is too sketchy to attempt any comparison.

The general picture which emerges is that ice flow from the South is largely diverted to the West by the range of nunataks of which Mt. Menzies is the most prominent peak. To the North of this region the ice acquires an Easterly component of flow towards the Lambert Glacier system and tends to flow North-East, but further damming of ice flow occurs at the submerged range crossed near SP12 and part of the ice is diverted towards the South-East. North of this submerged range ice flow is broadly at right angles to the rock surface contours.

6. RESULTS FROM THE SEMI-DETAILED TRAVERSES:

The results are presented as profiles in Plate 6. Two rock valleys are indicated on the profile of Traverse A, one centred at Station SP32 the other at Station G6. These valleys reach a depth of 90 metres below sea level, and are separated by ridges rising some 60 metres above sea level. This relief is very gentle in comparison with the ranges appearing through the ice (which reach an altitude of over 1000 metres), but comparison with the station location diagram (plate 5) shows that the ridges and valleys do not occur where they would be expected to if the alignments of the ranges were produced on to traverse A. For example, the centre of the Easterly rock valley is due North of the tip of

the David Range. It appears that the trends are in a roughly North East direction, so that for example the Southerly portion of the Masson Range and Mt. Henderson are connected by a sub-glacial ridge. A similar trend is seen in the islands lying close to the coast near Mawson, in the ice coastline itself and in the ice surface contours in the area between Mt. Henderson and Mawson. Thus it appears likely that ice flow in the region of the flow stake line is predominantly to the North-West and not approximately at right angles to the stake line as the mountain alignments would suggest. Flow rate measurements are presented in a paper by I. McLeod.

7. REFERENCES:

GOODSPEED M.J. 1958 - Preliminary Report. Commonwealth Bur. Kin. Res. Geol. and Geoph., Records 1958/40.

JESSON E.E., 1959 - Commonwealth Bur.Min.Res., Geol. and Geoph., Records 1959/74.

MATHER K.B. and
GOODSPEED M.J. 1959 - Thickness measurements and sastrugi observations,
MacRobertson Land, 1957-58,
Polar Record, Volv9, 62.

MELLOR M., 1959 - Ice flow in Antarctica, Jour. Glaciology, Vol.3, 25.

TABLE I.

RECIONAL TRAVERSES ICH TUICKETSS L'EASUREMENTS.

Stotion.	Latitude o S.	Longitude OE.	Ice Thickness Metros.	Mathod.
CDA	4-9401			
SP3 GG4	67 ⁰ 58°•0 67 ⁰ 56°	62 ° 30' •0	528	Seismic.
GG5	68 ° 00•	62°22°	625	Gravity.
G06	68 ° 33 °	52 6 13•	739	69
SP4	68934	62011.	1869	n
	68°36°.7	62507 .0	943	Seismic.
GG7 GG8	60°10°	62 ° 08•	1209	Gravity.
GG9	60 014 •	62°03°	1 00 5	10
GG1 0	68°18°	62°09•	1476	19
SP5	68 0 24 . 7	02009	3.705'	16
GG11	69,58	52°53°.7	1607	Seimic.
SPS	68°32°•7	68 0 08•	1409	Gravity.
GC12	68°32° • 7	62°06° •3	1668	Seismic.
SP7	68 °41° •3	62007	1516	Gravity.
GG13	68044	62°08' •0 62°07'	2403	Seismic.
CC14	68°48°	62°07°	2191	Gravity.
G C1 5	68°52°	62 ° 06 °	2055	19
GG16	68 0 55•	62°05•	1799	11
SP8	68059 .1	62004 .6	1633 1607	
GG17	69003	62006•	1905	Seismic.
GG18	69008	62°08°	1648	Grevity.
SP9	69°16° -3	62°10°.2	1387	**
SP10	69°83° 8	ິລ ^າ າປີ. 2		Seismic.
GC19	69025	82010.	1284	
GG20	69029•	62 0 10•	1450 11 90	Gravity.
G021	69039•	68010	1232	**
SP11	69°37° - 8	62°30°•8	155 5	
GG22	69°48°	62°03°	139 3	Seismic.
0023	69 ⁰ 46•	62°06°	974	Gravity.
GG24	69051	62005*	362	6
SP12	69 ⁰ 55•	63003 • 0	256	Seismic.
0025	70000	62005	7 52	Gravity.
G G26	70004	62°07°	779	01 G 12 Cy 6
GP13	70013.4	62°08°6	3327	Soismic.
GG27	70018	68 ° 09•	8259	Gravity.
CG28	70022	63 0 33•	2460	(1
1029	70027	68009	2425	10
5P14	70 ⁰ 34	62 <mark>9</mark> 09•.2	2655	Seismic.
11104	79036•	68°09	2151	Gravity.
1031 5632	70°40° 70°45°	62 0 09•	1614	19
5032 5 P2 5	70049 • 6	62°08°	1942	€
1633	70054	62 ° 08 ° . 0 62 ° 08 °	824 4	Seismic.
1024	72050	62 ⁰ 08•	2295	Gravity.
IG35	71°03°	62 0 37•	2164	#
P15	71008 .1	62°06°.9	2372	O
G36	71018	62°07°	2401	Seismic.
G3 7	71017.	62°07°	2400	Cravity.
P16	71021.6	62°07° •0	2145	0
G38	710260	62 007 •	2095	Scienic.
G39	71°30°	62°07•	2063	Gravity.
G40	71034.	62 037° •2	2091	n
G 41	71°58•	62 ⁰ 07'•3	1916	Ð
G42	71042	62 ⁰ 07* •4	2 134	48

Station. Latitude os. SP17 71°45' GG43 71°50' GG44 71°55' SP24 71°59' GG45 72°04'	о _{Е•}	Ice Thickness Metres. 2190 1926 2237 2495	Method Seismic Gravity
CC/13 71 501	.8 62°07'.5 62°07'.5 62°07'.6 62°07'.6 62°08'	1926 2237	
GG46 SP18 72°12' GG47 72°17' SP19 72°21' GG48 72°26' GG49 72°31' SP20 72°35' GG50 72°37' GG51 72°39' GG52 72°40' SP21 72°41' GG53 72°44' GG54 72°47' GG55 72°49' SP22 72°52' GG56 72°48' SP23 72°51' G43 70°50' G44 70°50'	62 08' 62 07' 8 62 08' 62 08' 62 08' 62 08' 62 08' 62 08' 61 54' 61 40' 61 26' 8 61 12' 2 61 01' 60 49' 60 38' 1 61 21' 0 61 55' 61 43'	2441 2275 2045 1935 1983 1417 1798 1757 1099 1007 1477 1991 1870 2137 1834 1475 1899 1297 2436 2699 2732	Seismic Gravity Seismic Gravity Seismic Gravity Seismic Gravity " Seismic Gravity " Seismic Gravity
G48 G49 G49 G47 G50 G47 G50 G61 G63 G63 G64 G65 G60 G61 G62 G63 G64 G63 G64 G65 G64 G65 G64 G65 G66 G67 G66 G67 G66 G67 G68 G69	61° 35' 61° 34' 61° 34' 61° 34' 61° 34' 61° 34' 61° 34' 61° 34' 61° 34' 61° 35' 61° 27' 61° 20' 61° 12' 61° 58'	2738 2778 2565 2414 2322 2287 2267 2315 2468 2535 2423 2395 2512 2557 2514 2718 2718 2718 2718 2718 2718 2718 2718	Seismic Gravity "" "" "" "" "" "" "" "" "" "" "" "" "

Station	Latitude	Longitude	Ice Thickness Metres.	Method
G884567898 G884567899900123456789012304567890123045679012333333333333333333333333333333333333	7 15 2 6 6 6 6 6 6 6 6 6 6 6 6 6	3 000000000000000000000000000000000000	221860 1988 1979 1979 1979 1979 1979 1986 1990 1980 1990 1990 1990 1990 1990 1990	Gravity Seismic Gravity Seismic Gravity """ """ """ """ """ """ """











