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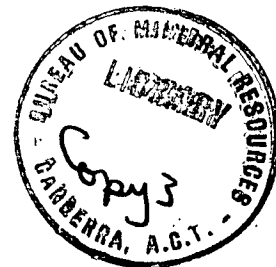
COMMONWEALTH OF AUSTRALIA

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DEPARTMENT OF NATIONAL DEVELOPMENT

BUREAU OF MINERAL RESOURCES, GEOLOGY AND GEOPHYSICS

RECORD No. 1962/15



WALTERS MARSH DAM SITES SEISMIC REFRACTION SURVEY, TASMANIA 1961

by

E.J. Polak

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SUMMARY

This Record describes a seismic refraction survey at the Walters Marsh dam sites on the Mersey River, requested by the Hydro-Electric Commission of Tasmania.

The 'overburden' consists of large boulders of dolerite or quartzite with pebbles, gravel, sand, and clay. Seismic velocities in the overburden range between 3000 ft/sec on the slopes above the water-table, to 7000 ft/sec on dam site No. 1, where boulders up to 2 ft in diameter were proved by drilling. The 'bedrock' consists of quartzite with seismic velocities ranging from 10,000 ft/sec in a shear zone to 17,000 ft/sec for fresh solid quartzite. Young's modulus of the bedrock ranges between 4.2×10^6 and 6.5×10^6 lb/in.².

1. INTRODUCTION

The Hydro-Electric Commission of Tasmania proposes to construct a dam on the Mersey River, approximately 20 miles south of Mole Creek (approximate co-ordinates E.420 N.862 on the Devonport four-mile map). The dam will store water to be used in the power stations located in the lower reaches of the Mersey and Forth rivers (Polak and Duggin, 1961). Three possible dam sites were chosen for investigation.

The Commission requested the Bureau of Mineral Resources, Geology and Geophysics to assist with the investigation, the problem being the determination of the depth to bedrock and the nature of both the 'overburden' and the bedrock.

The seismic survey was carried out between 13th and 24th February 1961. The geophysical party consisted of E.J. Polak (party leader), D.J. Harwood (geophysicist) and J.P. Pigott (geophysical assistant). The Commission provided additional assistants and did the topographical surveying of traverse lines.

The total lengths of the surveyed traverses were:

Dam site No. 1	3650 ft
Dam site No. 2	4350 ft
Dam site No. 3	5050 ft
Total	<u>13,050 ft</u>

2. GEOLOGY

The geology of the area has been described by Spry (1958) and is illustrated on Plate 1. A detailed geological examination of the dam sites was made by Paterson (1959). His results are shown on Plates 2, 4, and 6. Four holes had been drilled on site No. 1 (see Plate 2).

The three dam sites are in a narrow valley of the Mersey River. The floor of the valley is filled in with fluvioglacial deposits containing boulders, pebbles, sand, silt, and clay. The boulders are predominantly of fresh dolerite, with a few of quartzite, both ranging up to two feet in diameter.

The bedrock in the valley consists of Precambrian quartzite, mica-quartzite, and mica schist of the Howell Group. The rocks are jointed, predominantly along the foliation. Drilling on site No. 1 indicated that some of the joints are open and filled in with clay or limonite. The beds are dipping steeply northwards.

As used in this Record, the term 'bedrock' refers to unweathered rock, including jointed rock, in the refractor with the highest recorded seismic velocity. The term 'overburden' refers to river gravel, clay, scree material, and completely or partly weathered rock, all with seismic velocity less than 9000 ft/sec.

3. METHODS AND EQUIPMENT

The seismic method of exploration depends on the contrast between the velocities of the seismic waves through different rock formations. Hard unweathered rocks have higher velocities than their weathered counterparts, and these in turn have higher velocities than unconsolidated deposits. A detailed description of the method has been given in previous reports to the Commission (e.g. Polak and Moss, 1959).

The 'method of differences' was used for traverses along the slope, but 'broadside spreads' were used on steep banks.

The equipment used in this survey was a Midwestern 12-channel reflection-refraction seismograph with T.I.C. geophones of natural frequency about 20 c/s. A three-component geophone was used to record longitudinal and transverse waves.

4. RESULTS

Seismic velocities

The seismic velocities recorded may be grouped as follows:

Low velocities in surface layer : 1000 to 2000 ft/sec. These velocities are recorded in dry, unconsolidated soil and sand, above the water table.

Medium-velocity layers: 3000 to 5000 ft/sec. On sites No. 2 and 3 the Commission geologist (Paterson, 1959) indicated river terrace material (see Plates 4 and 6). These velocities indicate sand, gravel, and clay of the river terrace material; the lower velocity may indicate a higher content of clay material. A velocity of 4000 ft/sec recorded on the slopes high above the water-level probably indicates very weathered rock.

Velocities 5000 to 7000 ft/sec: Drilling on site No. 1 indicates that the overburden there consists of large boulders (up to two feet in diameter) of fresh dolerite and quartzite. These rocks are characterized by very high velocities (15,000 to 18,000 ft/sec); therefore unconsolidated deposits consisting of large boulders will show fairly high seismic velocity, especially below the water-table where all the space between boulders is filled with water-saturated material. The velocity of 5000 to 7000 ft/sec is therefore interpreted as indicating fluvioglacial deposits containing a high proportion of large boulders. On Traverse H, site No. 3 (Plate 7), a velocity of 9000 ft/sec was recorded; this velocity has been interpreted as representing weathered bedrock. This conclusion has been reached because the bed shows seismic velocity anisotropy (the velocity measured along the strike was 9000 ft/sec, whereas the velocity recorded at right angles to the strike was only 7000 ft/sec), and because no velocity higher than 7000 ft/sec has been recorded on unconsolidated deposits in other parts of Tasmania.

High-velocity layers: 10,000 to 17,000 ft/sec. These velocities were recorded in the deepest refractor, therefore they represent the seismic velocities in the bedrock. The lower velocities represent the sheared and fractured material, and the higher velocities represent solid massive rock.

Anisotropy of seismic velocities was recorded at all three sites, the east-west velocity (in the direction of strike) being generally higher than the north-south velocity. The seismic velocities recorded in the bedrock are shown on Plates 2, 4, and 6.

Dynamic properties of rocks

If transverse seismic velocities are measured in addition to the longitudinal velocities, the dynamic properties of the rocks can be calculated (Polak & Moss, 1959). The following table shows the locations, velocity data, and elastic properties of the bedrock. The velocities were measured using a 3-component geophone.

<u>Dam Site</u>	<u>Location</u>	<u>Longitudinal Velocity</u> (ft/sec)	<u>Young's modulus</u> (lb/in. ²)	<u>Poisson's ratio</u>
2	D3	15,000	6.5×10^6	0.29
2	D29	12,000	4.2×10^6	0.28
3	J9	15,000	6.5×10^6	0.285
3	K9	14,000	5.7×10^6	0.285
3	L9	13,000	4.8×10^6	0.285

The results indicate:

- (a) Poisson's ratio ranges between 0.28 and 0.29
- (b) Young's modulus ranges between 4.2×10^6 and 6.5×10^6 lb/in.².

Both these values are higher than those calculated for the Martha Creek dam site (Polak and Duggin, 1961) which is approximately 7 miles down-stream.

Dam site No. 1

Plate 2 shows the detailed geology of the area and the locations of seismic traverses; Plate 3 shows the interpretation of the data in the form of seismic profiles.

Site No. 1 is located farthest up-stream. The valley here is very broad, but a knoll of solid rock in the centre would facilitate the construction. The present bed of the river is solid rock.

In the westerly part of the area the valley is filled with fluvioglacial deposits. The thickness of the unconsolidated rocks there reaches 133 ft in drill-hole 5707.

In a deep narrow channel, like the one on site No. 1, the first arrival refracted waves would come from the steep sides of the valley. Thus the vertical depth to the bottom of the valley would be considerably greater than the distance to the refractor calculated from the records. This applies on Traverse A, between stations 11 and 14.

The following table shows a comparison of drilling and seismic data.

<u>DDH No.</u>	<u>Depth to Bedrock</u>		<u>Difference</u>	<u>Distance, DDH to seismic station.</u>
	<u>from drilling</u>	<u>from seismic survey</u>		
5707	133 ft	65 ft	68 ft	20 ft
5706	57 ft	54 ft	3 ft	5 ft
5703	99 ft	73 ft	26 ft	10 ft

Seismic velocities recorded in the bedrock are shown on Plate 2. An area of low velocities, suggesting the presence of a shear zone, crosses the traverses approximately at C3, A12, and P10. The shear zone was presumably deeply eroded during the last glaciation.

Dam site No. 2

Plate 4 shows the detailed geology of the area and the location of seismic traverses. Plate 5 shows the interpretation of the data in the form of seismic profiles.

Site No. 2 is located approximately $\frac{1}{3}$ of a mile down-stream from site No. 1. The river valley narrows here and it is covered with the river terrace. No drilling has been done on the site.

The seismic survey indicated that the river follows the centre of a U-shaped buried valley. The thickest overburden (about 92 ft) was indicated near D10, 50 ft west of the existing river. The velocities in the overburden are slightly lower than those at site No. 1, suggesting finer material.

An area of low velocities, suggesting the presence of a shear zone, crosses the traverses.

Dam site No. 3

Plate 6 shows the detailed geology of the area and the location of seismic traverses. Plate 7 shows the interpretation of the data in the form of seismic profiles.

Site No. 3 is located approximately $\frac{1}{2}$ mile down-stream from site No. 2. The river flows along a steep valley filled with fluvioglacial deposits in clearly defined river terraces. No drilling has been done on this site.

The seismic survey indicated that the thickest overburden (81 ft) is near H17. The velocities in the overburden there are lower than on the upper sites, indicating finer material. At the base of the overburden a bed with a velocity of 9000 ft/sec suggests the presence of weathered bedrock.

As on sites No. 1 and 2, a low-velocity zone between H18 and H22 may suggest the presence of a shear zone with a northerly strike in the centre of the valley.

5. CONCLUSIONS

The geophysical survey provided information on the thickness of overburden and the nature of both overburden and bedrock.

On each site a low-velocity zone was recorded in the centre of the valley. This zone suggests a shear zone eroded by ice and water, leaving deep and narrow valleys.

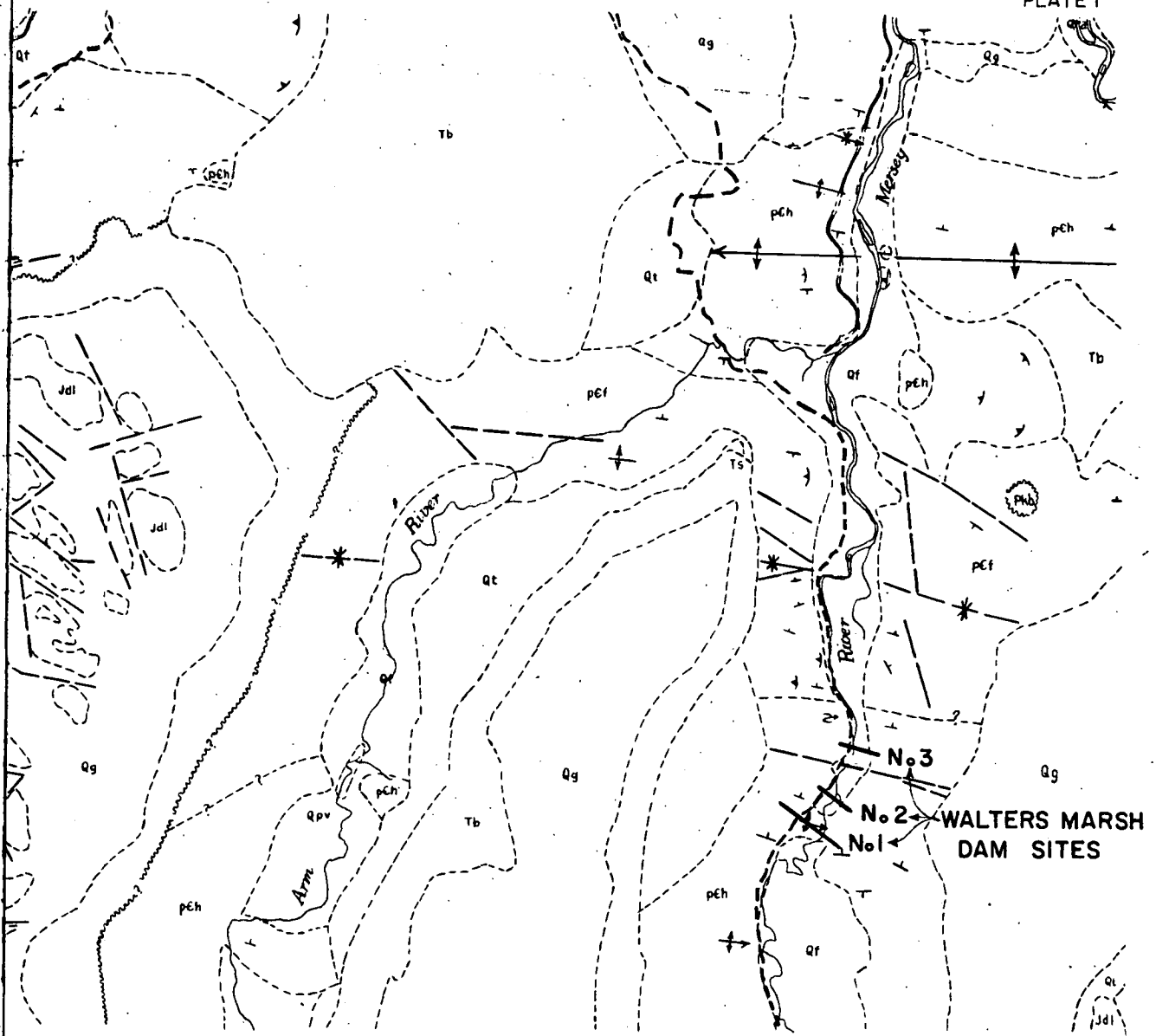
The difference in seismic velocities of the overburden for different sites indicates that overburden on site No. 1 contains coarser material than the overburden on the two lower sites.

The value of Poisson's ratio for bedrock is about 0.28. Young's modulus ranges between about 4.2×10^6 and 6.5×10^6 lb/in.².

Considering that the seismic work suggests the presence of shear zones, further drilling is recommended at: Site No. 1, P10; Site No. 2, D16; and Site No. 3, H17 and H21.

6. REFERENCES

- | | | |
|------------------------------|------|---|
| PATERSON, S.J. | 1959 | Unpublished maps, Hydro-Electric Commission, Tasmania. |
| POLAK, E.J. & DUGGIN, M.J.W. | 1961 | Mersey-Forth power scheme geophysical surveys, Tasmania 1960. <u>Bur. Min. Resour. Aust. Rec. 1961/29.</u> |
| POLAK, E.J. & MOSS, F.J. | 1959 | Geophysical survey at the Cluny dam site, Derwent River, Tasmania. <u>Bur. Min. Resour. Aust. Rec. 1959/87.</u> |
| SPRY, A.M. | 1958 | Precambrian rocks of Tasmania; Part 3 - Mersey-Forth area. <u>Pap. Roy. Soc. Tasm. 92, 117-137.</u> |



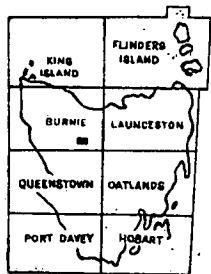
LEGEND

- ESTABLISHED BOUNDARY
- UNCONFORMITY
- GEOLOGICAL BOUNDARY (position approximate)
- STRIKE AND DIP OF STRATA
- TRANSCURRENT FAULT
- DIRECTION OF PLUNGE OF LINEATION
- LANDSLIPS SHOWING HEEL OF SLIP AND DIRECTION OF MOVEMENT
- RECENT ALLUVIUM Qra
- VARVED CLAYS Qpv
- KANSAS CREEK BEDS Pkp
- FLUVIOGLACIAL DEPOSITS, TILL and ALLUVIUM Qt

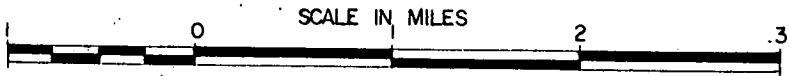
- ROAD
- STRIKE AND PLUNGE OF DRAGFOLDS
- ANTICLINAL AXIS (position approximate)
- ANTICLINAL AXIS
- SYNCLINAL AXIS (position approximate)
- SYNCLINAL AXIS

- BASALT, SCREE MATERIAL and LANDSLIP DEBRIS
- TILL, SCREE MATERIAL, MARSH DEPOSITS and GRAVELS
- HOWELL GROUP Pch
- TERTIARY BASALT Tb
- JURASSIC DOLERITE Jdl

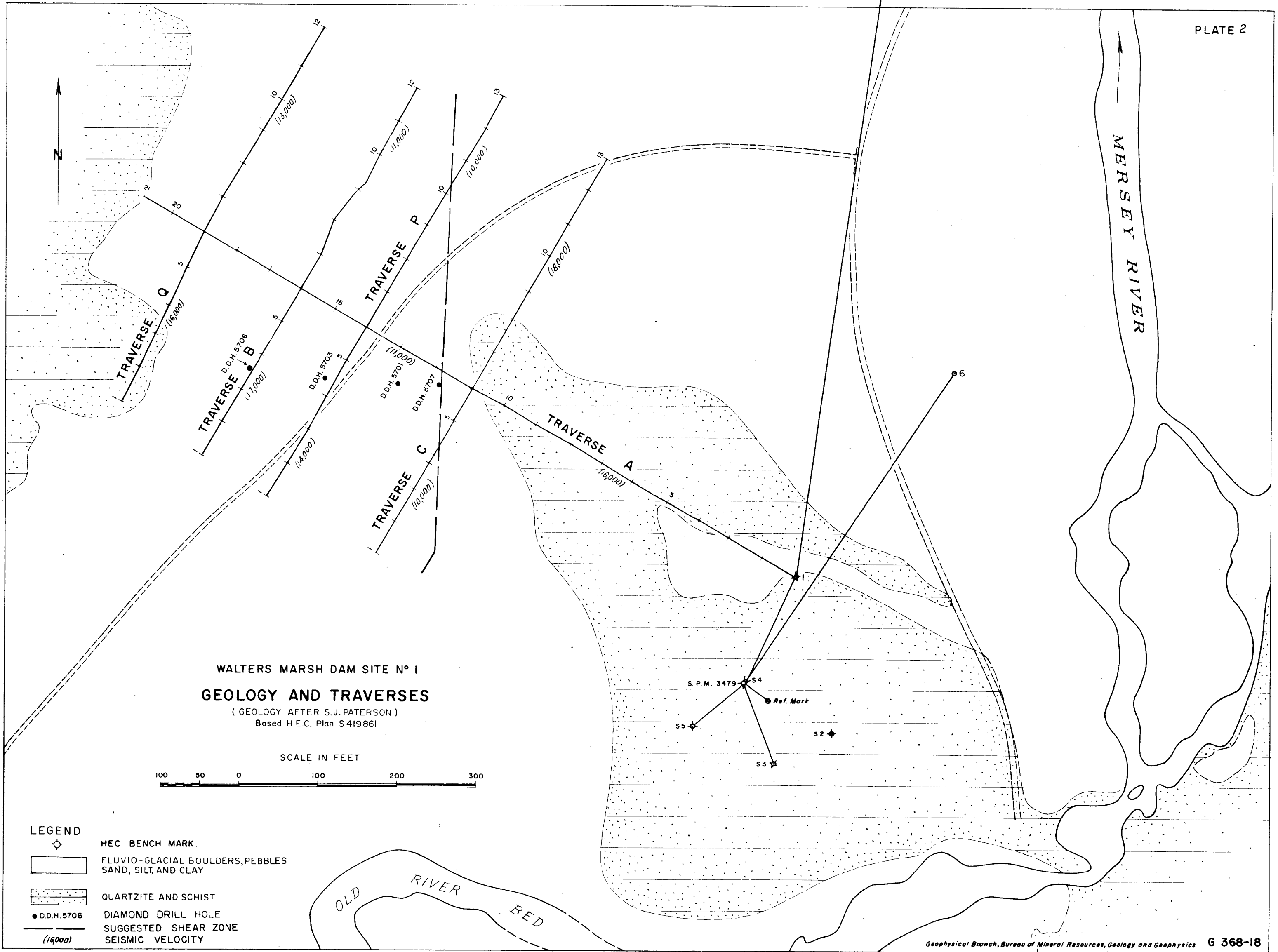
GEOLOGICAL MAPPING BY A.M. SPRY (1956).



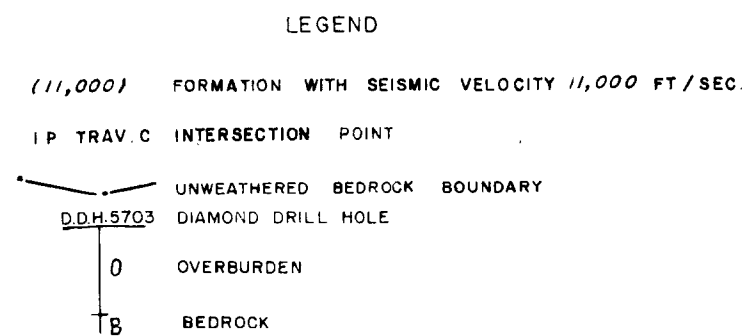
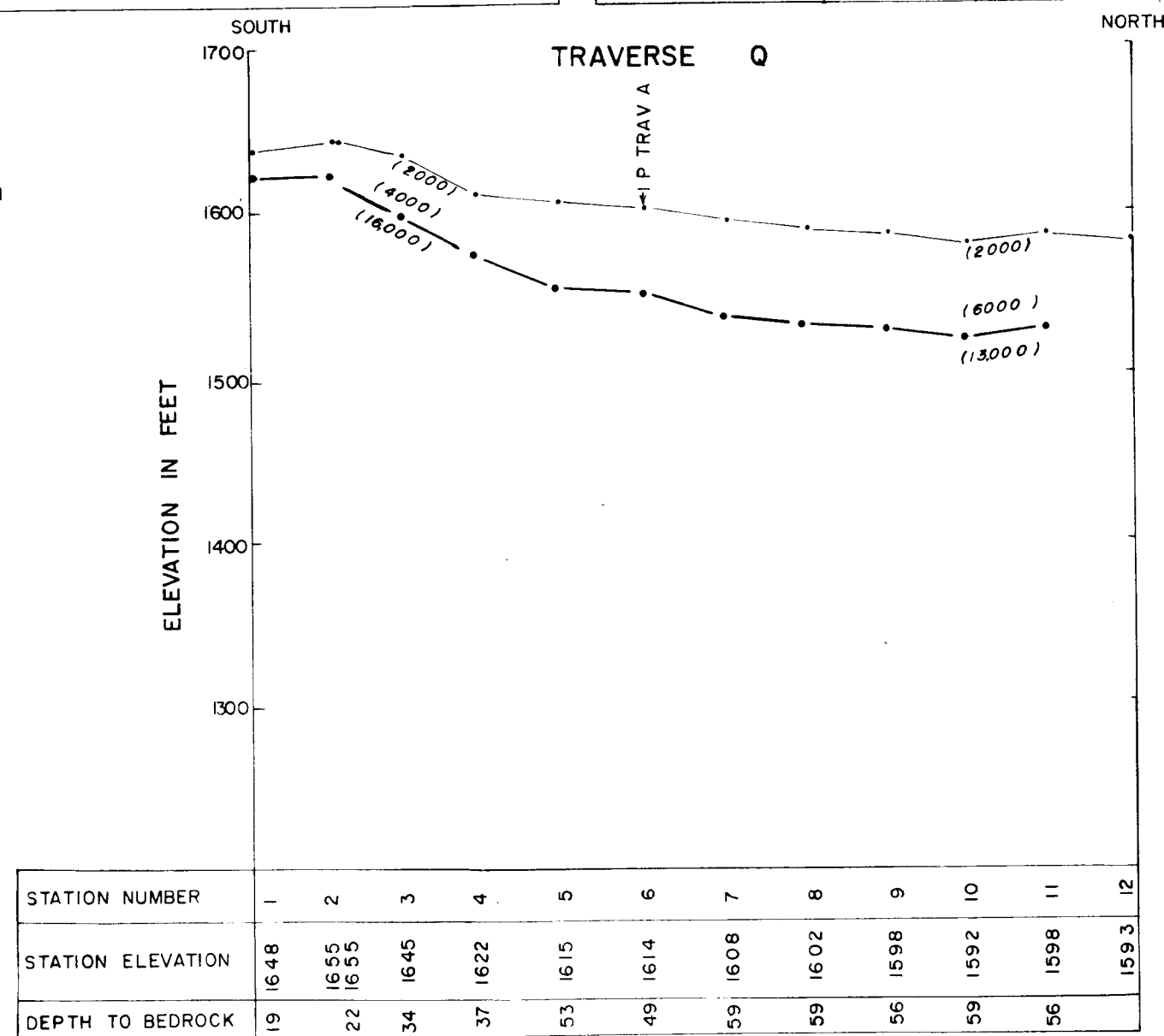
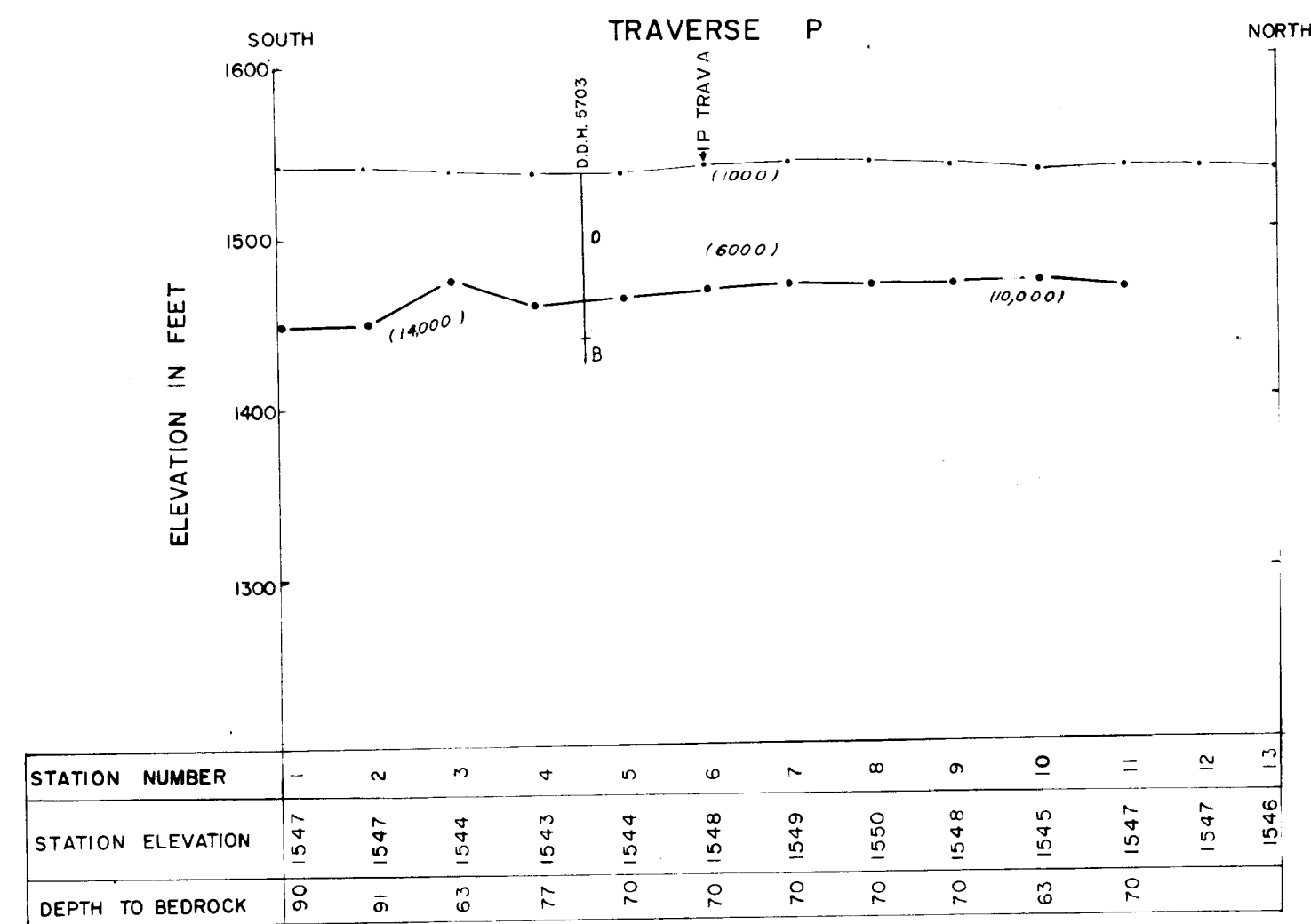
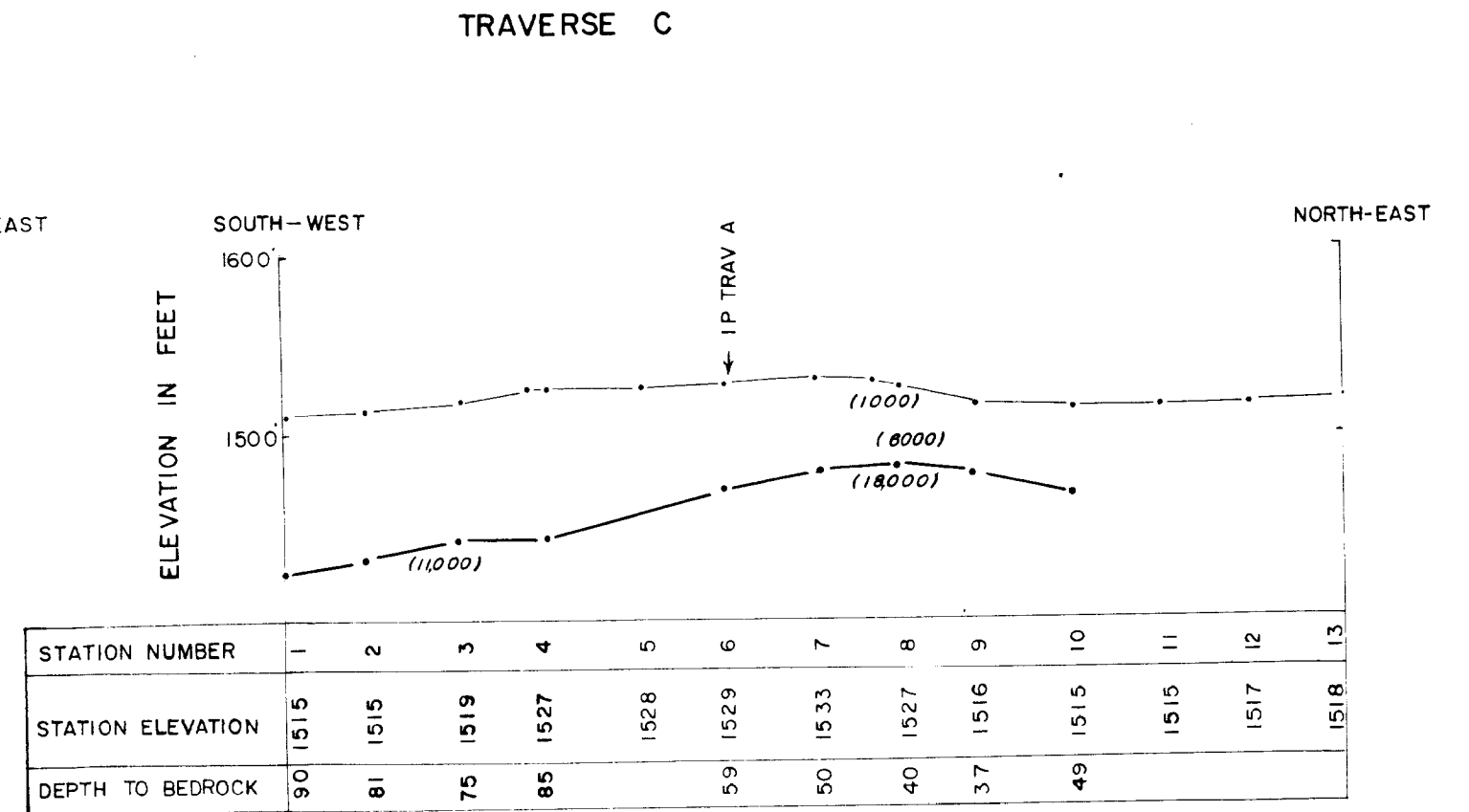
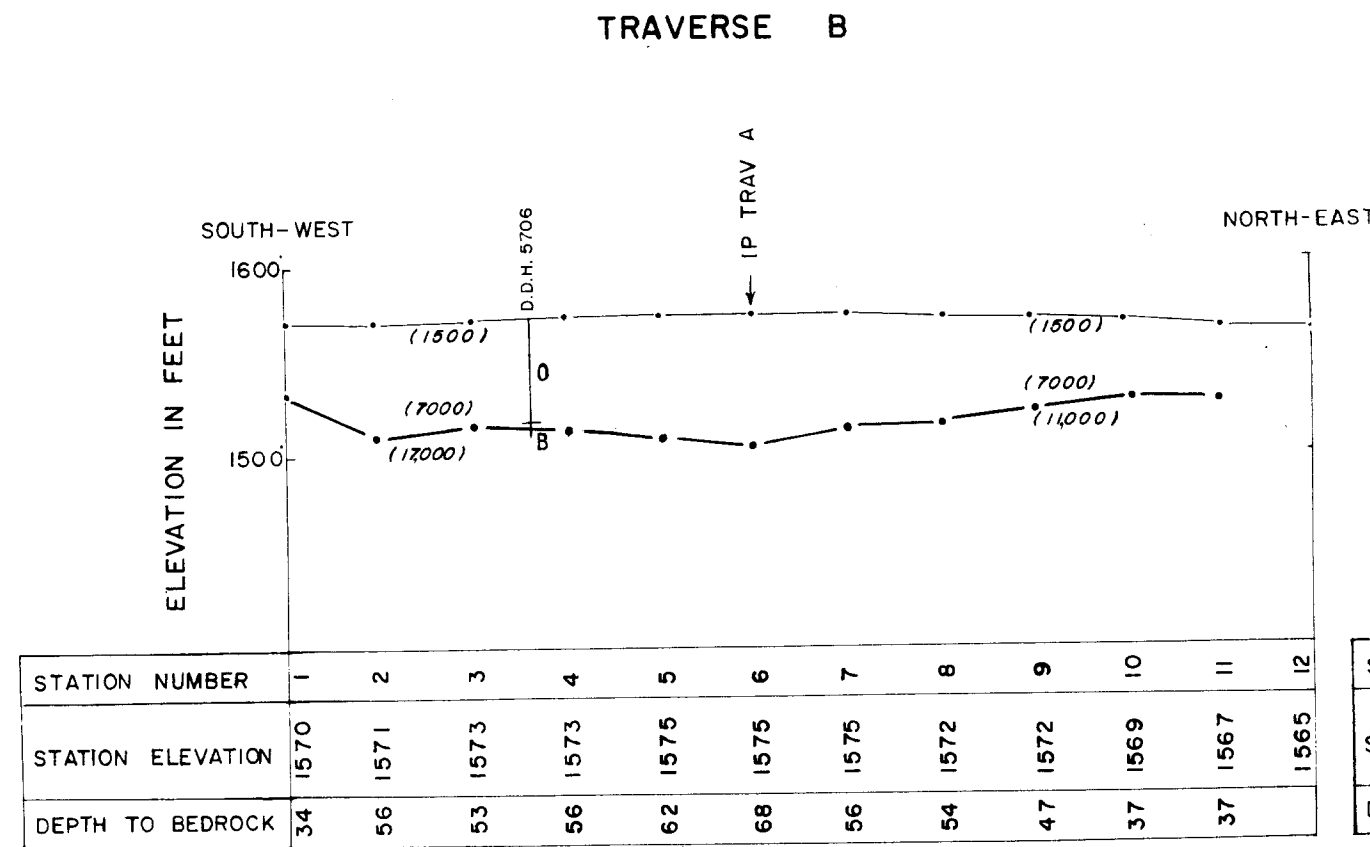
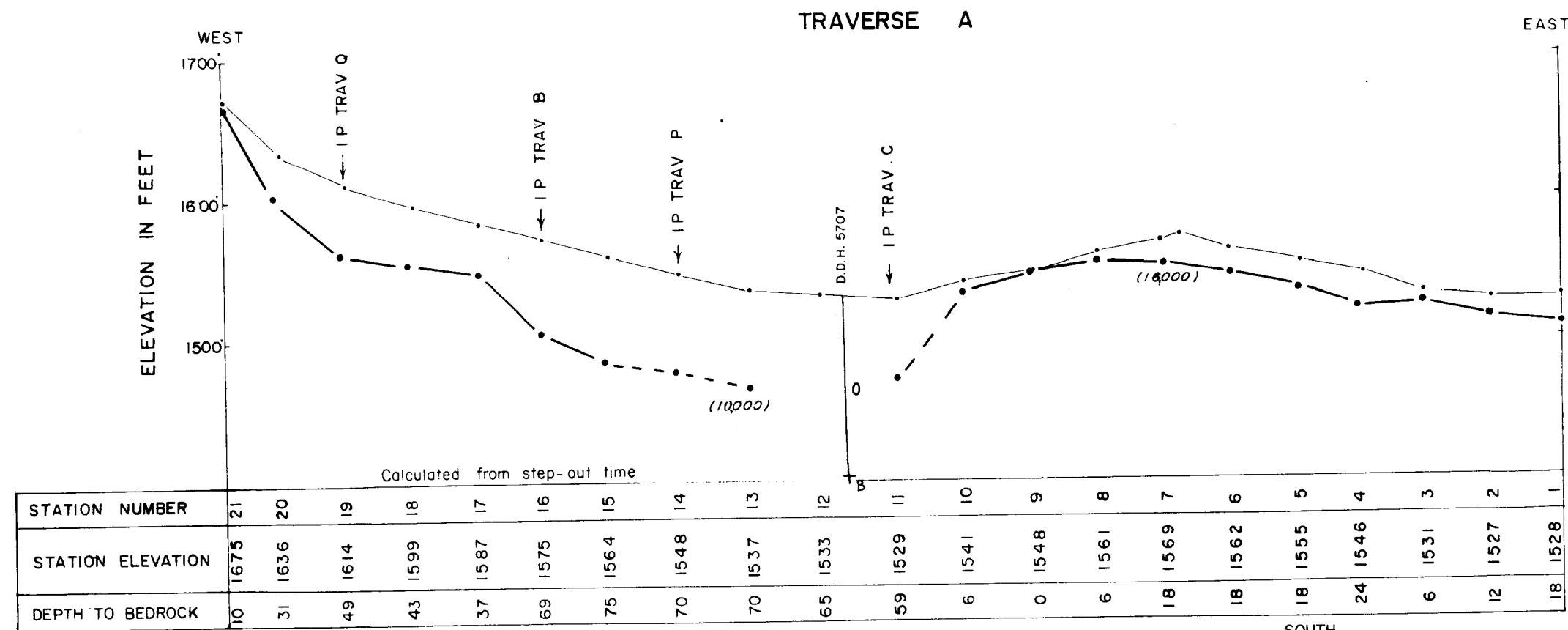
WALTERS MARSH DAM SITES
MERSEY RIVER, TASMANIA 1961
GEOPHYSICAL SURVEY



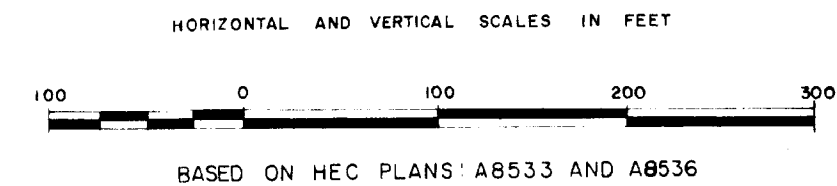
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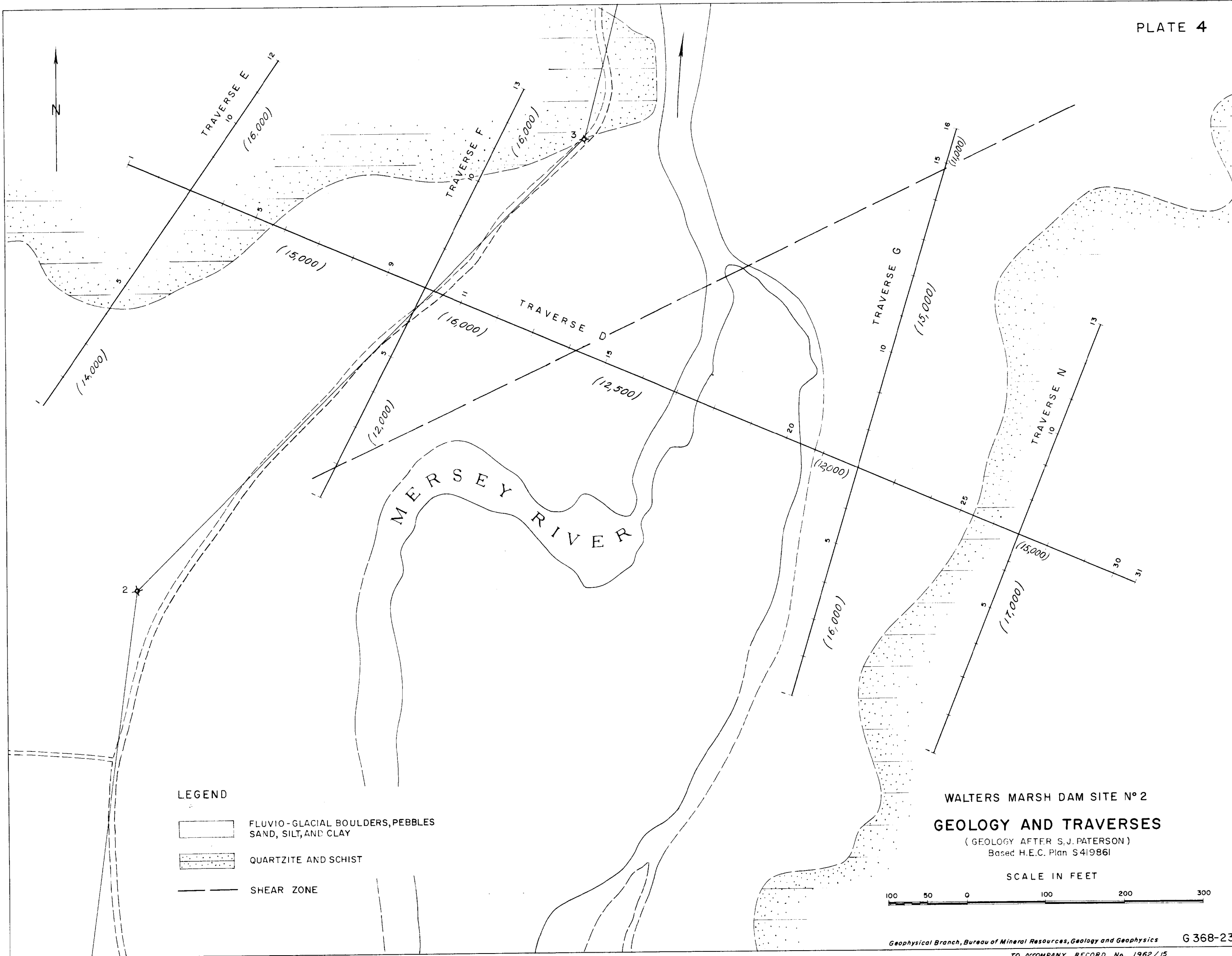


WALTERS MARSH DAM SITE



DAM SITE No. 1
TRAVERSES A, B, C, P, AND Q
CROSS-SECTIONS



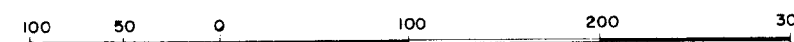


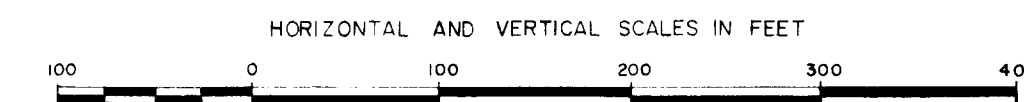
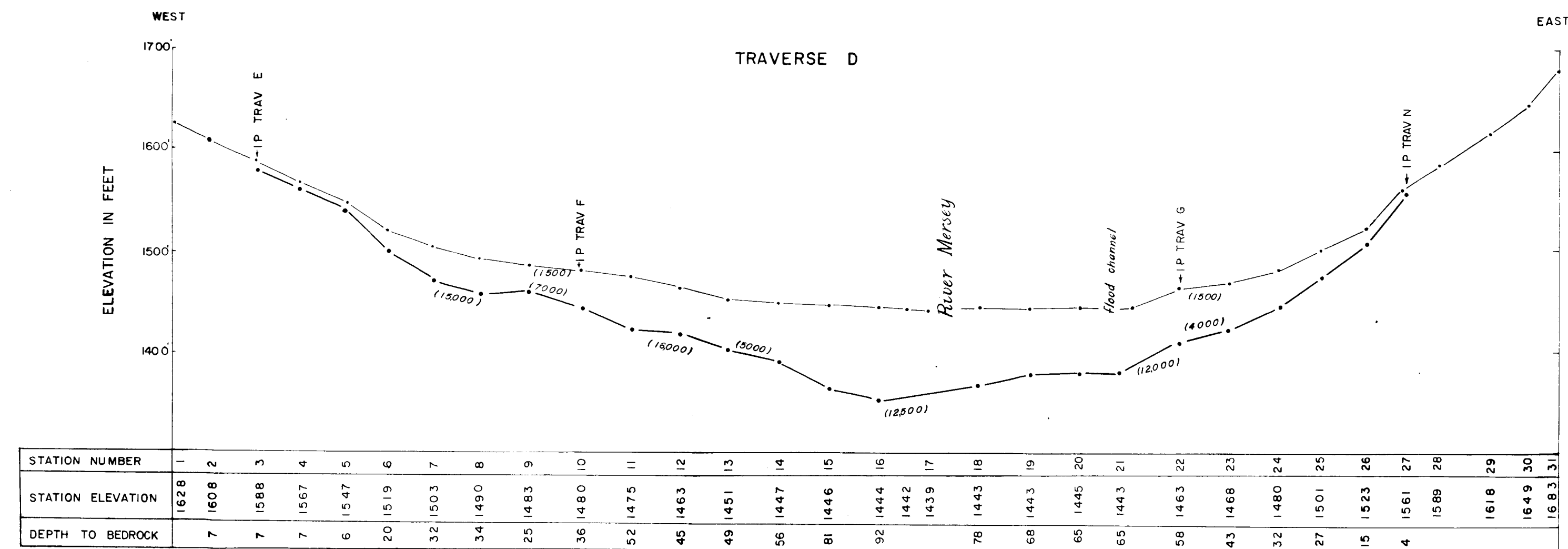
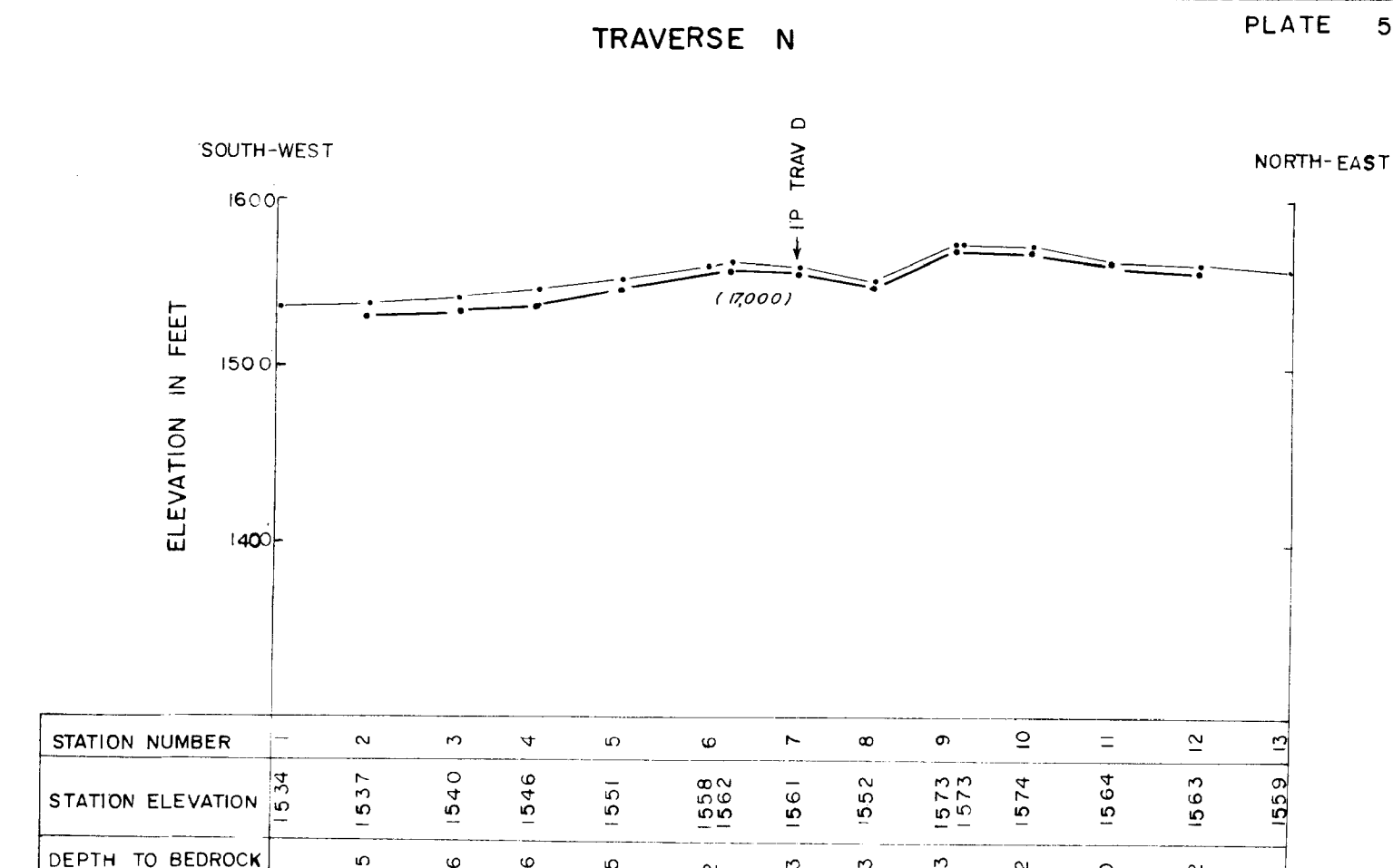
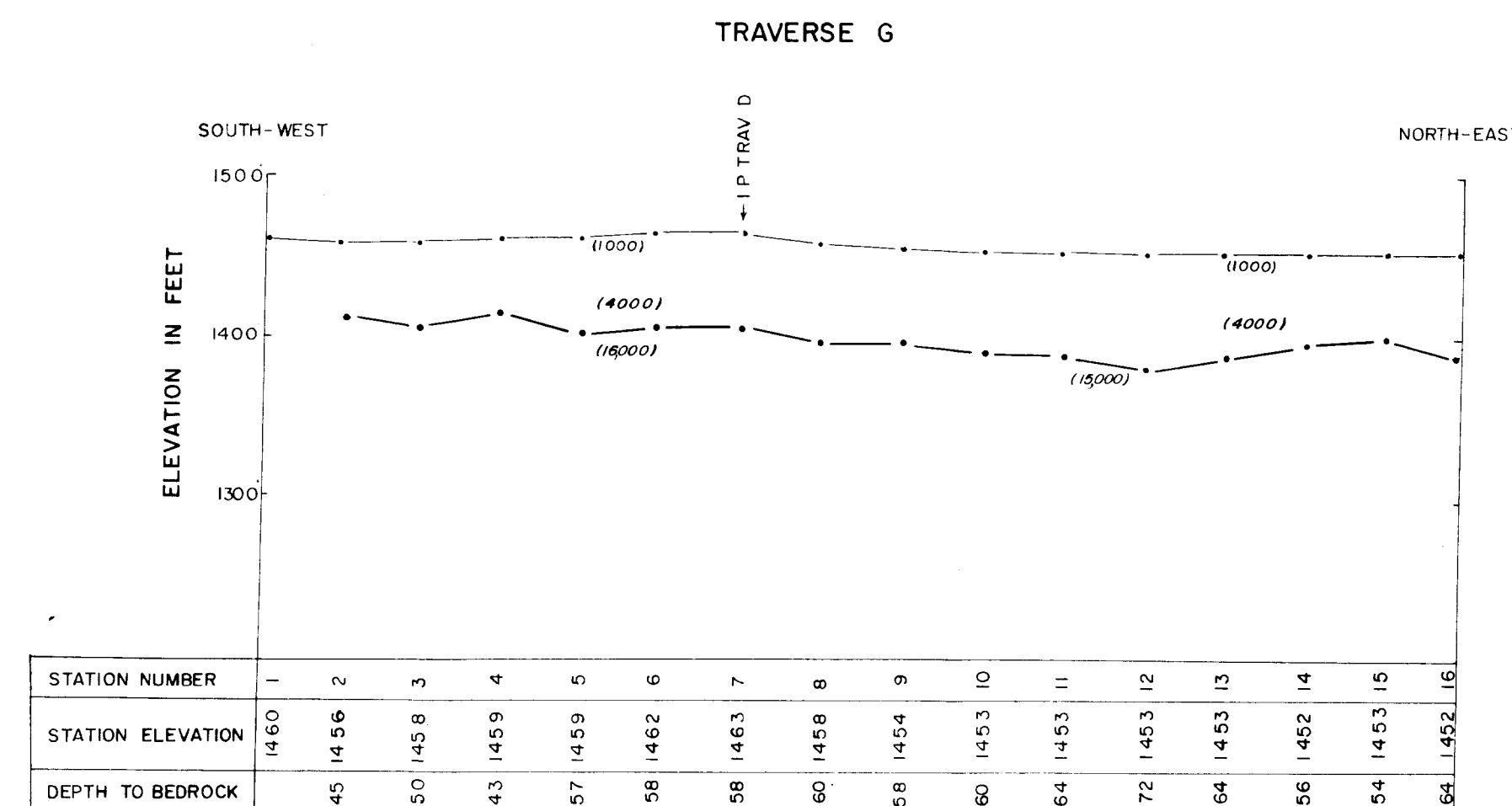
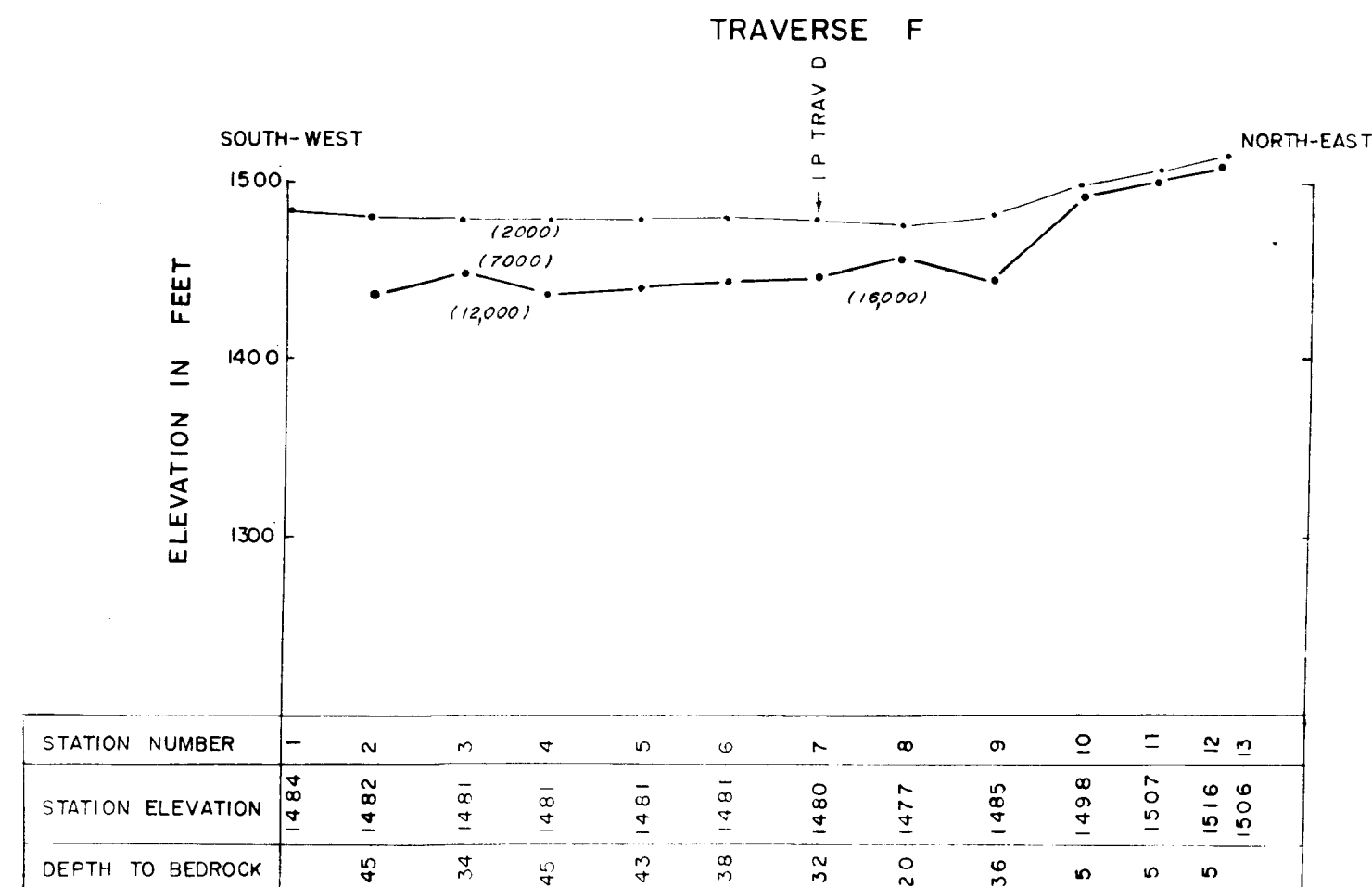
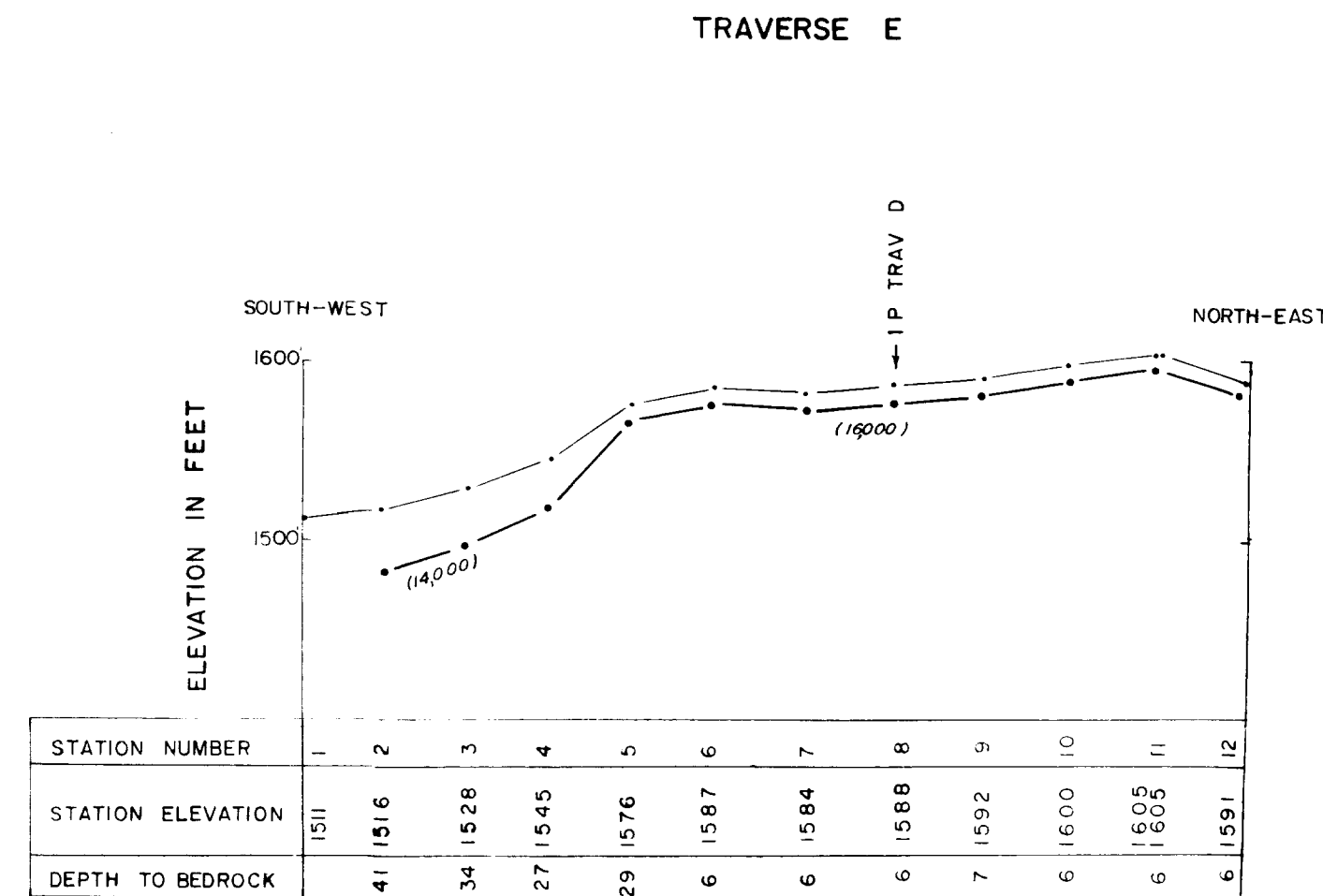
LEGEND

- FLUVIO-GLACIAL BOULDERS, PEBBLES
SAND, SILT, AND CLAY
- QUARTZITE AND SCHIST
- SHEAR ZONE

WALTERS MARSH DAM SITE N° 2
GEOLOGY AND TRAVERSES
 (GEOLOGY AFTER S. J. PATERSON)
 Based H.E.C. Plan S419861

SCALE IN FEET





BASED ON HYDRO-ELECTRIC COMMISSION TAS. PLANS A8533 & A8534

DAM SITE No. 2
TRAVERSES E, F, G, N, AND D
CROSS-SECTIONS

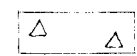
LEGEND

(10000) FORMATION WITH SEISMIC VELOCITY 10,000 FT/SEC

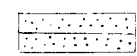
I.P TRAV. D INTERSECTION POINT

 UNWEATHERED BEDROCK BOUNDARY

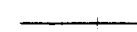
LEGEND



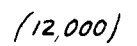
SCREE MATERIAL



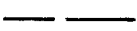
QUARTZITE and SCHIST



GEOPHYSICAL TRAVERSE



SEISMIC VELOCITY IN ft/sec.



POSSIBLE SHEAR ZONE



TRAVERSE H

TRAVERSE J

TRAVERSE K

TRAVERSE L

TRAVERSE M

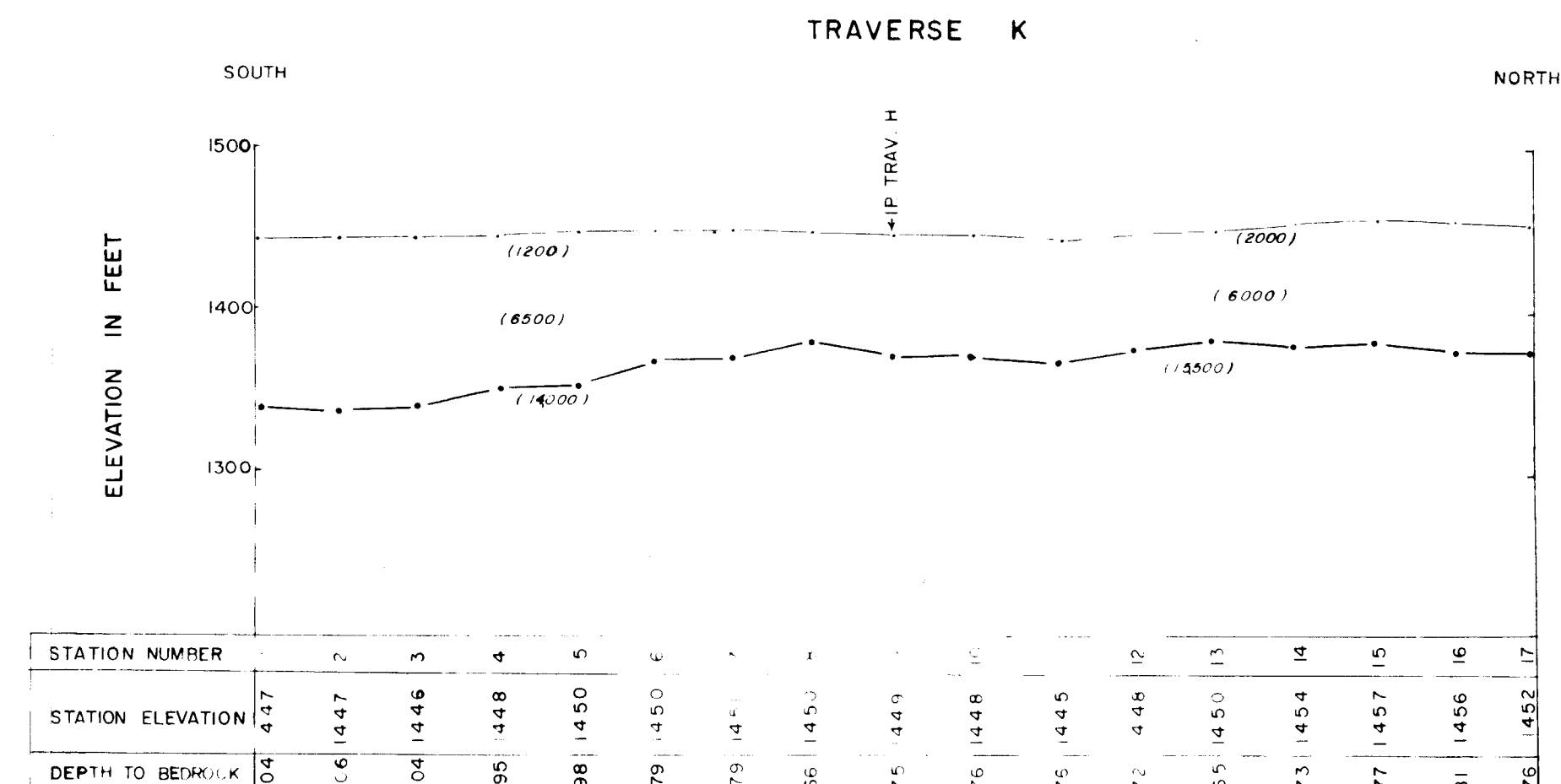
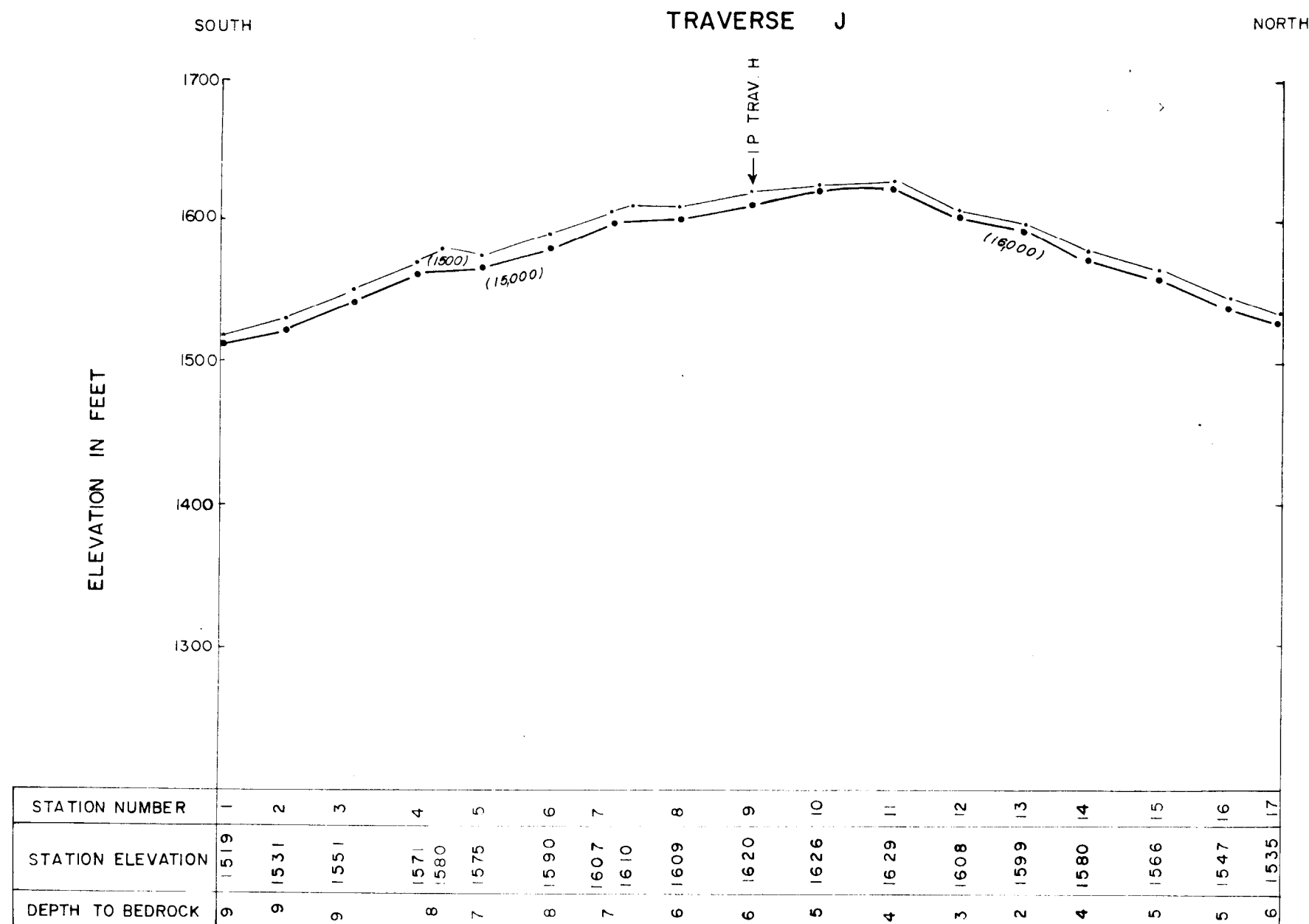
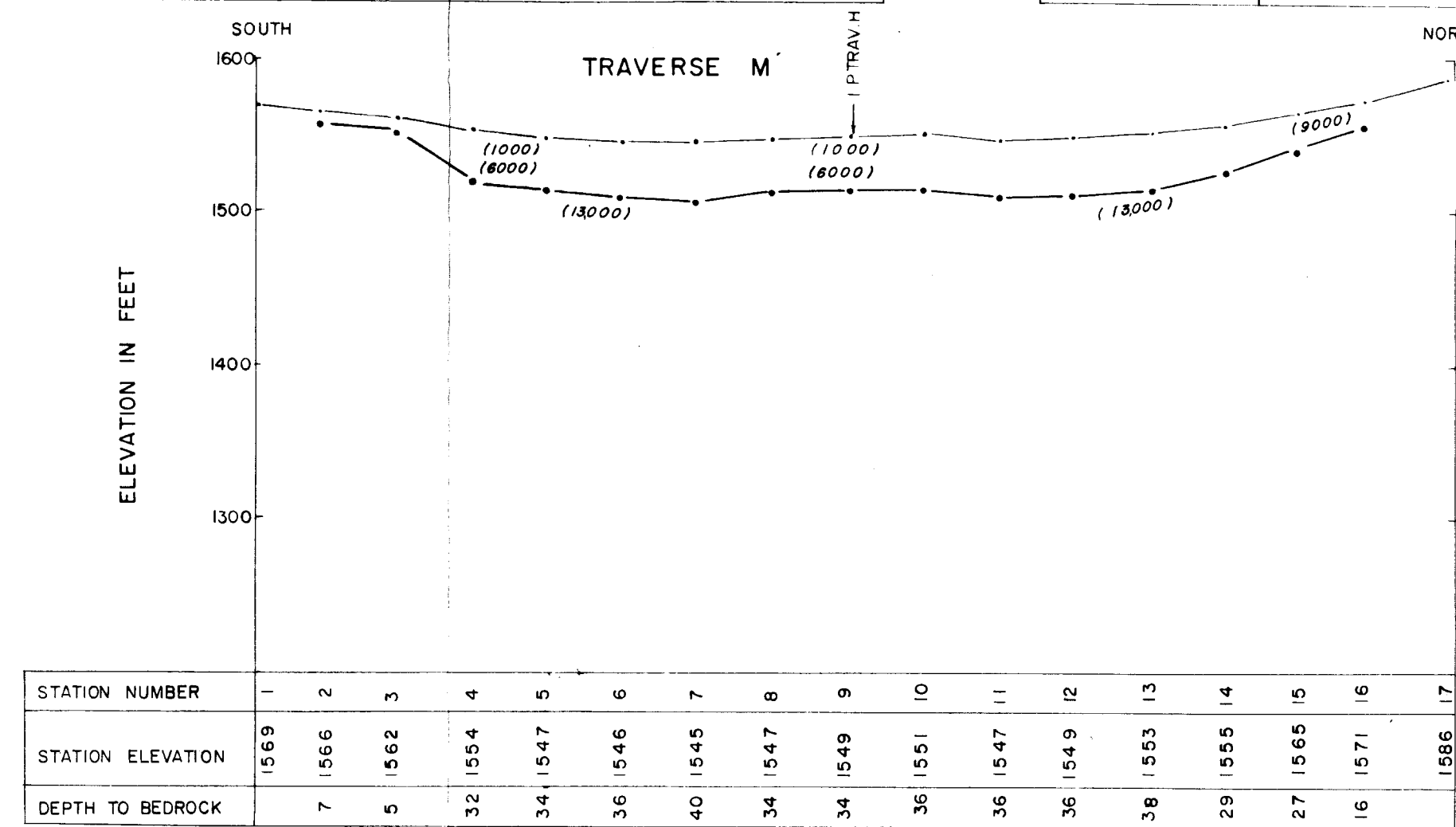
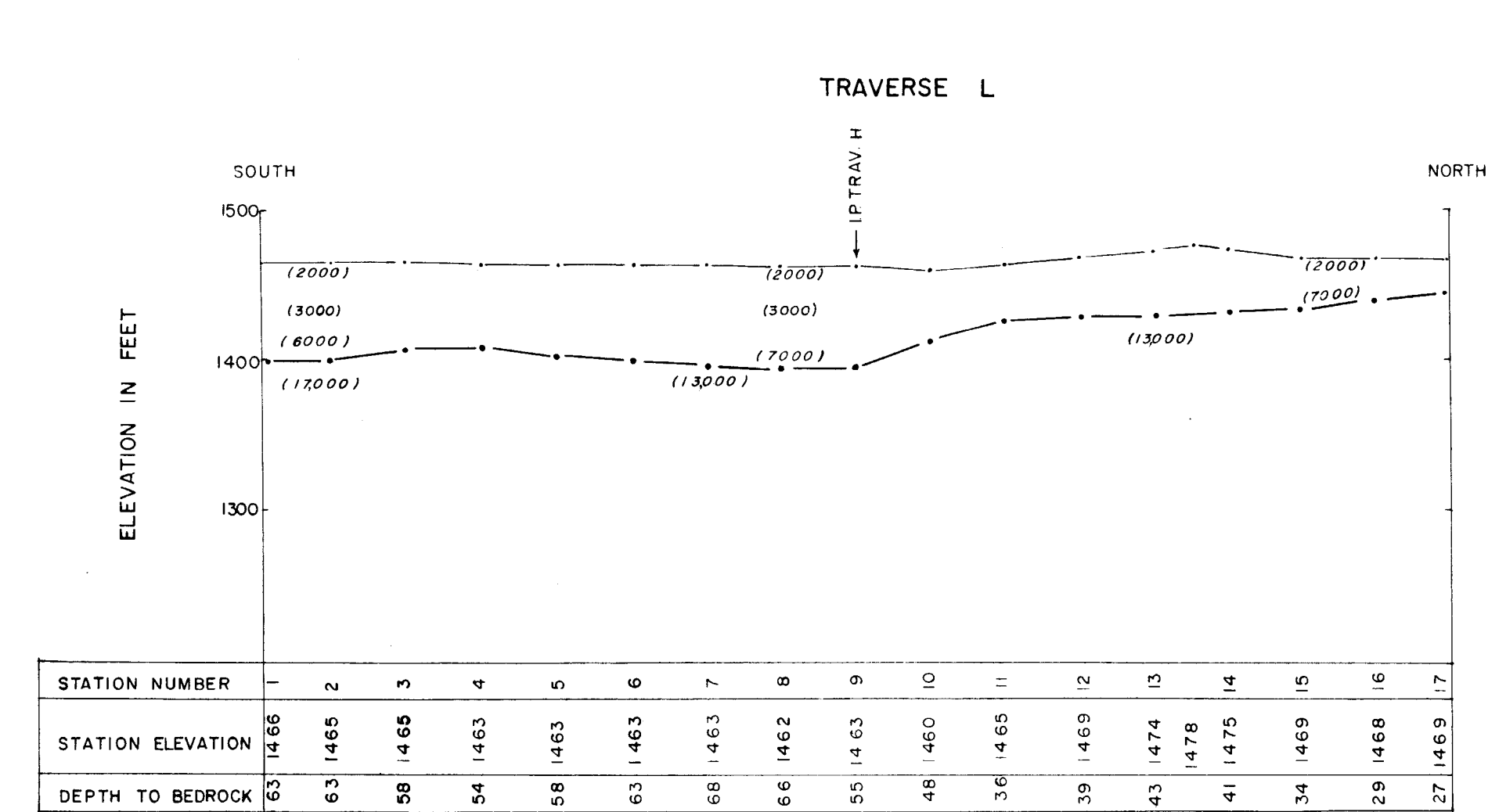
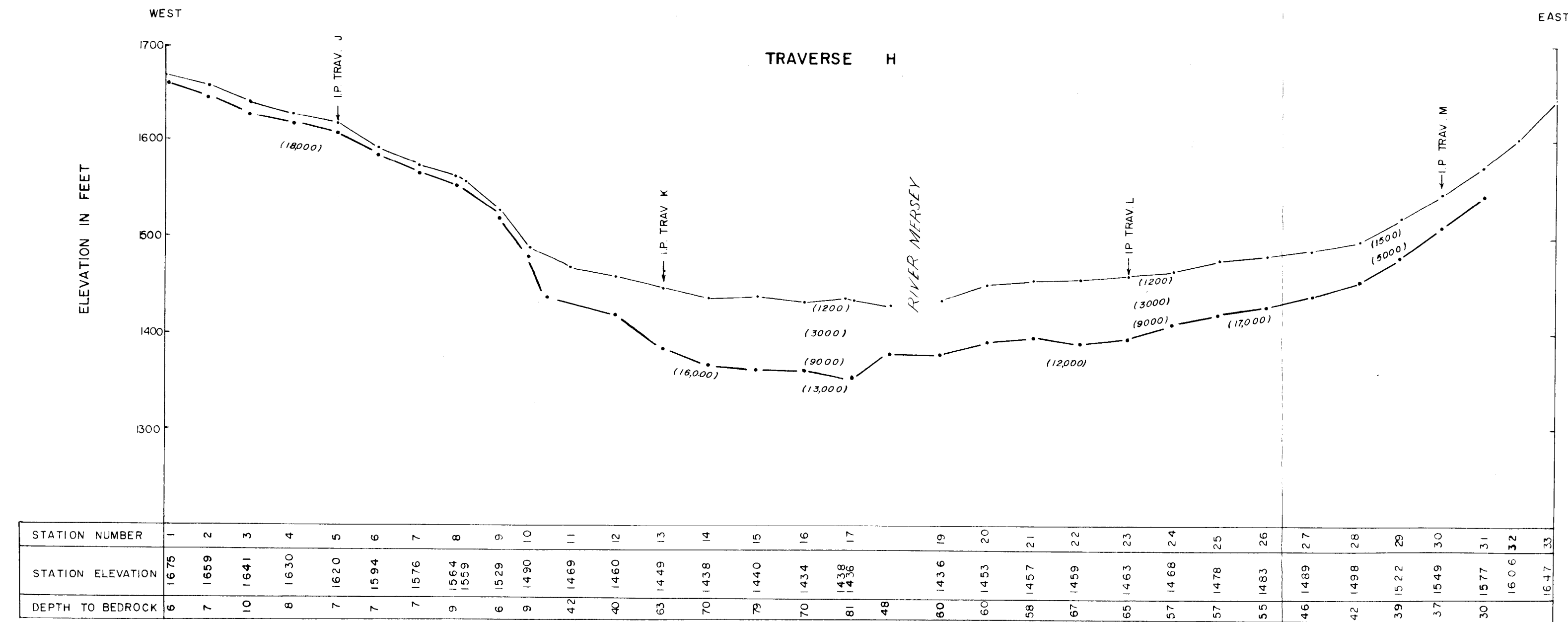
MERSEY RIVER

WALTERS MARSH DAM SITE N° 3

GEOLOGY, TRAVERSES, AND BEDROCK VELOCITIES

(Geology after S.J. Paterson)

BASED H.E.C. PLAN S4191/3 862



LEGEND

(15000) FORMATION WITH SEISMIC VELOCITY (15000) FT / SEC.

IP TRAV M INTERSECTION POINT

UNWEATHERED BEDROCK BOUNDARY

DAM SITE No.3

TRAVERSES H,J,K,L,M

CROSS-SECTIONS

HORIZONTAL AND VERTICAL SCALES IN FEET

100 0 100 200 300

BASED ON H.E.C. OF TASMANIA PLANS No A8535 AND 8536