## DEPARTMENT OF NATIONAL DEVELOPMENT

# BUREAU OF MINERAL RESOURCES, GEOLOGY AND GEOPHYSICS

RECORD No.1962/69

WILD RIVER/MILLSTREAM DAM SITE, GEOPHYSICAL SURVEY,

QUEENSLAND 1960

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#### SUMMARY

This Record describes a geophysical survey at a dam site which was undertaken in response to an application from the Irrigation and Water Supply Commission of Queensland. The dam will be used for water conservation.

The seismic method was used to measure the seismic velocities in different rock types and to determine the depth to rock discontinuities; some magnetic and resistivity work was done to determine the boundaries of a basalt flow, and to disclose the presence of shear zones.

Diamond drilling is recommended at several locations to check the geophysical results and to investigate whether rock formations in which the seismic velocity is lower than in the deepest refractor, could be used for dam foundations.

#### 1. INTRODUCTION

The Irrigation and Water Supply Commission of Queensland is investigating a site for a dam close to the junction of the Wild River and The Millstream near Mount Garnet, Queensland. The dam is to be used for water conservation and flood mitigation. In response to the Commission's request, the Bureau of Mineral Resources, Geology and Geophysics carried out a geophysical survey to determine the type of bedrock and depth to bedrock at a site which covered both streams. The site is referred to as the 174.0M/2.5M Wild River/Millstream dam site; its approximate co-ordinates are 318781 on the Atherton sheet of the Australia four-mile series.

A preliminary geological investigation of the dam site has been carried out by the Commission.

The Bureau's geophysical party consisted of P.E. Mann (party leader), D.J. Harwood (geophysicist), and J.P. Pigott (geophysical assistant). Four field assistants were supplied by the Commission. Field work at the site lasted from the 3rd August to the 25th August 1960.

The Commission carried out the topographical survey of the site and provided gelignite, detonators, additional transport, and some supplies for the party.

#### 2. GEOLOGY

A geological map is shown on Plate 1.

The dam site is located just above the junction of Wild River and The Millstream. The geology of the area has been described by Dunlop (1960).

The basement rock consists of a coarse-grained, pink, jointed granite and aplite (a fine-grained granitic rock in which quartz predominates). Granite forms the hill on the right bank of the Wild River; the hill on the left bank of The Millstream consists of aplite.

Tertiary basalt, overlying the granitic basement, forms the narrow plateau separating the Wild River and The Millstream. The slopes of the plateau are covered by basalt blocks and boulders, generally less than 9 in. in diameter. The basalt has a blocky fracture, is slightly vesicular, and occasionally shows columnar jointing.

High-level alluvial terraces of gravel, sand, silt, and clay are at the base of both banks of the Wild River; the left-bank terrace, which is 370 ft wide, commences abruptly with a steep face 25 ft high. Smaller, low-level terraces, 50 to 100 ft wide, border the Wild River stream beds.

At The Millstream, high-level alluvial terraces are found up to 50 ft above stream level and the left-bank terrace is over 200 ft wide.

Stream beds of both rivers consist of alluvium. Eluvial hill-wash material covers the slopes.

<sup>\*</sup> looking down-stream.

#### 3. METHODS AND EQUIPMENT

#### <u>Seismic</u>

General descriptions of the seismic refraction methods are outlined by Polak and Mann (1959), and Heiland (1946, p. 522).

The seismic equipment used on the survey was a 12-channel portable seismograph, which was designed for shallow seismic reflection and refraction work and manufactured by the Midwestern Geophysical Laboratory of Tulsa, Oklahoma, and geophones with natural frequency of about 20 c/s to record the vertical motion of the ground.

The following types of geophone spread, based on the 'method of differences', were applied:

- (a) Normal spreads the geophones were spaced 50 ft apart in a straight line and shots were fired 50 ft and 200 ft beyond each end of, and in line with, the spread.
- (b) Weathering spreads these spreads were used to obtain the seismic velocity and thickness of the soil and near-surface layers. The geophone interval was 10 ft and shots were fired at distances of 10, 50, and 200 ft beyond each end of, and in line with, the spread.

On several traverses a shot was placed in the centre of a normal spread to provide additional information about the seismic velocity and thickness of the near-surface layers.

Past experience has shown that depth determinations in this type of terrain are probably not better than within 20 per cent.

### Resistivity

In the resistivity method (Heiland, 1946, p. 741; Wiebenga, 1955) an electric current is applied to the ground through two current electrodes and the potential difference is measured between two additional electrodes which are located between the two current electrodes. The ratio of the measured potential to the applied current is a measure of the ground resistivity. In the Wenner configuration the electrodes are equally spaced. The depth of current penetration is of the same order as the electrode spacing. If the electrode arrangement as a whole is moved along a traverse, lateral resistivity variations are recorded.

Resistivity variations from place to place may be caused by variations in rock porosityand in salinity of pore solutions. Hence resistivity variations may be correlated with variations in rock type; e.g. geological formation boundaries and shear zones may be indicated. On this survey, resistivity traversing with electrode spacing of 50 and 100 ft was used to indicate and confirm the presence of shear zones and geological formation boundaries.

#### Magnetic

The magnetic method (Heiland, 1946, p. 334) depends on the susceptibility contrast between geological formations, and should therefore provide a means of disclosing discontinuities between formations carrying different amounts of magnetic minerals. Magnetic methods may be used to estimate the depth to basement rock and to indicate the presence of dykes, orebodies, shear zones, basalt flows, etc.

Usually the vertical magnetic intensity is measured. The instrument used on this survey was a Watts variometer. The magnetic method, when used alone, is rarely adequate, but in combination with other geophysical methods, it will often render valuable information.

On this survey, magnetic work was done along Traverse XX to indicate the boundaries of the basalt flow.

#### 4. RESULTS

Plate 1 shows the locality plan and the geology of the dam site, Plate 2 the traverse plan with topographical contours and bedrock velocities, and Plates 3, 4, and 5 the seismic profiles. Vertical magnetic intensity and resistivity profiles for Traverse XX are also shown on Plate 3.

#### Seismic Velocities

Table 1 was prepared from available geological information, recorded velocities, and general experience. Together with the available geological information, Table 1 may be used to interpret the seismic velocities in geological terms. Sometimes this interpretation may be ambiguous. For instance, a layer with a seismic velocity of 13,000 ft/sec could be either vesicular basalt or weathered granite. A layer with a velocity of 5000 ft/sec could be either very weathered granite or riverterrace material.

#### TABLE 1

Seismic velocity (ft/sec)	Rock type
1000 to 2000	Soil
2000 to <b>30</b> 00	Eluvium, hill-wash material; unconsolidated, near-surface sediments, un-saturated; clays.
3000 to 4000	Very weathered or decomposed igneous rocks; water-saturated alluvial sediments, predominantly clay.
4000 to 6000	Very weathered igneous rocks; water-saturate alluvial sediments or river-terrace materia
6000 to 10,000 .	Weathered, jointed, fractured, or sheared igneous rocks.
10,000 to 13,000	Vesicular basalt; jointed, fractured, or sheared, slightly weathered igneous rocks.
13,000 to 16,000	Jointed or fractured igneous rocks.
16,000 to 18,000	Unweathered igneous rocks.

### Traverses XX, F, G, H, and J (Plates 3 and 5)

The following discussions of the geophysical results are given to assist interpretation of the cross-sections and profiles. Only points which may lead to ambiguous interpretation will be mentioned.

Between XX2 and XX27 the 10,400-ft/sec layer probably is slightly weathered granite; the 5000-ft/sec layer probably is water-saturated river-terrace material.

Between XX27 and XX34, the 12,000-ft/sec bedrock probably is fractured or sheared granite and not a remnant of basalt. If it were basalt, the magnetic profile would not be so smooth. The 5000-ft/sec layer probably is river-terrace material, but it could be very weathered igneous rock.

Between XX43 and XX68 the 11,000 to 13,000-ft/sec layer is interpreted as vesicular basalt as suggested by the geological report. This interpretation is confirmed by many irregular changes on the magnetic profile. The width of the basalt as suggested by the magnetic profile (approximately XX35 to XX75) is larger than that suggested by the seismic work. This can be explained by assuming that erosion caused basalt boulders to cover the sides of the basalt ridge. The 5000-ft/sec layer probably is very weathered basalt. The seismic results indicate that contacts between basalt and granite are near XX43 and XX68. Between XX68 and XX103 (near The Millstream) the bedrock probably consists of unweathered granitic rocks with a shear or fault zone between XX85 and XX87. The 8000 to 10,500-ft/sec layer most probably is a weathered or slightly weathered granitic rock, and not basalt, because the magnetic profile is relatively smooth. The 4000 to 6000-ft/sec layers may be either river-terrace material, alluvium, or very weathered material.

South of The Millstream, between XX108 and XX159, the 5000-ft/sec layer may be either river-terrace material or very weathered igneous rock; further south and on higher ground the 4000 to 7000-ft/sec material probably is weathered igneous rock. Bedrock profiles or bedrock velocities suggest the presence of shear zones between XX115 and XX118, between XX135 and XX137, between XX149 and XX152, and between XX171 and XX177.

According to the geological plan, the southern end of Traverse XX is on aplite. However, because aplitic and granitic rocks are very similar and belong to the same family, seismic velocities cannot be used to determine the boundary between these two rock types. Between XX187 and XX202 the seismic data suggest the presence of a high-velocity layer (18,000 ft/sec) at about 240 ft depth.

#### Traverses R and K (Plates 4 and 5)

According to the geological plan, Traverses R and K are located on basalt except for the part between R122 and R160.

Between R7 and R46, most of the 11,000 to 13,500-ft/sec velocity layer probably consists of basalt, the 9000-ft/sec layer of jointed basalt, and the 5000-ft/sec layer of weathered basalt.

Farther north-east, e.g. between R70 and R94, the intermediate layer with 7000 to 10,000-ft/sec velocity may be jointed basalt, overlying higher-velocity granitic rocks. The profile along Traverse K may be explained in the same way.

Between R94 and R101 the seismic velocity in bedrock (11,500-ft/sec) is considerably lower than in the neighbouring area. This suggests the presence of a shear zone in the elongation of a valley, marked by the 2150-ft contour line (see Plates 1 and 2).

Other shear zones are suggested between R112 and R122, between R127 and R132, and between R143 and R145.

### 5. CONCLUSIONS AND RECOMMENDATIONS

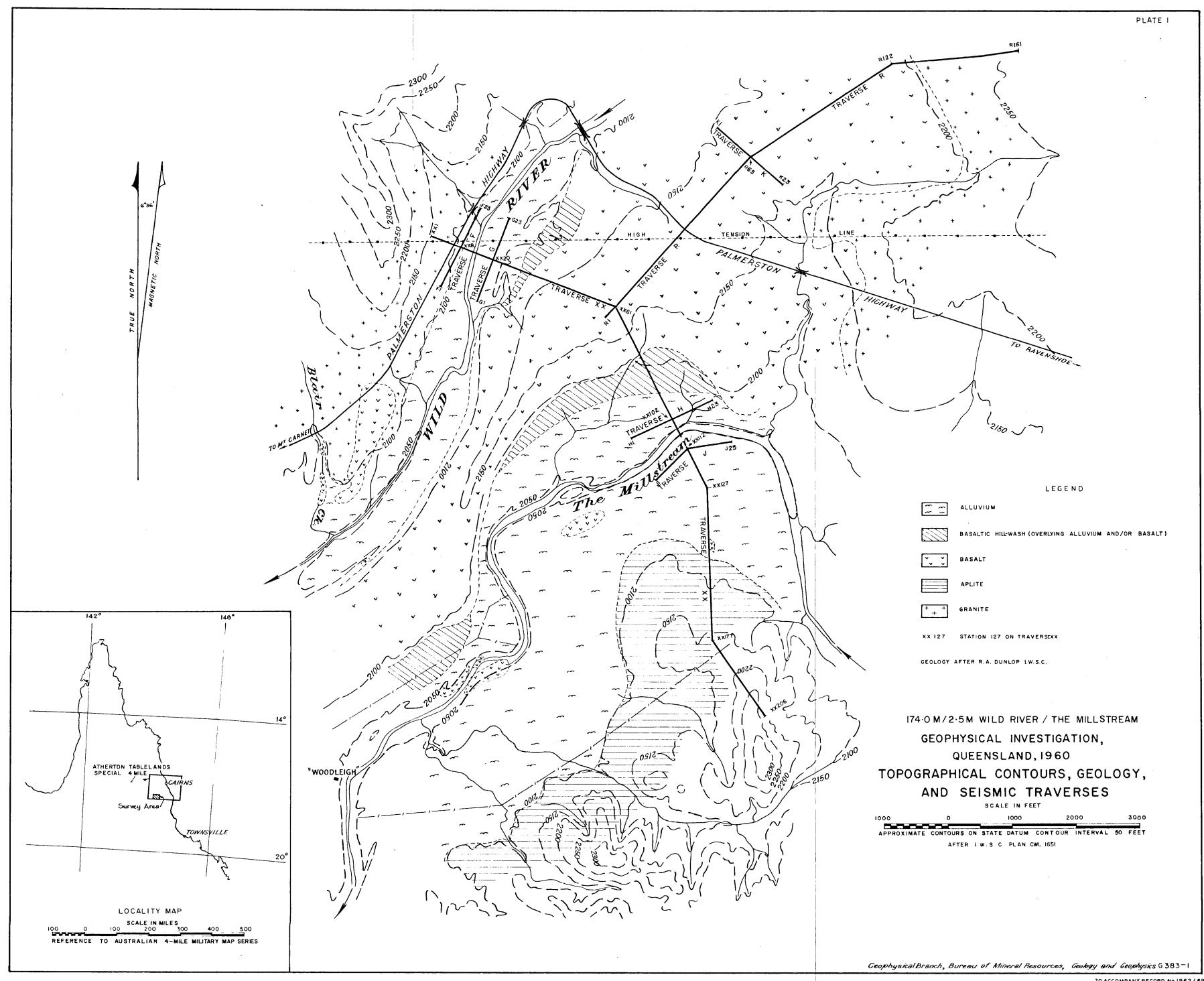
With the exception of the area between XX186 and XX202, the deepest refractor detected has been called bedrock. However, it may be possible to use a shallower refractor as a foundation rock. Seismic velocity can be used as an indicator of rock strength: in general, the higher the velocity the stronger the rock. In many places weaker rocks in which the seismic velocity will be lower than in unweathered rock, may still serve as good foundation rocks, depending on the design requirements and engineering techniques used. Therefore, the velocities of the rock layers above the deepest refractor (bedrock) are shown on the cross-sections wherever possible. For instance, a rock of 10,000-ft/sec velocity may be suitable for foundations, but this should be checked by drilling and laboratory tests. Several diamond-drill holes are recommended to check the results of the geophysical survey and to provide coresfor the laboratory investigation of some relevant rock properties. Table 2 shows the location of these drill holes and the information it is expected the holes would provide.

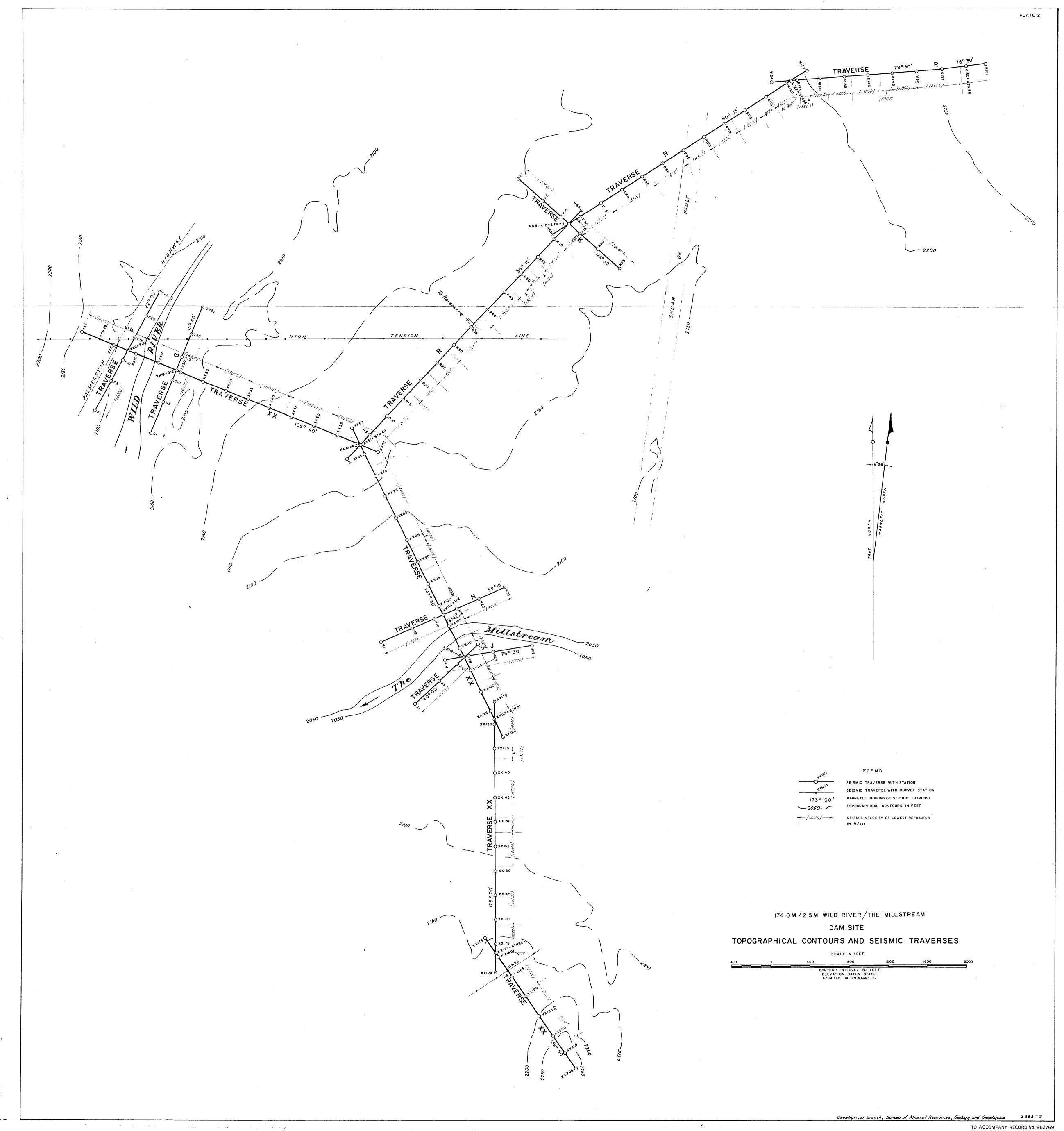
#### TABLE 2

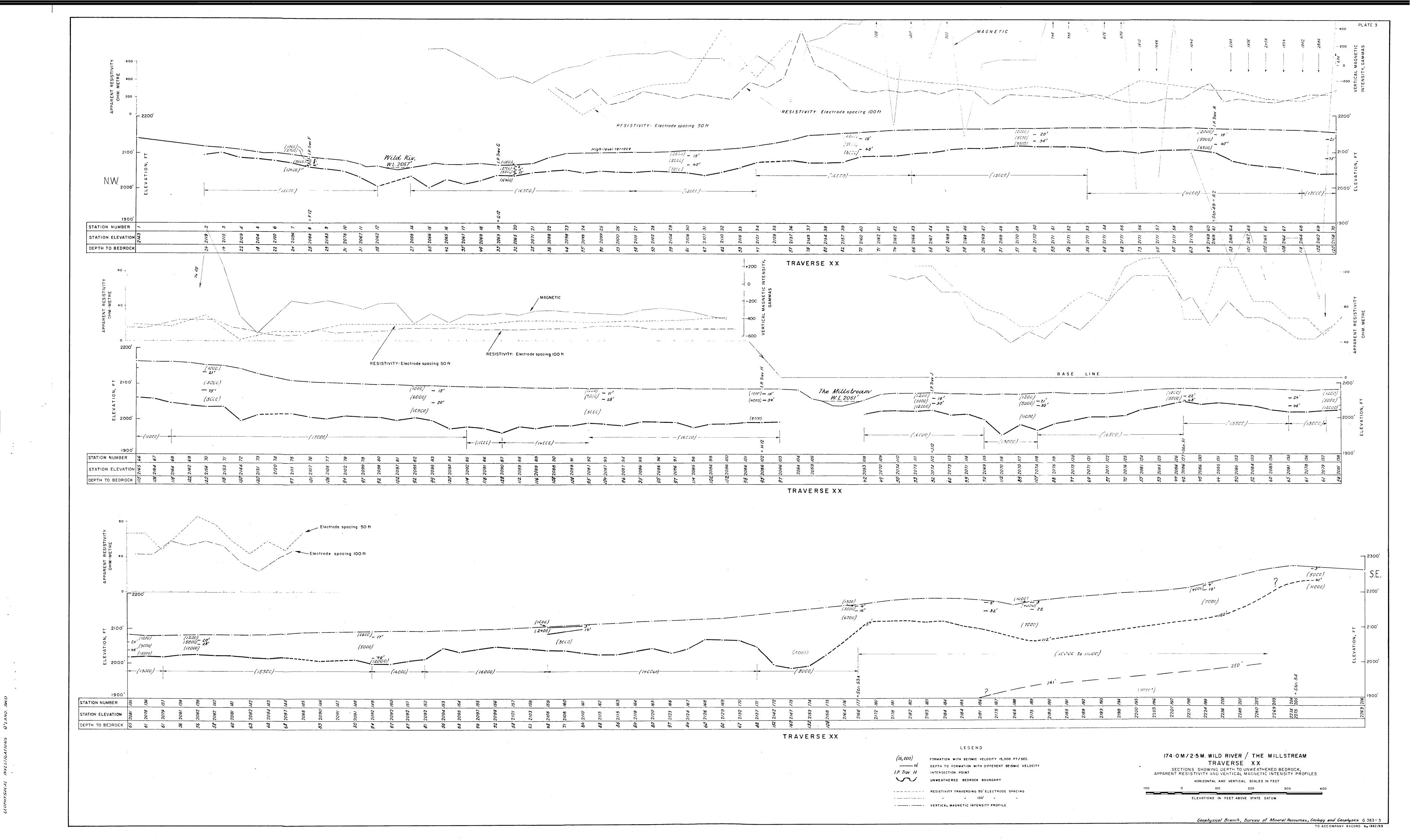
Location Dri	<u>ll Hole</u>	Information expected from	om drill hole
XX 173	I	epth and quality of fau bedrock.	lted or sheared
XX <b>1</b> 50	I	epth and quality of the	12,000-ft/sec layer.
XX 116	C	Quality of the 12,000-ft, presence of fault or sh	
XX 87	I	Depth and properties of 10,500-ft/sec layers, a shear or fault zone.	
XX <b>31</b>		Quality of the 7200-ft/s	
R 113	I	Depth and quality of the	9000-ft/sec bedrock.
R 129	I	Depth and quality of the	7000-ft/sec bedrock.
R 53	I	Depth and quality of the ft/sec layers.	8000 and 11,000-
R 97	C	Quality of the 11,500-ft the presence of shear o	

# 6. REFERENCES

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HEILAND, C.A.	1946	GEOPHYSICAL EXPLORATION. Prentice Hall, New York.
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WIEBENGA, W.A.	1955	Geophysical investigations of water deposits, Western Australia. <u>Bur. Min. Resour. Aust. Bull.</u> 30, pp. 4 - 7.







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