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#### COMMONWEALTH OF AUSTRALIA

# DEPARTMENT OF NATIONAL DEVELOPMENT

# BUREAU OF MINERAL RESOURCES, GEOLOGY AND GEOPHYSICS

RECORD No. 1965/203



NORSEMAN DETAILED
AEROMAGNETIC SURVEY,
WA 1964, WITH AN
INTERPRETATION OF DATA FROM A
REGIONAL AEROMAGNETIC SURVEY,
1959

010688

by

B.A. DOCKERY and W.A. FINNEY

The information contained in this report has been obtained by the Department of National Development as part of the policy of the Commonwealth Government to assist in the exploration and development of mineral resources. It may not be published in any form or used in a company prospectus or statement without the permission in writing of the Director, Bureau of Mineral Resources, Geology and Geophysics.

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Plate 2.	Total magnetic intensity contours, detailed survey (I51/B1-20)
Plate 3.	Total magnetic intensity contours, regional survey (I51/B1-21)

1965/203

#### SUMMARY

A detailed aeromagnetic survey was made over an area of about fourteen square miles to the north-east of Norseman, WA, in August 1964 at the request of New Consolidated Gold Fields (Aust) Pty Ltd. The object of the survey was to aid geological mapping in the soil-covered portion of the area, and it was expected that detailed geological information would become available from the co-operating company and would enable an adequate interpretation to be made. Owing to delays in the company's programme, no detailed geological data were produced and therefore no significant interpretation of the results of the detailed survey was possible. However, the results of a Bureau of Mineral Resources regional aeromagnetic survey of 1959 have been examined more closely for a larger area and interpreted with the aid of recent geological information. The results in this Record are mainly derived from this interpretation.

At Norseman the rocks form the western limb of a regional anticline of Precambrian age and are divided stratigraphically into the Mount Thirsty Beds, the Abbotshall Beds, the Woolyeenyer Group, the Noganyer Group, and the Penneshaw Beds.

Two large amplitude anomalies are attributed to metajaspilite or ultrabasic beds within the Mount Thirsty Beds. The Abbotshall Beds cause an anomaly of moderate amplitude. The general magnetic expression of the Woolyeenyer Group is a featureless area of low magnetic field. A semicircular structure within the Woolyeenyer Group is the source of a zone of high magnetic field and warrants further investigation. The western limb of the anticline in the Noganyer Group shows up as a very large amplitude anomaly. It is suggested that the eastern limb of the Noganyer Group hybridized by the Eastern Granite, can also be traced in the magnetic pattern. The Penneshaw Beds might form the core of the regional anticline.

#### 1. INTRODUCTION

During the first fortnight of August 1964, a detailed aeromagnetic survey was made over an area of about fourteen square miles north-east of Norseman (Plate 1). The survey was made at the request of New Consolidated Gold Fields (Aust) Pty Ltd, who held Temporary Reserves over the surveyed area, and was supported by the Geological Survey of Western Australia. The aim of the survey was to aid geological mapping in the soil-covered portion of the survey area.

Detailed geological information covering the survey area was expected to be produced by the co-operating company and to be available at the time the survey took place. However, delays in the company's programme prevented such information being obtained. No significant interpretation of the detailed aeromagnetic data is therefore possible.

Recent regional geological information for the Norseman area has become available from work done by Western Mining Corporation Ltd (Hall & Bekker, 1964), and this Record, although presenting the results of the detailed aeromagnetic survey, deals mainly with the interpretation of part of the results of a Bureau of Mineral Resources (BMR) regional aeromagnetic survey of the Norseman area in 1959 (BMR, 1965) in relation to this more recent geological information.

No other geophysical work is known to have been done in the survey area.

#### 2. GEOLOGY

Norseman lies in the Dundas Goldfield of Western Australia and is about 100 miles south of Kalgoorlie. An account of the geology of Norseman was given by Campbell-(1904), but very little has been published since this work. Two short articles were written by Ellis (1953) and O'Driscoll (1953). The most recent paper on the goldfield is by Hall and Bekker (1964), whose results are shown in Plate 1 and quoted and summarized in this section.

#### General geology

"Norseman lies near the southern end of a large mass of Precambrian greenstones and meta-sediments, which extends northward through Coolgardie and Kalgoorlie. This greenstone area is surrounded by the extensive granitic rocks of the Western Australian Precambrian Shield. The rocks at Norseman consist of sediment and basic lavas which strike NNE and dip uniformly west at about 55°, forming part of the west limb of a regional anticline which plunges flatly north. Basic and ultrabasic sills and dykes are numerous throughout the area which has also been intruded by a swarm of albite porphyry dykes. The rocks show a low to moderate degree of regional metamorphism.

The last phase of igneous activity was the intrusion of east-west trending doleritic and noritic dykes, the principal example being the Jimberlana Dyke.

Marine Tertiary and younger sediments occur on the lower-lying areas of the Precambrian surface, especially in the vicinity of the salt lakes."

Two types of gold mineralization are present in the field. Minor bedded quartz-sulphide lodes occur as replacements of folded or brecciated structures within the jaspilites (see below). The more important mineralization occurs as auriferous quartz reefs which strike approximately north and dip about 45° east. "Their main development occurs over a length of 7 miles in the stratigraphic zone which embraces the Crown-Bluebird sill sequence and the Mararoa Pillow Lava".

#### Precambrian rock succession

"The Archaean basic meta-igneous rocks and meta-sediments of the Norseman area have a fairly uniform dip to the west.

From the tope of the succession to the bottom, that is, from west to east in the field, the stratigraphic column is as follows:-

Name of Unit

Thickness (in feet)

Lithology

# Western granite contact

A. Mt. Thirsty Beds

25,000<sup>+</sup>

Slates, greywackes, quartzites, breccias, meta-jaspilites, meta-basaltic lavas and probably conformable ultrabasics.

B. Abbotshall Beds

800 to 4,500

Meta-jaspilites with interbedded basaltic rocks and clastic meta-sediments.

C. Woolyeenyer Group 28,000 to 32,000

l. Desirable Pillow Lava

20,000 +

Meta-basaltic pillow lavas with minor graphitic slate members.

2. Crown Basalt

600

Multiple sill of massive alphanitic to strongly ophitic meta-basalt, with remnant bands of slate.

3. Nulsen Slate

3 to 6

4. Royal Amphibolite 370 to 650

Graphitic slate.

Complex sill of metaquartz gabbro, metadolerite and remnant extrusive meta-basalts. Characterized by a weak schistosity and by the occurrence of small carbonate masses. 250 feet below the top of the formation occurs the Empress Slate Member (maximum thickness 10 feet).

5. Bluebird Gabbro 200 to 350

Sill of massive metaquartz gabbro characterized by a coarsely porphyritic lower layer.

	6.	Gee Cee Slate	½ to 6	Graphitic slate.
	7.	Mararoa Pillow Lava	4,200	Meta-basaltic pillow lava.
	8.	Venture Slate	3 to 10	Graphitic slate, sometimes partially replaced by quartz and dilated by intrusions of doleritic rocks.
	9.	Kingswood		
	<i>,</i>	Basalt	1,400	Meta-basaltic rocks, including pillow lavas
D.	Nogs	anyer Group	4,500 +	A sedimentary sequence with well defined beds of meta-jaspilites, meta-conglomerates, sandstone, graphitic slate, biotite-andalusite schist, garnet-chlorite schist and mica schist.  The Lady Mary Formation is a multiple sill of meta-gabbro and meta-dolerite.
E.	Peni	neshaw Beds	10,000 +	Meta-basaltic pillow lavas and other basic and acid metamorphic rocks.
		<b></b> .		

Eastern Granite Contact"

#### 3. RESULTS

The results of the survey are shown as a contour map of the total magnetic field intensity over the survey area (Plate 2). This map may be compared directly with the contour map of the total magnetic field intensity obtained from the BMR regional aeromagnetic survey of 1959 (BMR, 1965), part of which is shown in Plate 3.

From a comparison of the two contour maps it is apparent that the detailed survey has confirmed the results obtained by the regional survey and shows the various magnetic features in greater detail and greater contrast, but, because of the limited geological knowledge of the area within and adjacent to the detailed survey area, the results of the detailed survey contribute little to what was already apparent from the results of the regional survey.

A study of the regional aeromagnetic contour map shows that the main expression of the Mount Thirsty Beds is a large amplitude magnetic anomaly cutting the Jimberlana Dyke at approximately 121 38 E and striking north-north-east. A parallel, large amplitude anomaly in the Mount Thirsty Beds cuts the Dyke about 3½ miles further east. Presumably these anomalies represent the major meta-jaspilite or ultrabasic beds of the Mount Thirsty Beds. Although not apparent on the contour map, low amplitude anomalies are present on the magnetic profiles over the Mount Thirsty Beds. These might be the

magnetic expression of the ultrabasic or meta-basaltic lavas, and the intermediate areas of low magnetic field might represent the areas of slate, greywacke, quartzites, and breccias.

Mount Thirsty Beds occur on the west limb of an anticline, which plunges gently north. The formation to the east of the Mount Thirsty Beds is the Abbotshall Beds of meta-jaspilite, which show up as a moderate amplitude anomaly on the contour map. An unusual feature is that the magnetic anomaly associated with these Beds occurs only over the parts of the Beds that crop out. Anomalies appear over the two islands in Lake Cowan that have been mapped as Abbotshall Beds, but not over the intermediate parts of the Beds that presumably underlie the Recent sediment in Lake Cowan. The most likely explanation is that the magnetic anomalies over the Beds are due to an ironstone capping over the exposed part of the Beds, and not to the unweathered rock existing at depth.

The magnetic expression of the Woolyeenyer Group south of the Jimberlana Dyke is a featureless area of low magnetic field. A study of the regional survey profiles for this area did not reveal any trends, although there were isolated low amplitude anomalies throughout. From a consideration of the lithology of the WoolyeenyerGroup, some magnetic trends were expected. In particular, the Royal Amphibolite and the Bluebird Gabbro were expected to contrast magnetically with the remainder of the Group. Such a feature might, however, be detected by a detailed magnetic survey in this region or in the vicinity of the Princess Royal Mine to the north of the Jimberlana Dyke.

A study of the profiles north of the Dyke revealed one trend in the Desirable Pillow Lava extending for five miles in a north to north-north-east direction from the point marked A in Plate 3. This might represent a more basic flow amongst the meta-basaltic pillow lavas of the Desirable Pillow Lava. A feature that occurs over what is presumed to be Desirable Pillow Lava is the semicircular zone of high magnetic field to the north of the peninsular that contains the Princess Royal Mine. An expression of the south-east end of this feature appears in the north-west corner of the detailed survey area. From a consideration of the lithology and magnetic characteristics over the known outcrop of the Desirable Pillow Lava, this semicircular feature must be due to an unidentified source within this formation. It is probably stratigraphically controlled, representing the north-pitching nose of the major anticline that affects the whole area discussed in this Record.

Four separate anomalies, B, C, D, and E, occur. The greatest amplitude is at D, which coincides with a small island in Lake Cowan. It seems likely that the magnetic anomalies might again be due to weathered ironstone cappings. If this is so, the anomaly of greatest amplitude could well mark an ironstone capping in outcrop, and the other three anomalies might arise from ironstone cappings now buried under lacustrine sediments.

The nature of the source of this semicircular magnetic feature warrants some investigation. It could be due to an untrabasic flow or a meta-jaspilite bed within the Desirable Pillow Lava, although no such beds have been described in the

lithology of this formation. Another possibility that could be of economic interest is that the source of the magnetic feature is a zone of sulphide mineralization emplaced into the nose of the major anticlinal structure. This possibility could be investigated by inspecting the island on Lake Cowan that coincides with anomaly D.

The most important formation economically, is the Crown-Bluebird sill sequence. From the lithological descriptions, a magnetic anomaly might be expected over this formation, but no such anomaly was detected by the regional survey. From the trend evident in Plate 1, the Crown-Bluebird sill sequence might cut across the north-west corner of the detailed survey area, but owing to practical limitations in the use of the magnetometer (see Appendix), a low amplitude magnetic feature striking north-east to east across the area would probably not have been detected in the detailed survey.

The Mararca Pillow Lava and Kingswood Basalt are no more evident on the magnetic profiles than the other beds within the Wollyeenyer Group. Much of the contour map for the detailed survey is almost featureless, which would suggest that a large part of the detailed survey area (i.e. the north-western part, except for the extreme north-west corner) consists of the Woolyeenyer Group.

The large amplitude magnetic anomalies in the southern and south-western parts of the detailed survey area are attributed to the Noganyer Group. From a comparison of Plates 1 and 3 it is obvious that the Noganyer Group is strongly magnetic. There appears to be a discrepancy between Plates 1 and 3 where the magnetic data north of the Jimberlana Dyke indicate that the Noganyer Group extends half a mile to the north-west of the geological boundary as mapped in Plate 1.

South of the Jimberlana Dyke, the profiles of the regional survey show a north-trending magnetic 'high' at about longitude 121 54. This is interpreted as being due to the Noganyer Group on the eastern limb of the major structural anticline, where it has been hybridized by the Eastern Granite. To test this proposition the large amplitude anomaly F should be investigated as there may be a xenolith of the Noganyer Group within the Eastern Granite at this point.

It is envisaged that the nose of the anticline in the Noganyer Group has been replaced by the Jimberlana Dyke, the magnetic 'high' at G representing a highly contorted off-shoot possibly having a tight fold of 180° in the bedding.

This interpretation implies that the Penneshaw Beds form the core of the anticline with the Eastern Granite contact falling along the axial plane of the anticline. The Penneshaw Beds are magnetically featureless on the regional contour map and are not within the area of the detailed contour map.

The area to the east of the Penneshaw Beds and south of the Jimberlana Dyke in Plate 3 is presumably mainly of Eastern Granite. North of this area and north of the Dyke, the geology is presumably a repetition on the eastern limb of the major anticline of the geology to the west. For example, the area of high magnetic field in the east of the detailed survey area may be due to a repetition of the Abbotshall Beds. There is an east-west lineation near Buldonia, which cannot be explained at present and the source of this feature may make the eastern limb of the anticline more contorted than the western limb.

Reference has been made above to the prominent anomalies over the Mount Thirsty Beds on the western limit of the anticline. The large amplitude anomalies near Buldonia, which strike north-west across the north-eastern corner of the area shown in Plate 3, could be the expression of part of the Mount Thirsty Beds occurring on the eastern limit of this anticline.

The Jimberlana Dyke is obvious in Plate 3 as a large amplitude anomaly which strikes east across the anticlinal structure described above and which corresponds approximately with the position mapped in Plate 1.

#### 4. CONCLUSIONS

The magnetic characteristics of the various geological units have been determined from a comparison of the regional aeromagnetic results with the known geology. Also, from the regional aeromagnetic results, the major structure is seen as an anticline pitching flatly north with the axial plane at about the position of the contact between the Penneshaw Beds and the Eastern Granite. The eastern limb of the anticline, north of the Jimberlana Dyke, appears to be more contorted than the western limb. In particular, there is an east-west lineation near Buldonia, which cannot be explained at present. Geological mapping in this eastern section is needed to clarify the situation.

A semicircular magnetic feature in the Desirable Pillow Lava north of the peninsular that contains the Princess Royal Mine should be investigated further. It may be due to sulphide mineralization in the nose of the major anticlinal fold. As a first step, the island coinciding with anomaly D should be examined geologically to attempt to determine the nature of the underlying rock. It should be possible to delineate this feature more completely by means of detailed gravity and aeromagnetic surveys.

The detailed aeromagnetic survey was not successful in assisting the geological mapping within the survey area. It may have been far more useful if the survey area had included the peninsula that contains the Princess Royal Mine, as the Crown-Bluebird sill sequence may then have been detected and the known geology could have been used as a basis for the interpretation of the remainder of the survey area. Unfortunately, the geological map shown in Plate 1 was not available at the time of planning this survey. Also, owing to practical considerations, it may not have been possible to detect the low amplitude anomaly expected over the Crown-Bluebird sill sequence.

The source of the large amplitude anomaly F should be investigated because it may be due to a xenolith of the Noganyer Group within the Eastern Granite. If this were so, it could be confidently proposed that the eastern limb of the Noganyer Group has been hybridized by the Eastern Granite and is now represented by a magnetic 'high' along longitude 121 54' E.

# 5. REFERENCES

- BMR 1965 Maps showing the results of an airborne magnetic and radiometric survey of the Norseman 1:250,000 area. <u>Bur.Min.Resour.Aust.Rec.</u> 1965/36.
- CAMBELL, W.D., 1904 The geology and mineral resources of the Norseman district. Geol. Surv. W. A. Bull. 21.
- ELLIS, H.A., 1953 Norseman Gold Mines N.L. In GEOLOGY OF AUSTRALIAN ORE DEPOSITS. Aust. Inst. Min. Metall.
- HALL, H.I.E., and BEKKER, C., 1964 The gold deposits of Norseman, Western Australia. Paper presented to 1964 Annual Conference of the Australian Institute of Mining and Metallurgy.
- O'DRISCOLL, D., 1953 Operations on the Norseman Field. <u>In</u>
  GEOLOGY OF AUSTRALIAN ORE DEPOSITS. <u>Aust.Inst</u>.
  Min.Metall.

## APPENDIX

## Operational details

## Survey specifications

Detector altitude

Line spacing

Line direction

Recorder sensitivity

Diurnal correction

250 feet above ground level

One tenth of a mile between

adjacent flight lines

Parallel to Eyre Highway, that is,

approximately 077°

1st recorder, 20 gammas per inch

2nd recorder, 2000 gammas per inch

Applied correction rounded off

to the nearest gamma

Equipment

Aircraft

Magnetometer

Recorder

2 x Mosely Autograph

l x Texas Instruments (two-channel)

Camera

Modified Vinten frame, 35 mm, 186°

BMR proton precession type MNS1

field of view

Cessna 180

Radio altimeter

AN/APN-1

:

#### Method

A correction for diurnal variation was determined by flying a baseline at the beginning and end of each survey flight. The mean level of the magnetic field readings obtained along the baseline was compared with a standard level to determine the correction. The diurnal correction was applied on the assumption that the diurnal magnetic field varied in a linear The diurnal correction was applied on the manner during any one survey flight. No total mrecorder was available to check this proposition. No total magnetic field

Notification of magnetic storms was given by BMR Mundaring Observatory.

The airborne magnetometer records accepted for survey data showed a noise envelope of approximately 20 gammas. noise level limited the detection of low amplitude anomalies. Also, it imposed a limitation on the direction of the flight lines, which had to be in an easterly direction for minimum noise, whereas flight lines in directions between north-west and north would have been preferred.

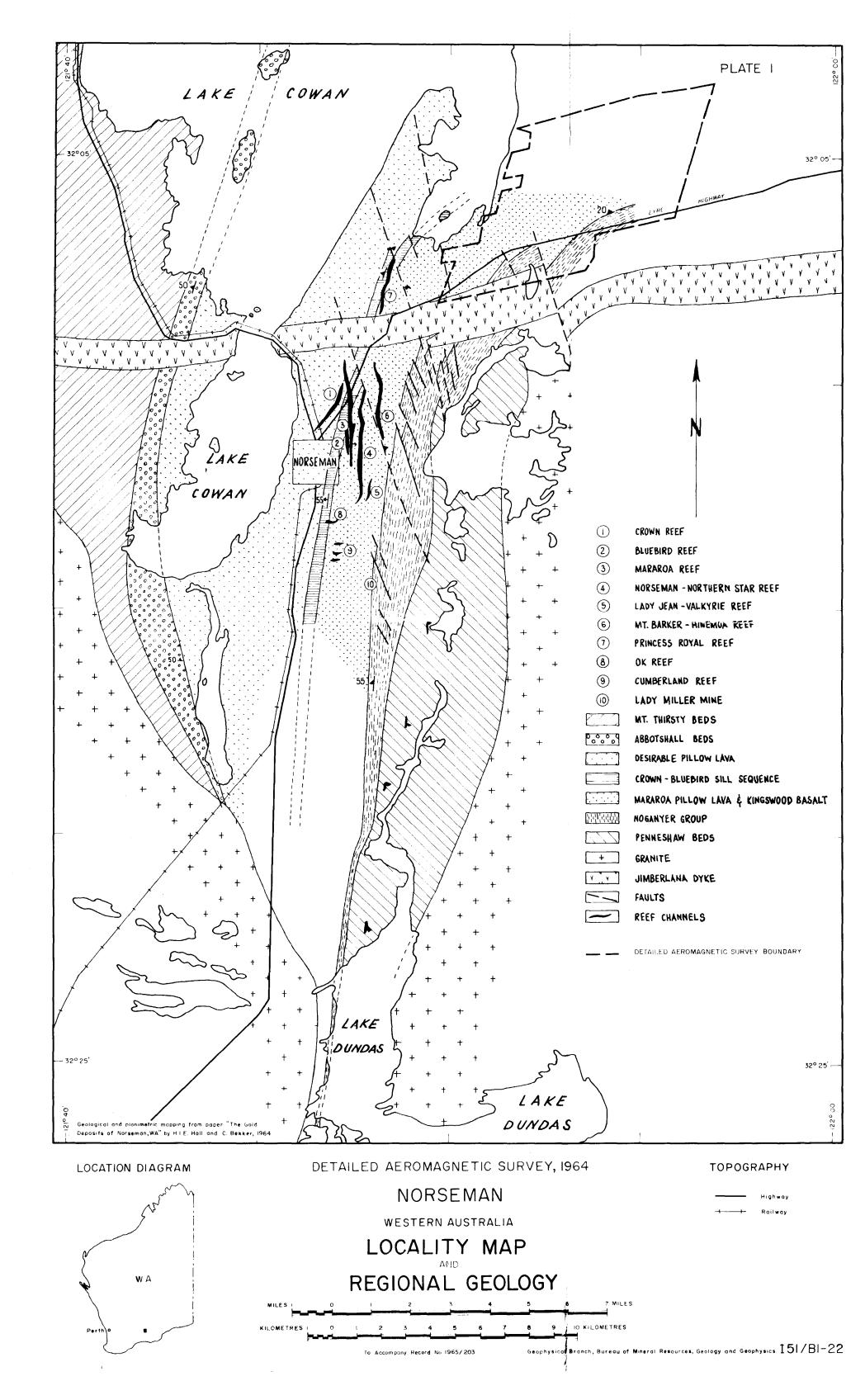
No samples of rocks were collected during the survey. no measurements were made of either magnetic susceptibility or remanent magnetization.

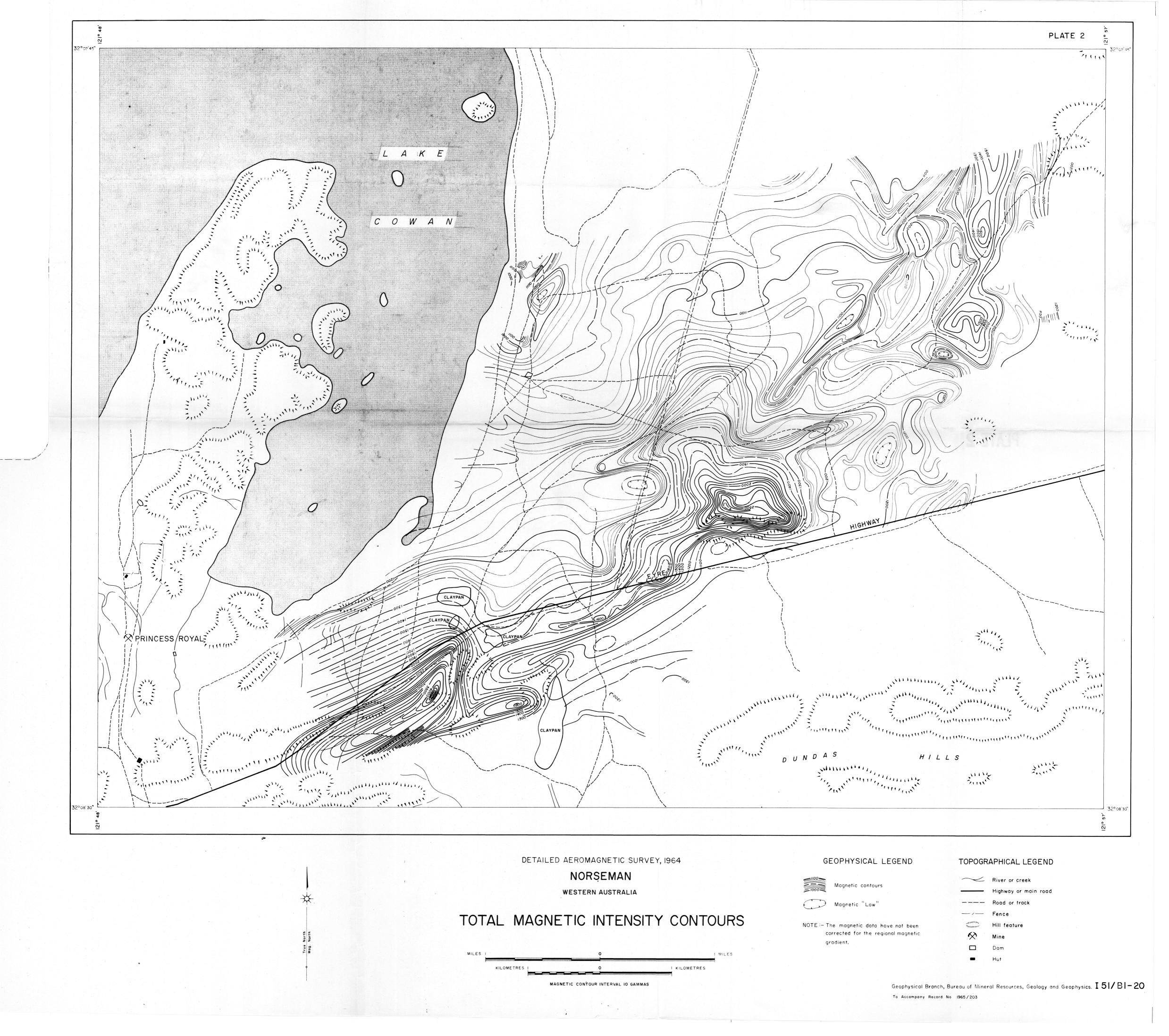
#### Personnel

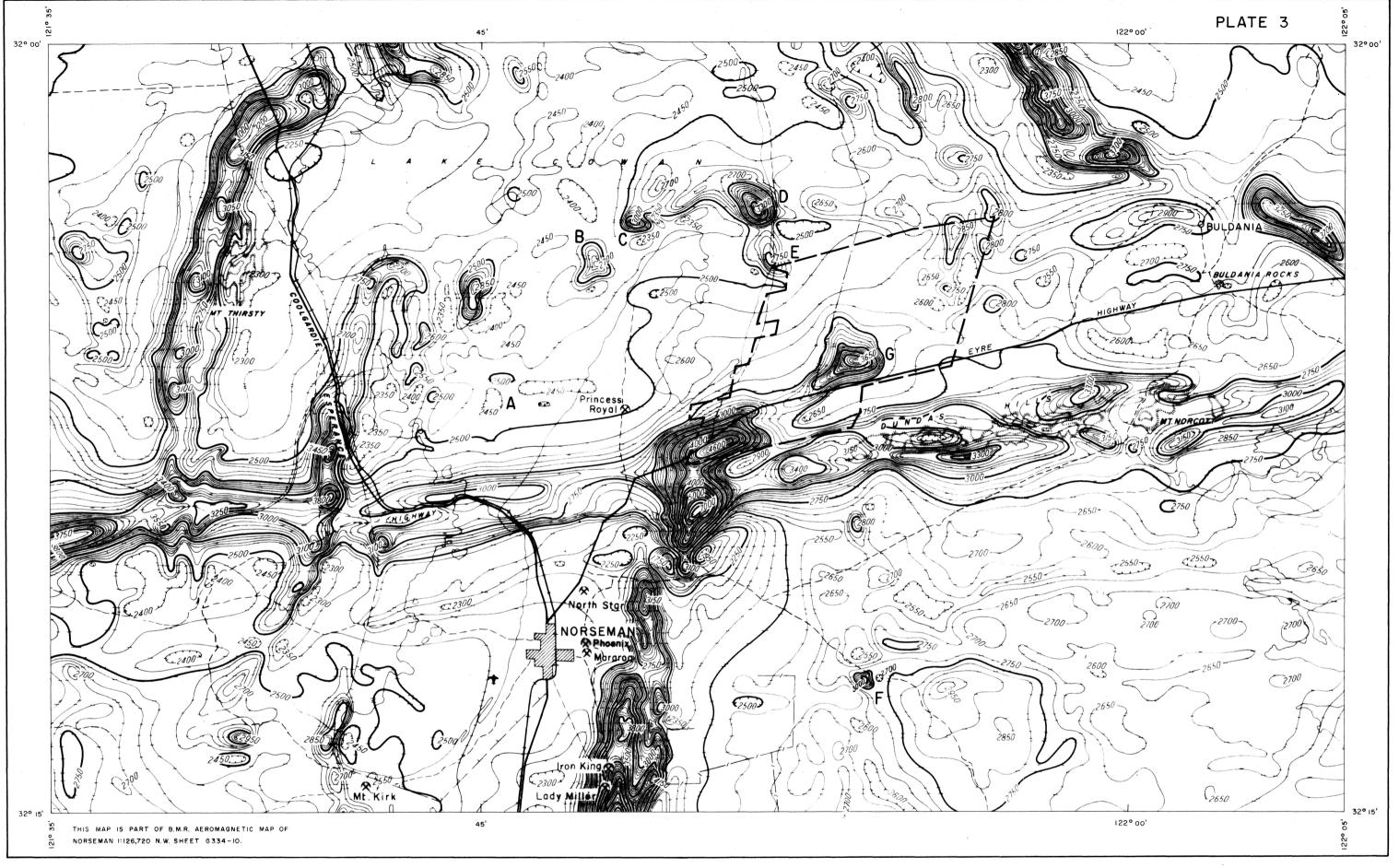
Personnel engaged in the survey were:

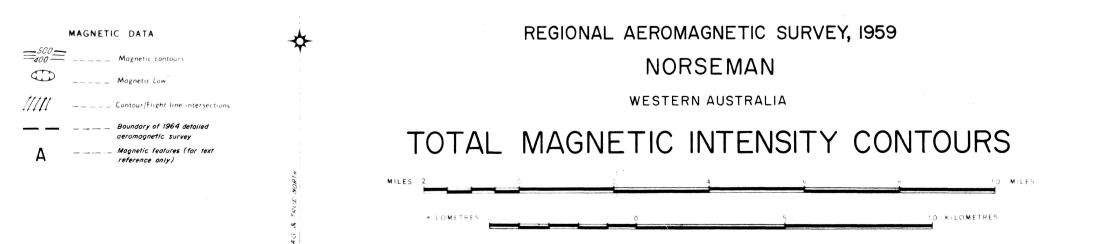
BMR: B.A. Dockery, W.A. Finney, A.S. Scherl, P. Zerial C.J. Braybrook, B. Tregallas.

First Officer G.B. Litchfield.









River or Creek

Highway or Main road

Road or Track

Railway

Fence

Rock outcrop

Named place

Hill feature

Aerodrome or landing growth

TOPOGRAPHICAL DATA

Geophysical Branch, Bureau of Mineral Resources, Geology and Geophysics. I51/B1-21

To Accompany Record No 1965/203

MAGNETIC CONTOUR INTERVAL 50 GAMMAS