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STRANGWAYS RANGE AEROMAGNETIC SURVEY,

NORTHERN TERRITORY 1965 .

007542

by

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RESTRICTED

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SUMMARY

Between mid-September and mid-November, 1965, the Bureau of Mineral Resources flew a detailed aeromagnetic survey of five areas in the Strangways Range to the north and north-north-east of Alice Springs. A total of 115 square miles was surveyed with the primary purpose of outlining possible extensions to known copper and apatite deposits and any anomalous areas nearby. A second aim of the survey was to assist geological mapping.

This Record outlines the regional and local geology and briefly reviews previous investigations in the general area before presenting and analysing the aeromagnetic data. The operational details of the survey and an outline of the interpretation methods employed are given as Appendices.

Of the sixteen known mineral occurrences, five have strong magnetic expression, three others are probably correlatable with two magnetic anomalies, and a further two occurrences are very close to a fault interpreted from the magnetic data. Further ground investigation is recommended in the vicinity of six of the deposits. Many other anomalies were recorded in all areas, and the majority are due to bodies either outcropping or within 50 feet of the surface. Many, if not most, of these bodies are believed to be steeply dipping lenses and prismatic bodies of amphibolite interjacent with less magnetic gneissic and schistose rock. Ground investigation is necessary to differentiate between mineralised and non-mineralised magnetic bodies.

Thirty-one faults have been interpreted, many of which coincide with photo-linear features, and three fold structures have been tentatively proposed.

1. INTRODUCTION

Between 20th September and 16th November, 1965 the Bureau of Mineral Resources flew a detailed aeromagnetic survey of parts of the Strangways Range in the Northern Territory. The region surveyed, totalling 115 square miles, comprises five small areas around the fringe of the ranges. These areas, shown in Plate 1, extend from the northern part of the ALICE SPRINGS 1:250,000 area into the ALCOOTA area.

The survey was requested by the Bureau's Resident Geological Staff in Alice Springs. The primary purpose of the survey was to outline the probable extensions of known mineralised deposits and any anomalous areas nearby. In Areas 1, 2, 4 and 5 limited copper mineralisation has been found associated with magnetite, and in Area 3, three mineral deposits are known, each consisting of magnetite, apatite and zircon. A second aim of the survey was to assist geological mapping, especially where the solid geology is obscured by Quaternary alluvium.

Almost all previous investigations involved the mapping or study of a large region, of which the aeromagnetic survey areas constitute only small parts. In a paper on the geology of the Northern Territory, Hossfeld (1954) described the metamorphics of the Arunta Complex which he divided into the older Arunta Series and the younger Riddock Series, the latter forming the Harts Range to the immediate east of the survey area. The first comprehensive regional map was made by Joklik (1955) during his survey of the Harts Range. His more detailed mapping, involving the subdivision of Hossfeld's Riddock Series, did not extend westwards to include the Strangways Range. During 1957 and 1958 the Commonwealth Scientific and Industrial Research Organisation made a regional study of an extensive part of the Northern Territory including the Strangways Range. As part of this investigation Quinlan (1962) reported on the geology, Ryan (1962) on the mineral occurrences, and Mabbutt (1962) on the geomorphology. Quinlan has also described the regional geology of the Alice Springs area in a Bureau Record (Quinlan, 1959). The greater part of the HERMANNSBURG, ALICE SPRINGS and ILLOGWA CREEK areas has recently been mapped by Forman, Milligan and McCarthy (1965).

More detailed mapping of the five survey areas has recently been commenced by the Bureau's Resident Geological Staff in Alice Springs. By photo-interpretation supported by ground investigation, the rocks have been broadly grouped into lithological sub-divisions of the Arunta Complex and geological strikes and faults have been delineated. All known mineral occurrences have been inspected in the field. The geology shown in the plates which accompany this report, and the geological descriptions of the five areas, are based on

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this work. Previous workers have examined some of the mineral deposits; Owen (1944) and Jensen (1944) reported on the occurrence of apatite in Area 3, Bell (1953) described the Pinnacles copper workings in Area 4, and Thomson (1950) reported on the copper exposure in Area 5.

In 1961 a helicopter gravity reconnaissance survey was made of much of the Northern Territory. The data obtained from the Amadeus Basin area have been presented as a Bureau Record (Langron, 1962). As this was a reconnaissance survey the gravity pattern is correlated with broad regional features and the degree of resolution is insufficient to be of great assistance to the detailed aeromagnetic survey. A ground magnetometer survey was made in 1965 by Geopeko Ltd. over a chalcopyrite-magnetite deposit 4,000 feet north-west of Southern Cross Bore in Area 4. The results of this survey are not known.

An airborne magnetic and radiometric reconnaissance survey of the Amadeus Basin was flown in 1965 (Young and Shelley, 1966). The magnetic field was found to be highly disturbed in the northern half of the ALICE SPRINGS area, where many anomalies of over 2,000 gammas were detected. Most of the radiometric data were correlated with broad geological features.

2. GEOLOGY

Regional geology

The five survey areas represent parts of the Arunta Block composed of ancient metamorphics of the Arunta Complex. The rocks of the Strangways Range originated as sediments with some interbeds of volcanics. During at least two periods of regional metamorphism these rocks were altered to metamorphosed calcareous sediments, amphibolites, quartzo-feldspathic gneiss and garnet-mica-feldspar gneiss with some bands and lenses of quartzite, (Youles, 1965). Various schists have also been found in subordinate amounts. Following the procedure adopted by the geologists in Alice Springs, the term 'amphibolite' is applied throughout this Record to all unspecified basic igneous rocks known to be metamorphosed. These occur as lenses, plugs, sills and flows. Small isolated outcrops of ultra-basic intrusives are also known.

At least two orogenies are evident in the north-eastern margin of the Amadeus Basin and there is evidence to suggest that both were preceded by an earlier period of folding and metamorphism. The Arunta Orogeny pre-dated the deposition of the Amadeus Basin sediments and caused the deformation and moderate to high grade metamorphism of the Arunta Complex. The Complex was isoclinally folded about north-south axes and was later tightly refolded about steeply dipping east-west axes. The later Alice Springs Orogeny caused recumbent folding of the older rocks, possibly with some metamorphism in the Strangways Range area, (Forman, Milligan and McCarthy, 1965). Faulting, mostly of unknown age, is common throughout the Arunta Block.

Detailed geology

Area 1. This area which is mostly soil-covered has a few scattered outcrops of medium-grained amphibolite, the foliation of which strikes at 070° and dips at 55° northwards (Plate 2). East of Redrock Bore an outcrop of banded quartz-garnet-magnetite rock has been mapped trending 130° and dipping very steeply north-eastwards. The rock consists of narrow alternating bands of quartz and quartz/magnetite with scattered garnet throughout.

Two mineral deposits, each of magnetite and chalcopyrite, have been found in the centre of the area. The minerals occur in veins which are up to 300 feet long and 10 feet wide, and whose strikes appear to be parallel to the regional trend of the foliation. There are also substantial areas where magnetite floaters occur. Weak copper mineralisation has also been observed in a shear at the contact of the amphibolite-rich zone and gneiss to the north.

Area 2. Almost all the area is of gneiss and amphibolite, the latter mostly occurring as bands concordant with the gneiss (Plate 4). These bands are generally 10 to 30 feet wide, too narrow to be marked on the geological map. The strike of the foliation is commonly northerly but becomes eastwest in the extreme north and south. Many amphibolitic and ultrabasic intrusives, some with associated magnetite, also have exposures too small to be shown. Faulting of variable strike is common.

A copper exposure in the central part of the area is a secondary ironstone body 500-600 feet long and up to 60 feet wide, which strikes due north and dips at 55° east, parallel to the regional trend in the area. The body occurs in metamorphosed limestone. Secondary copper minerals are prominent and smaller amounts of galena and haematite have been noted. It is not known whether the mineralisation is associated with nearby basic intrusives. At the two copper localities in the north-east of the area, secondary copper minerals and limited chalcopyrite have been found in a basic intrusive elongated north-south, parallel to the strike of the country rock.

Area 3. Three major rock groups within the Arunta Complex have been recognised in this area (Plate 6). The most important of these economically is a metamorphosed limestone in which three mineral deposits are known, each consisting of magnetite, apatite and zircon. The apatite is believed to have formed by the pneumatolytic replacement of limestone by fluoride-rich fluids. The three deposits occur in a belt, one mile long and 500 feet wide, trending north-east. The apatite is present throughout the limestone probably in amounts of the order of 10%, and its tonnage is estimated to be 200,000 tons per vertical foot. The attitude of the deposits is not yet known with certainty because of the masking effect of Tertiary weathering. At the eastern outcrop, however, the limestone is overlain, probably conformably,

by quartz-muscovite-schist which trends north-north-east and dips at 70° to the west-north-west. In addition, from photo-interpretation supported by ground investigation, the probable strikes of the central and western outcrops are 045° and 085° respectively, and the dips are steep in a northerly direction. Small bodies of pegmatite, consisting almost wholly of feldspar with minor quartz, have intruded the limestone. The largest outcrop measures 24 feet by 10 feet. Small masses of phlogopite have also been seen in the limestone.

The remaining two rock groups consist principally of amphibolite and gneiss. The division is based on the relative amounts of these two components, and there is therefore, a gradation between the two groups. The trend of their foliation is variable. In the extreme south this is easterly; in the north-west there is a swing through north-east to north, whilst in the south-east of the area the swing is towards the north-west. Faulting is common, the dominant trends being north-west and west. North-east faulting has also been observed.

Area 4. The rocks of this area have been divided into two broad lithological groups, one of gneiss, amphibolite and schist and the other of marble and hornfels (Plate 8). The two groups are generally well-defined and only in the south-east is there any evidence of gradation. The strike of the foliation is very variable, the major changes being either side of fault planes. The long meridional fault in the central west of the area is a good example; to the west, the trends are north-north-west to north-west whereas on the eastern side they swing from north-north-east to east. In the north-east of the area, the trends are north-north-west, becoming westerly across a fault plane. From the curvature of the trend lines shown, folding has obviously occurred; the major fold axes are believed to strike generally north whilst minor cross-folding is thought to exist throughout the area.

Copper mineralisation has been observed at a number of localities. Half a mile north-west of Southern Cross bore, massive magnetite-chalcopyrite has replaced a lens of metamorphosed limestone occupying an area 400 feet long and 20 feet wide. The magnetite is concentrated in the noses of steeply plunging folds along the contact with gneissic country rock. A small basic boss, heavily stained with secondary copper minerals, has intruded the limestone. The other deposits, including the Pinnacles workings, are found north-east of Southern Cross bore and are veins of bornite, ankerite and quartz in marble. The mineralisation occurs in a belt about 1,000 feet wide and two miles long trending north-north-west approximately parallel to the strike of the country rock. Included within this belt are two ironstone capped hills each about 500 feet long and 10 to 20 feet wide. It is th ought that the ironstone may have been derived from sulphide deposits but its significance is not understood. Secondary copper ore has been won from an opencut on the hanging wall of the easterly of these bodies.

Area 5. The whole area is shown as undifferentiated Pre-Cambrian rocks either outcropping or beneath Cainozoic cover (Plate 10). Three faults strike north-west to west-north-west, and there is evidence of weaker faulting striking north-east. By photo-interpretation the rocks north of the central fault and those south of the southern fault are believed to be similar, consisting probably of granitic gneiss with banded amphibolite. Between the two southern faults the rocks have different photo-characteristics and are thought to be predominantly schist with some intrusive (?) amphibolite. Tight isoclinal folding about north-west trending axes occurs west and east of the map area, and is believed to be present also within the survey area. There are large open folds and domes of variable trend in the north and east.

A copper prospect in the north-central part of the area has been examined by Thompson (1950) who stated that quartz-magnetite-haematite lodes up to 14 feet wide occur in altered sillimanite-sericite-schists. The magnetite and haematite are found either as disseminated crystals, or with quartz in more clearly defined lode channels. Lodes of both types crop out for approximately 1,300 feet along a 115° strike and dip of 80° northwards, and they appear to be closely associated with amphibolite sills. The only surface copper mineralisation is sporadic and weak carbonate staining. A pit 12 feet deep has been dug in the centre of the lode. Assay values of 0.5% and 1% copper were obtained at the 5 feet and 9 feet levels respectively.

3. RESULTS AND INTERPRETATION

The aeromagnetic data are presented in Plates 2, 4, 6, 8 and 10 as contour maps of total magnetic intensity. Plates 2, 4, 6 and 8 are at a scale of 1:22,500 whereas the scale for Plate 10 is 1:24,000. A basic contour interval of 50 gammas has been used for Areas 1, 2, 4 and 5. In Area 3 the relatively low degree of magnetic disturbance, in the eastern part particularly, has allowed for contours at a 10 gamma interval. Corrections have been made for the diurnal variation of the total field but the earth's regional gradient has not been removed. All intensity values are relative only and are referred to a single arbitrarily chosen datum.

Plates 3, 5, 7, 9 and 11 show the qualitative interpretation of the magnetic data in the form of magnetic trend-lines both positive and negative, interpreted faults, and zones of magnetic distribunce. The geological significance of these magnetic features is discussed for each area. The zoning is based on differing overall magnetic patterns, and although the zones have geological significance, no implication is made that the zonal boundary exactly reflects a lithological boundary. Positive and negative trend-lines represent the axes of magnetic 'ridges' and 'troughs' respectively.

Generally a magnetic body produces a magnetic high with an adjacent magnetic low. It should be noted at this stage that the apical axis of an inclined prismatic magnetic body does not generally coincide with either a positive or negative trend-line but is usually displaced to the negative side of the positive trend-line.

Selected anomalies have been quantitatively analysed to provide information on the depths of burial and widths of the causative magnetic bodies. All areas surveyed show a considerable degree of complexity in the magnetic pattern. The majority of the anomalies are themselves complex, being the resultant of the magnetic contributions from more than one source. Consequently, the number of anomalies suitable for accurate analysis represent a small percentage of the total. The contour maps show a series of lines, A-A', B-B' etc. where magnetic profiles were constructed from the contours, normal to the magnetic strike of the anomalies. The data obtained by analysing the constructed profiles are given for each area. Some of the selected anomalies have been further analysed to yield dip angles of bodies and apparent magnetic susceptibility contrasts. The calculated depths are considered accurate to + 15%. The calculated widths of bodies can in general be considered to represent maximum values. The magnetic bodies need not necessarily be homogeneous, but may consist of a series of bodies sufficiently narrow and closely spaced that the magnetometer was unable to even partially resolve their individual magnetic contributions.

Area 1

This area divides naturally into three zones of contrasting magnetic character (Plate 3). Zone A in the north is represented by a major complex anomaly with a maximum amplitude of 2600 gammas. The anomaly is partially resolved into at least two simpler anomalies but its northerly continuation was not covered by the survey. Zone B is characterised by a series of subelongate anomalies with an average amplitude of 100 gammas although larger anomalies of up to 700 gammas were detected. Zone C in the south contains two major anomalies of 4600 gammas and 2250 gammas. Almost all magnetic trends lie between east-north-east and east-south-east. The trends in zones A and C are biased to the north and south respectively of due east. Only in the extreme eastern part of zone B is there a large divergence from east-west.

Data derived by quantitative analysis of constructed profiles (Plate 2) are tabulated below :-

Section	Depth of magnetic body below ground level (in feet)	Half-width of magnetic body (in feet)
A-A 1	200	400
B-B1	250 approximate	1200 maximum
C≟D	150	600
D⊷E	50 to 100	500
मृ_्म ।	500 maximum	1400 maximum
GG'	500 maximum	1600 maximum
H-H ,	100	400
J_J'	0	400
K-K 1	50	500
L-L	50	600
M-M	0	700
N-N *	200	200
P-P†	0	100
Q-Q'	less than 50	100

The limited known geology affords some control on the interpre-The amphibolite-rich zone (use of 'zone' here tation of zonal significance. is geological) in the south-east of the area is clearly reflected by an anomaly of 700 gammas crossed by J-J' (Plate 2). Although the contours suggest a single anomaly, the source-body is probably a series of parallel amphibolitic bands too closely spaced to be resolved. Similar anomalies in zone B, e.g. those crossed by N-N', P-P' and K-K', are also expected to reflect amphibolitic bands and lenses. The generally smaller amplitudes that that of J-J are partially due to the greater depth of burial of the magnetic bodies, but are also though to signify smaller percentages of amphibolite relative to interjacent strata, probably gneiss. Some of the anomalies in zone B have been analysed to give the apparent susceptibility contrast between the magnetic bodies and the surrounding rock. The average value obtained, of the order of 2.0 x 10^{-3} c.g.s. units, is consistent with the contrast expected detween amphibolite and gneiss. As most of the magnetic trends in znne B are generally parallel, both with respect to each other and to geological strikes, the metabasics are concordant with the meta-sediments and were therefore originally lava-flows or sills. Their attitude is near-vertical.

The largest anomaly in zone C was detected over an outcrop of banded quartz-garnet-magnetite rock which strikes parallel to the anomaly trend.

The known outcrop, measuring 750 feet by 1300 feet, appears to extend for a considerable distance beneath Cainozoic cover; the dimensions of the magnetic body causing the anomaly crossed by L-L⁰ and M-M⁰ are slightly more than one mile long and approximately 1400 feet wide. Its dip has been calculated to be 80° to the north-east. This rock-type is probably responsible for all anomalies in this zone although those crossed by sections C-D and D-E are lower in amplitude. In the case of the former anomaly this can be partially explained by the increased depth of burial of the causative body. The anomaly crossed by D-E is due to a slightly narrower body, and its amplitude is possibly reduced also by the negative effect of the magnetic bodies to the north.

The complex anomaly in zone A is due to at least two magnetic bodies up to 500 feet below ground level. If these were outcropping, the resultant anomalies would have amplitudes as great as those detected in zone C, and it is more than possible that the magnetic bodies in both zones are of the same rock-type. This could signify a fold structure about an east-west axis such that zones A and C represent the two limbs of the fold. This concept is supported by a curvature of both magnetic trend-lines and known geological strikes. In the eastern part of zone B, the curvature is particularly noticeable and in the centre of the area between faults f_1 and for the positive trend-lines appear to form a closure to the east. Assuming the depth determinations in zone A to be reasonably accurate, the suggested fold would have to be synform. However, the dip directions shown on the geological map in the vicinity of section P-P' (Plate 2) imply that the fold is antiform, but this could be a local fold within a major synclinorium. This fold concept is only tentatively proposed and further mapping is necessary to test its validity. Unfortunately the anomalies in zone A are insufficiently complete to enable reliable dip-angles to be determined, but a southerly dip is indicated.

Three faults have been interpreted from the magnetic data. Fault f₁ is based on a displacement of magnetic trend-lines, and the difference in strike and depth of burial between the bodies causing the anomalies crossed by L-L' and C-D. Both of the known copper prospects occur very close to a second fault f₂ interpreted from a displacement of magnetic trend-lines and the co-linear termination of a number of anomalies. The south-eastern part of this fault, and its continuation south-east of the survey boundary, is co-incident with a strong photo-linear feature. Further ground investigation should be made where the interpreted fault crosses the large outcrop south-east of the southern copper exposure, to test for faulting and mineralisation.

Any extensions to the known deposits are expected to occur along the fault plane. The northern deposit is also close to, and approximately parallel to, the anomaly crossed by Q-Q'. The significance of this will also be known only by ground investigation, but the anomaly could be due to a basic intrusive perhaps associated with ore-genesis. Fault f₃ is interpreted from a change in magnetic trenddirections, although folding could also account for the observed pattern.

Area 2

Four magnetic zones have been delineated in Area 2 (Plate 5). Zone A in the north-west includes a number of ovate anomalies of high amplitude and variable trend. A similar pattern is seen in zone B in the north-east. In the extreme south of the area, zone C represents the northern part of a third region of high magnetic intensity. In the remainder of the area, designated zone D, relatively low-amplitude anomalies of various shapes, trends and amplitudes constitute a complex magnetic pattern.

Data derived by quantitative analysis of constructed profiles (Plate 4) are shown below. Depth calculations indicate that all magnetic bodies either outcrop or lie within 50 feet of the surface.

Section	Half-width of magnetic body (in feet)
A-A	130
B-B*	150
C-C'	250
D-D'	400
E-E:	350
P_F'	150
G.=G •	350 approximate
H-H.	300 approximate
J_J'	350
KK *	200
L-L'	200
MM *	200
N=.N •	200 approximate
P⊷P†	200 approximate
Q-Q'	300 approximate
R-R	200

The whole area is mapped as amphibolite and gneiss, either outcropping or beneath a veneer of alluvium. Individual bands are too narrow
to be shown on the geological map. For these reasons it was particularly
necessary to attempt to correlate the magnetics with features on aerial photographs, and to use any observable correlation in the further interpretation
of the magnetic pattern. The general agreement was found to be good.

Zone A coincides approximately with a dark area on the photographs which differs from surrounding country in topography and vegetation. As is known from geological mapping, this dark area is faulted at its southern, north-eastern and eastern contacts; the southern boundary fault coincides with f_5 , a fault interpreted from a change of magnetic trend directions. The western boundary of the zone is clearly seen on the photographs as an abrupt change in rock-type. This contact could be faulted, but the magnetic evidence is insufficient for the fault to be shown in Plate 5.

The abrupt variations in magnetic trends within the zone signify that other faulting has occurred but they do not indicate the positioning of the faults. All magnetic trends are parallel to known geological strikes. The individual anomalies crossed by G-G', H-H' and J-J' are due to lenses or irregularly shaped masses of basic material, each with a maximum width of 600 to 700 feet. The magnetic susceptibility contrast between the source rock of the anomaly crossed by G-G' and the rock to the west and east, has been calculated to be 3.5 x 10⁻³ c.g.s. units. This is approximately equivalent to a change in magnetite content of 1%.

Zone B, similarly correlated with an area of dark hilly country, is confined by a series of faults shown on the geological map (Plate 5). The southern boundary fault coincides with f8 a fault interpreted from a change in magnetic trend directions. Trend variations within the zone have enabled two other faults to be interpreted. North-east of fault fin a number of anomalies trend north-east. That crossed by Q-Q' is due to a body, parallel to the geological strike, which has a maximum thickness of 600 feet and minimum length of 3000 feet. The rock type is probably meta-basic and similar to that in zone A. West of fault f10 two copper localities are known where secondary copper minerals have been found in a basic intrusive elongated north-south. The meridional anomaly crossed by sections M.M., N.N. and P.P. is almost certainly due to this basic intrusive. The anomaly, which coincides with a steep ridge seen on the photographs, is due to a north-south body calculated to have a maximum thickness of 400 feet. The calculated length of 2900 feet is exactly the length of the ridge. The area south of the known exposures is considered worthy of detailed ground investigation. On magnetic evidence the north-north-west trending anomaly immediately to the south-east could also be due to a similar basic intrusive, but this is not supported by any similarity in topographic form. West of fault fo the magnetic trends revert to a north-easterly direction and the anomalies are probably due to the same rock type as that in the north-east of the zone.

The boundary of zone C, marking the approximate limit of a third region of high magnetic intensity, coincides with a photo-linear feature clearly representing an abrupt change in rock-type. The anomalies in zone C have similar amplitudes to those in zones A and B/are probably due to meta-basics. The interpreted fault f6, coinciding with a known fault, separates north-west and north-north-east trending magnetic bodies.

Zone D, comprising the remainder of Area 2, is interpreted as a much-faulted region composed of at least two rock types, probably amphibolite and gneiss, the proportion of amphibolite being lower than in the other zones. A number of anomalies e.g. those crossed by section F-F', K-K' and L-L' may be directly correlated with dark bands on the aerial photographs. It is expected that most elongated anomalies trending parallel to the geological strike are concordant bands and lenses of amphibolite interjacent with gneiss. Anomalies which cross obliquely to the geological strike are probably basic intrusives occurring as dykes. Two examples of oblique anomalies are seen south of the thrust fault in the centre of the area.

There is no obvious correlation between the magnetic data and the copper exposure in the central part of the area. However, the anomaly immediately to the east could be due to a basic intrusive, possible associated with ore genesis. Zone D includes a number of interpreted faults many of which merely serve to confirm faults already deduced by photo-interpretation. Most of those interpreted from the magnetic data separate anomalies of conflicting trend directions. The exception is f2 which causes the displacement of a number of magnetic trend-lines. This fault is also seen as a photo-linear feature.

Area 3

Four magnetic zones have been delineated in this area (Plate 7).

Zone A has a complex magnetic pattern formed by anomalies of variable trends, shapes and amplitudes. The dominant trend direction is north-north-east and the average amplitude is approximately 150 gammas, although anomalies of up to 400 gammas were detected. Zone B is a simple magnetic trough which trends 005° and widens towards the north. Zone C includes the three largest anomalies in the area having amplitudes of 750, 1000 and 1800 gammas. The western two trend north-easterly and the eastern one trends slightly east of north. Zone D, forming most of the remainder of Area 3, is characterised by north-westerly trending anomalies of low amplitude superimposed on a gradually changing field. In the south east, zone E encloses an east-west anomaly of 250 gammas.

Data dericed by quantitative analysis of contructed profiles (Plate 6) are shown below. Depth calculations indicate that all magnetic bodies either outcrop or lie within 50 feet of the surface.

Section	Half-width of magnetic
	body (in feet)
A-n:A *	350
BB*	Not known
C and C	500
DD •	700
E-E	500
	300
G_G'	500
·HanaJ	300
J⊷K	150
K=L	460
I:M	250
M-N	200

A general correlation is observed between the zones and known geology. Zone A includes two outcrops of amphibolite with minor amounts of gneiss. Where the geological strike is known, the magnetic trends are near-parallel. Most, if not all, north to north-north-east trending anomalies are therefore interpreted as being due to amphibolitic bands and lenses. Susceptibility calculations on some of the anomalies gave values between 1.7 x 10⁻³ and 2.4 x 10⁻³ c.g.s. units. The north-west trending anomalies probably reflect discordant basic intrusives or their metamorphic equivalent. The complexity of the contour pattern in this zone could be due in part to basic plugs and irregularly shaped masses of basic material. There is no magnetic indication that deposits of magnetite-apatite-zircon, comparable to those in zone C, occur in zone A. Beyong the limit of known outcrop, the thickness of Cainzoic overburden is calculated to be less than 50 feet throughout the zone.

The geological significance of zone B is not fully understood. It clearly reflects a rock of low susceptibility relative to those in zone A and C. Although not shown as such in Plate 7, its western boundary could be a fault contact, parallel to the negative trend-line. Of known rock-types in the Strangways Range, zone B could represent acidic gneiss or a continuation of the limestone in zone C but without the magnetite. The possibility of an acidic intrusion should not however be discounted especially as pegmatite has been observed in zone C to the east. The widening of the trough at its northern part is probably due to the negative effect of a magnetic body to the north in zone A.

Zone C is the most important economically, enclosing the three outcrops of limestone with apatite, magnetite and zircon. Each outcrop is clearly reflected by a positive magnetic anomaly to the Morth, and a negative anomaly to the south (Plate 6). The western anomaly, crossed by F-F', is due to a body 600 feet wide and approximately 1400 feet long trending parallel to, and centred midway between, the positive and negative magnetic trend-lines. The dip of the body is calculated to be $80^{\circ} + 5^{\circ}$ to the north-north-west, and a value of 4.0 x 10⁻³ c.g.s. units was obtained for the magnetic susceptibility contrast between the body and the rocks to the south. The anomaly crossed by E-E' and D-D' is due to a body approximately half a mile long trending northeast. Its width in the south-west is about 1000 feet increasing to 1400 feet to the north-east. The longitudinal axis of the body is approximately midway between the positive and negative trend-lines, and its dip is about 60° to the north-west. The section C-C! crosses the largest anomaly in Area 3 which trends north-north-east. As this is close to due north, the positive part of the anomaly coincides more closely with the causative body. This is calculated to be about 1000 feet wide and a third of a mile long, dipping steeply westwards. The dimensions of the known limestone outcrops and those calculated for the magnetic bodies are compared below :-

Outcrop	Dimensions of known outcrop	Dimensions of magnetic body
Western	1800 ft. x 800 ft.	1400 ft. x 600 ft.
Central	1800 ft. x 500 ft. (average)	2600 ft. x 1200 ft.(average)
Eastern	2000 ft. x 1600 ft. approx.	1700 ft. x 1000 ft. approx.

It seems apparent therefore that not all the limestone outcrop is magnetic. The central outcrop, being appreciably smaller than the calculated dimensions, is the only one where hitherto unknown extensions could be found beneath Cainozoic cover. The extension would be towards the north and northwest. The anomaly in the northweastern corner of the zone could also be due to a similar rock type. The causative body is near-surface and its relatively low amplitude is indicative of a magnetite content well below that of the three major deposits. If the apatite content is found to be approximately proportional to the magnetite content then the smaller anomaly is not considered worthy of further investigation.

Zone D occupies an area known to have a much smaller percentage of amphibolite with a corresponding increase of gneiss. The contact of this group in the east with the more amphibolitic group is believed by geologists to be faulted, (Plate 7). The low-amplitude anomalies in zone D are interpreted as being due to localised concentrations of amphibolite, the strike of which is

north-west throughout. The zone includes the country south of a major negative trend-line in the south-west of the area. The magnetic gradient in this region is positive to the south; this gradient, however, being a part of the anomaly to the north, does not imply an increase in magnetite content towards the south. Nor does it conflict, therefore, with the geological belief that this region represents a transition between rocks predominantly amphibolitic to the north and those predominantly gneissic immediately to the south of the survey area.

Zone E comprises a magnetic anomaly considerably higher in amplitude than any in the surrounding zone D; it is the effect of an east-west body less than 50 feet below ground level. A geologist (Youles, 1965) has inspected the site of the anomaly, where he found an outcrop of amphibolite considerably smaller than the apparent dimensions of the magnetic body; it is probable that it was only a part of a much larger basic body increasing in horizontal extent with depth.

Three faults have been interpreted in Area 3 by analysing trend-line displacements and changes in trend directions. There can be little doubt that the fault f₁ exists. At the fault plane, negative trends meet positive ones, there is an obvious displacement of the major negative anomaly in Zone D, and a general change in trend direction is also apparent. If the relative movement was wholly horizontal, the direction was western block to the north-west. Fault f₂ separates the central and eastern limestone outcrops and is based on geological evidence and the change in anomaly trend direction. Fault f₃ is based solely on the large angle between the trend in zone E and that immediately to the north.

Area 4

Plate 8 clearly shows that the magnetic relief varies greatly from one part of this area to another. As a generalisation the eastern half is considerably more disturbed magnetically than the western half. Seven zones of differing magnetic character have been recognised (Plate 9). Zone A in the north-east is characterised by linear anomalies of up to 2300 gammas. Zones B, D and E in the south-east quadrant have a similar magnetic pattern to that of zone A, and are separated from zone A by a deep magnetic 'trough' striking approximately south-east (Zone C). Zone F in the south-west contains anomalies of about 800 gammas. Zone G, forming the remainder of Area 4, is characterised by relatively low-amplitude anomalies of variable trend.

Data derived by quantitative analysis of constructed profiles (Plate 8), are shown below. All depth calculations indicate that the magnetic bodies either outcrop or lie within 50 feet of the surface.

Section	Half-width of magnetic body (in feet)
A.A.*	700 very approximate
B,B ⁸	500 approximate
C-C1	200
D.=:D 6	, 350
E-E	400
Fr∞F. 8	400
G_G'	650
H~H;	600
J_J '	400
K-K	100
L.⇒L [†]	300 approximate
M—M *	200
N-N •	100
P_P'	100
QQ*	200
R-R°	100 approximate
S-S'	. \$ 300
. Մ	150

Over most of the area a good correlation has been found between the magnetic data, the mapped geology and photo-features, and the interpretation which follows is based upon this correlation. Zone A is characterised by linear, highamplitude anomalies which are generally only partially resolved from neighbouring anomalies. Magnetic trends and known geological strikes vary from one part of the zone to another, but there is general parallelism between the two. The zone largely coincides with an extensive outcrop of gneiss, amphibolite and schist (Plate 9), and the magnetic pattern supports the concept of three dominant rock-types. Calculations to determine the magnetic susceptibility contrast between adjacent rock-types gave values between 4×10^{-3} and 8×10^{-3} c.g.s. units. These figures are compatible with the contrast expected between amphibolite and either gneiss or schist, and the magnetic 'ridges' are therefore believed to be due to bands and lenges of amphibolite, the magnetic 'troughs' reflecting interjacent bands of gneiss and/or schist. Some anomalies, e.g. those crossed by C-C' and D-D', may be directly correlated with parallel dark bands on the aerial photographs. Others, particularly in the north-east of the zone, have less obvious correlation. A number of faults have been interpreted from abrupt changes in the magnetic pattern within the zone. Fault f_1 separates north-west trending anomalies to the south from westerly trending anomalies to the north. Similarly f_3 separates north-west and north-easterly trending anomalies. Faults f_2 and f_4 were interpreted from the co-linear termination of a number of anomalies. All four faults have moderate to strong topographic expression.

The curvature of the magnetic trend-lines east of fault f₅ clearly reflects arcwate rock outcrops of alternating light and dark bands as seen on the aerial photographs; both patterns suggest a truncated fold structure.

Zone B, occupying the south-east corner of the area, has a similar magnetic pattern to that of zone A, being a series of near-parallel, sub-elongate anomalies of high amplitude. However, whereas the highest anomalies in zone A are fairly evenly dispersed, in zone B there is a gradual rise in both magnetic intensity and magnetic 'relief' towards the south-east. A rock-type transition from marble/hornfels to amphibolite/gneiss/schist probably exists in the southeast quadrant of the area (Youles, 1965); the gradation in magnetic intensity supports such a concept. The magnetic 'ridges' are again believed to be due to amphibolitic bands and lenses concordant with the country rock. The magnetic trends are seen to swing from northerly through north-east to east, parallel to the known geological strike, and major folding has obviously occurred in this zone. A number of the trend lines are quite short (down to 400 feet minimum) suggesting that either the amphibolite is very variable in width pinching out altogether in places, or that the magnetite content is very variable. The former alternative is more feasible as both lava flows and intrusive sills can be expected to be reasonably homogeneous along strike.

The northern and north-eastern boundaries of the zone are almost certainly faulted. That part of fault f₈ which forms the northern zonal boundary, separates easterly and north-easterly trending anomalies and it is seen as a very strong photo-linear feature. Fault f₇ forms the contact between the arcuate trend-lines of zone B and the straight north-westerly trends of zone C. This interpreted fault, also seen as a photo-linear feature, is nearly coincident with a positive north-west trend line, and it is possible that a section of the fault plane is filled with magnetic material. Two other interpreted faults are shown within the zone. Fault f₉, interpreted from the co-linear termination of a number of anomalies, is seen as a photo-linear feature, and f₁₀ was interpreted from an abrupt change in mangetic trend-directions.

Zones A and B are separated by zone C, a strong linear magnetic 'trough' which represents a major geological feature, believed to be tectonic in origin. The trough, and all anomalies superimposed on it, trend consistently at 140° , oblique to the trends outside the zone. The eastern boundary is probably faulted along its entire lenth (f_{14}) , and as has already been shown, at least one section of the western boundary is believed to be faulted. The zone is therefore expected to represent either a double shear, graben, or horst structure. Anomalies within the zone are of relatively low amplitude averaging approximately 200 gammas. They are due to parallel bodies either outcropping or within 50 feet of the surface. In the north-west, the generally simple magnetic pattern within the zone is disturbed by two co-linear cross-trending positive anomalies. This disturbance is interpreted as being either a fault (f_5) or two basic dykes along a plane of structural weakness. The significance of the deepening of the magnetic 'trough', south-east of fault f_{13} , is not understood.

It is interesting to note that the configuration of zones A, B and C is reflected by a similar configuration of radiometric highs and low (Young and Shelley, 1966). A region of low radiometric intensity, coinciding approximately with zone C, separates two radiometric highs coinciding with zones A and B. This radiometric configuration could be partially explained by an increase of overburden in zone C, but it is believed that an abrupt lithological change is also necessary to explain the pattern.

Zone D is a region of high magnetic intensity which represents the partial resolution of two similar and parallel magnetic bodies striking 025° and dipping very steeply. The zone is correlated with an obvious topographic feature which appears to be a denuded anticline pitching to the north and south. The two magnetic trends are therefore believed to represent an amphibolitic band outcropping on the two limbs of a fold. The anomaly crossed by H-H' coincides with this band where it strikes eastwards and dips north, and the calculated width of 1200 feet represents, not the stratigraphic thickness, but the width of the anticline at its northern end measured between the two outcrops of the magnetic body. The stratigraphic thickness is approximately 400 feet maximum.

Zone E, consisting of a number of linear anomalies, is separated from zone D by a magnetic 'trough', and from zone B by a fault contact. The causative bodies within the zone outcrop and are correlated with parallel bands on the aerial photographs. By comparison of anomaly amplitudes and topographic forms between zones D and E and the northern part of zone B, the rocks appear to be similar throughout. The magnetic trough between zones D and E could therefore be due to relatively non-magnetic rock within a deep syncline complementary to the anticline of zone D.

The long and narrow zone F includes four anomalies with an average amplitude of 800 gammas; each anomaly has one axis slightly longer than the shorter axis, and is therefore due to a body with a near-gircular cross-section. These are possibly basic plugs in a zone of structural weakness.

The remainder of the area, designated zone G, is characterised by relatively low amplitude anomalies of various trends, shapes and amplitudes. The linear anomalies generally trend parallel to the geological strike, but apart from this there is desappointingly little correlation between the magnetic pattern and the strong features seen on the aerial photographs. It would seem that the very large areas of dark rock, although possibly more basic, do not have any appreciable increase in magnetic iron. Most linear anomalies are due to steeply dipping prisms of rock having a magnetite content of up to 1% greater than the surrounding rock. Their general parallelism with the geological strike implies that they are interbeds of basic material, probably amphibolitic sills and lava-flows. Some of the anomalies, e.g. that crossed by T-T' and the one immediately east of fault fig, are directly correlatable with conspicuous bands on the photographs. A few anomalies, e.g. the short anomaly immediately south of the western end of f6, trend obliquely to the known geological strike. Such anomalies are thought to be due to either basic dykes or irregularly-shaped basic intrusions.

The known copper-magnetite body north-west of Southern Cross Bore is clearly the cause of an anomaly of 1200 gammas crossed by R-R' (Plate 8). The causative body is rather less than 1000 feet long and dips near vertically. The breadth of the body, being less than the terrain clearance, cannot be accurately determined, and the value given of 200 feet probably represents a maximum. The susceptibility contrast of the body is approximately 3 x 10⁻² c.g.s. units, which is equivalent to a change in magnetite content of 5-10%.

The copper exposures north-east of Southern Cross Bore have little, if any, magnetic expression, but this is not surprising considering their geological environment. The western of the two ironstone bodies lies at the northern end of a long anomaly of 300 gammas (Plate 8), and this body could theoretically continue southwards for a considerable distance. No anomaly comparable to that crossed by R-R' was detected in zone G. Three faults have been interpreted from the co-linear termination of anomalies, but only f6 has any sign of topographic expression.

Area 5

Three magnetic zones have been delineated, one of which forms most of the area (Plate 11). Zone A, in the north-west, is a small region of high magnetic intensity which includes an anomaly of at least 2000 gammas. Zone C in the south includes a number of anomalies of intermediate amplitude trending east-south-east. Zone B, representing the remainder of Area 5, is a series of parallel linear anomalies of relatively low amplitude trending east-south-east to south-east.

Data obtained by quantitative analysis of constructed profiles (Plate 10) are shown below. All depth calculations indicate that the magnetic bodies either outcrop or lie within 50 feet of the surface.

Section	Half-width of magnetic body
	(<u>in feet</u>)
A-A *	300
B-B*	200
C=C ·	400
DD.	600
E-E	200
F_F*	400 very approximate
G-G*	400 maximum

Zone A represents only a corner of an area of greatly disturbed magnetic field. The southern boundary is well defined by the magnetic contours which could reflect a fault contact. The zone is part of an area tentatively described as being granitic gneiss with banded amphibolite. The magnetic anomalies are therefore probably due to the different magnetic susceptibilities of these two rocks, the amphibolite expected to have the greater value. Intrazonal faulting is also believed to be necessary to explain the magnetic pattern but its location cannot be deduced because of the incomplete surveying of the anomalous area. In the west of the zone two anomalies trend east and eastnorth-east. The latter, crossed by A-A', is due to a body up to 600 feet wide and approximately 1300 feet long. To the north-east the magnetic trends are not known with certainty. However, be referring to the original magnetic profiles, more information was gained than can be shown in the Plates. boundary of any survey area represents the limit beyond which the film, used for flight path recovery, could not be accurately plotted due to the turning attitude of the aircraft. Nevertheless it was possible to ascertain that larger anomalies of up to 3,000 gammas are evident immediately north of the survey boundary, and furthermore that the probable magnetic trend in this region approximates to the geological strike shown in Plate 11. A sharp deviation in magnetic trends is therefore apparent between the north and south-west of the zone.

The magnetic pattern in zone C suggests that the source rocks are similar to those in zone A, although in the southern zone they appear to have a greater length-to-width ratio. The anomaly crossed by B-B' and C-C' is due to an east-west body, or possibly two separate co-linear bodies, with a width of 400 feet in the west and up to 800 feet in the east. The combined length is approximately 3,000 feet.

Zone B occupies an area tentatively described as predominantly schist with some amphibolite. The magnetic pattern confirms the existence of two interbedded rock-types striking generally east-south-east. The apparent magnetic susceptibility contrast has been calculated to have an average of 0.9 x 10⁻³ c.g.s. units. This value does not include the anomaly crossed by secton E-E' which is due to the known copper prospect. The near-vertical causative magnetic body is approximately 400 feet wide and 2,000 feet long and its skis is midway between the positive and negative trend-lines. The anomalies crossed by F-F' and G-G' to the east, could be due to a folded extension of this body, although the lower amplitude signifies a lower percentage of magnetite.

Most of the anomalies in this zone are due to steeply dipping bodies. Although the magnetic pattern suggests the occurrence of south-east striking fold-axes, dip-angles of bodies cannot be calculated with sufficient accuracy to delineate the folds.

A north-east trending fault, f₁, has been interpreted from a displacement of magnetic trend-lines in the south-east of the area.

4. CONCLUSIONS AND RECOMMENDATIONS

The survey had two aims, both of which have been partially achieved. The first was to outline any extensions to known mineralised deposits and any anomalous areas nearby. Sixteen mineral occurrences are known in the region surveyed, and of these :-

Five have strong magnetic expression producing clearly correlatable anomalies (C-C', D-D' and E-E', F-F' of Area 3, S-S' of Area 4 and E-E' of Area 5)

Two are porbably correlated with a magnetic high (M-M', N-N' and P-P' of Area 2)

One is tentatively associated with a magnetic high (K-K' of Area 2) Two occur in, or close to, an interpreted fault (f2 of Area 1) The remaining six exposures, all in Area 4, have no obvious magnetic effect.

Six of the occurrences on magnetic evidence could have lateral extensions. These are

Both copper exposures in Area 1 (along the interpreted fault plane)
The southern copper exposure in the north-east of Area 2 (southwards)
The central apatite outcrop in Area 3 (north and north-westwards)
The copper-magnetite deposit in Area 4 (northwards)
The copper occurrence in Area 5 (general vicinity but principally north-west, plus the bodies to the south-east).

It is recommended that detailed ground investigation be made around the above six exposures. Drilling will probably be necessary in Areas 1 and 3.

Many other anomalies have been recorded in all areas. Unfortunately all known mineral deposits occur in bodies striking parallel to the regional trend of the country rock. It is therefore not possible to differentiate between mineralised and non-mineralised magnetic bodies. Ground investigation in anomalous areas is necessary to distinguish between the two.

The secondary aim of the survey was to assist geological mapping especially where the solid geology is obscured by Quaternary alluvium. The aeromagnetic results should be of considerable use in the subsequent mapping of the survey areas. Thirty-one faults have been interpreted, many of which were not already known. Although some may be found to be misinterpretations, most are given with some confidence. The trend of the magnetic disturbance is given in some detail, and as most anomalies are believed to be due to metabasics concordant with the country rock, this information should be of assistance in the structural, as well as the lithological mapping of the areas. Three fold structures in Areas 1 and 4 have also been tentatively suggested.

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APPENDIX 1. OPERATIONAL DETAILS

Survey Specifications

Ground clearance Nominally 280 feet

Detector altitude Nominally 250 feet above ground level

One-tenth of a mile Line spacing

Line orientation East-west (Areas 1, 2 and 4) Northwest-

southeast (Areas 3 and 5)

Area surveyed 115 square miles

Equipment

Aircraft Cessna 180

MNS-1 Proton precession type of B.M.R. Magnetometer

design

deflection

Recorders 1 single channel Mosely-Autograf, sensitivity 1,000 gammas full scale

1 two-channel Devar, sensitivity 10,000

gammas full scale deflection.

Modified Vinten, frame-type, 35mm., with 186° fish-eye lens.

AN/APN-1 Radio-altimeter

Storm Warning Detector MFD-3 fluxgate type of B.M.R. design, sensitivity 100 gammas full scale

deflection.

Method

Camera

A correction for diurnal variation was determined by flying a pre-selected baseline at the beginning and end of each flight. The baseline was chosen for its ease of precise reflying and relatively flat magnetic field. Each baseline profile was compared with a reference or standard profile, and the diurnal correction was applied by assuming that the variation was linear throughout the flight. This assumption was checked daily by reference to a ground magnetometer which also served as a magnetic-storm warning device.

The output of the magnetometer is a step function, and the necessary smoothing of the recorded profile also smoothed through the electronic 'noise' envelope.

Flight path recovery was made by plotting terrain features from the 35mm. film onto aerial photographs and then on to a photo-scale planimetric base-map of each area. Control points were plotted every 0.7 miles approximately. The aircraft's ground speed was assumed constant between any two adjacent control points. The maximum probable positional error is estimated to be - 150 feet.

Personnel

The personnel engaged on the survey were - D. B. Tipper (Geophysicist), A. Crowder (Senior Drafting Officer), I. Heath (Senior Radio Technician), and J. Boyd and A. E. Busuttil (Geophysical Assistants). The aircraft was flown by First Officer J. Lord of Trans-Australia Airlines.

APPENDIX 2 INTERPRETATION METHODS

The interpretation methods employed may be conveniently grouped into qualitative and quantitative analyses.

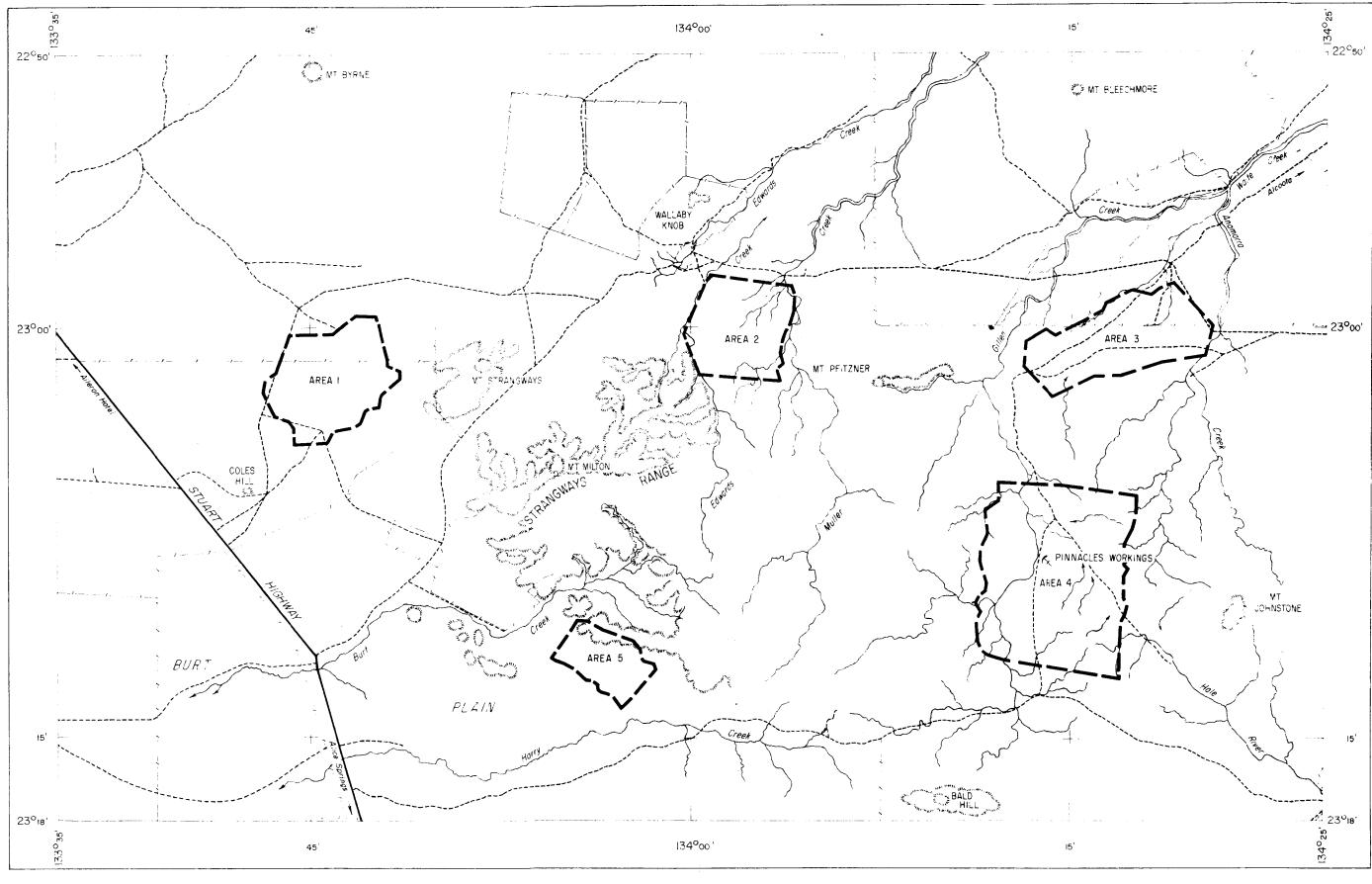
Qualitative interpretation involved the delineation of positive and negative magnetic trend-lines, being the axes of magnetic 'ridges' and 'troughs' respectively. This procedure, particularly in the case of positive trend-lines, indicates the strikes of magnetic bodies. From the displacement of anomaly trend-lines and the co-linear termination of anomalies, faulting was detected. All areas were divided into zones based on the degree of magnetic disturbance, and the significance of the zones is given by reference to known geology. Correlation was also sought between the magnetic pattern and terrain features. Where anomalies were not of ideal shape for quantitative analysis the approximate attitudes of the causative slab-like bodies of various strikes were obtained by constant reference to series of standard curves given by Gay (1963) and Reford (1964).

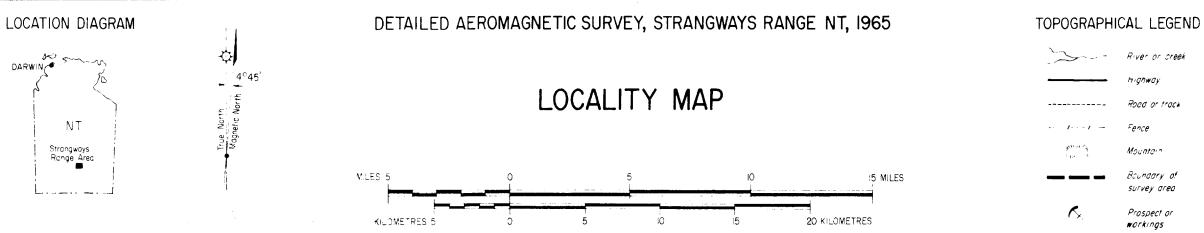
The quantitative interpretation involved the determination of depths, widths, dips and susceptibility contrasts of the magnetic bodies and was based on the assumption that the magnetization is wholly induced. Depths to magnetic bodies (below detector level) were obtained by more than one method. With anomalies of simple form, i.e. showing no partial resolution, depths were rapidly calculated by the half-maximum-slope technique advocated by Peters (1949) and extended by Moo (1965). Moo has given three independent methods for obtaining the depth of prismatic bodies. The anomaly crossed by F-F' in Area 3 was sufficiently well-defined to be analysed by all three of these methods. The values obtained, 230 ft, 250 ft and 270 ft below detector, give an average of 250 ft ± 8%. Most of the other depths quoted are considered to have an accuracy of ± 15%.

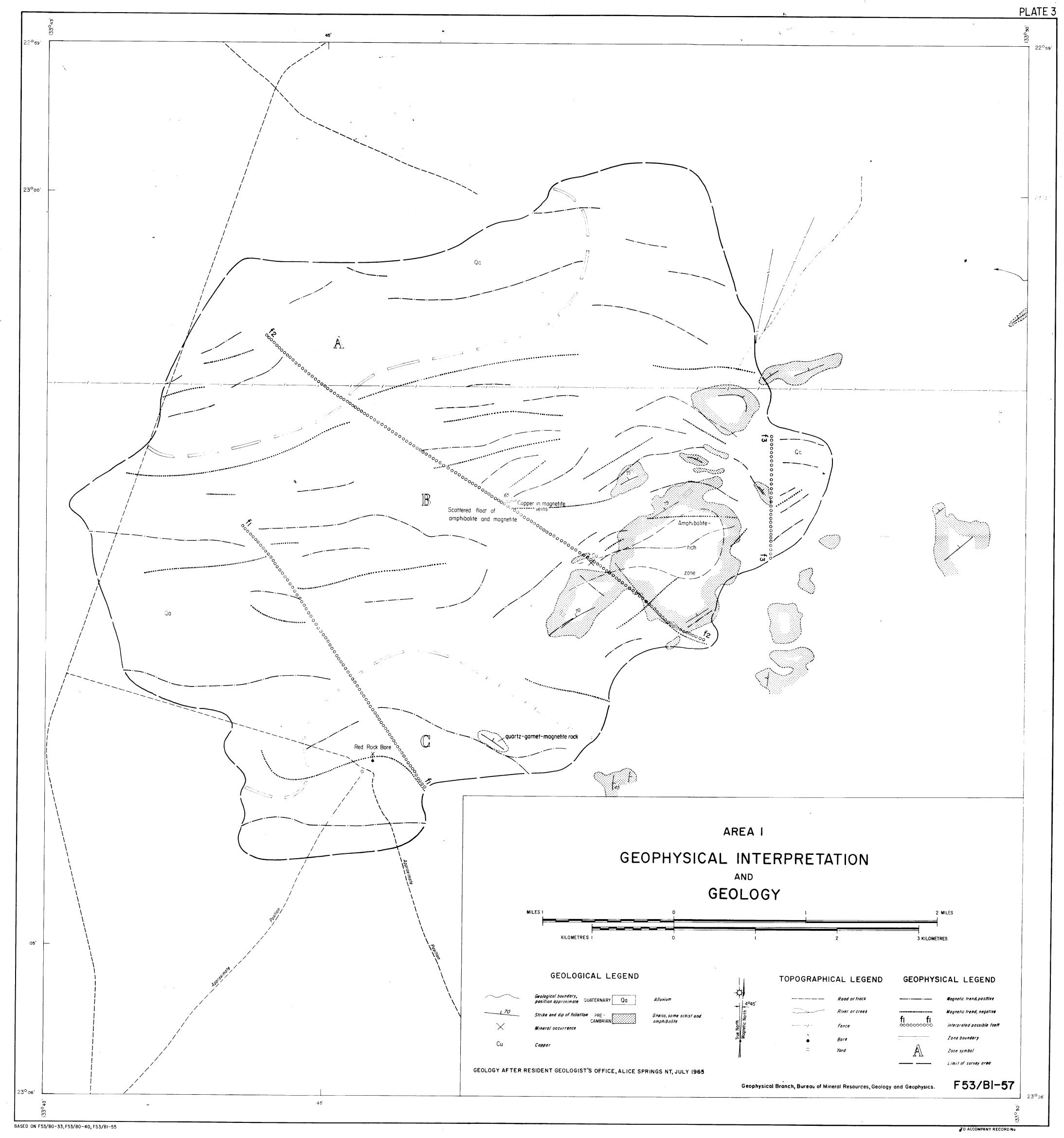
to be analysed by measurements on discrete sections of the profiles. A method of matching the whole profile with a series of standard curves was considered more reliable and those constructed by Gay (1963) were used. The depth-of burial of magnetic bodies was in all cases obtained by subtracting the ground clearance of the detector from the calculated total detector-to-body distance. Half-widths were obtained by measuring the horizontal distances between the peak of each anomaly and the position of steepest; slope. This method is reasonably accurate where the half-width is greater than the depth of the body below the detector. Where the actual total width of an outcropping body is less than 500 ft, the calculated width represents a maximum. Where possible, more accurate determinations of half-widths were made by reference to Moon(1965).

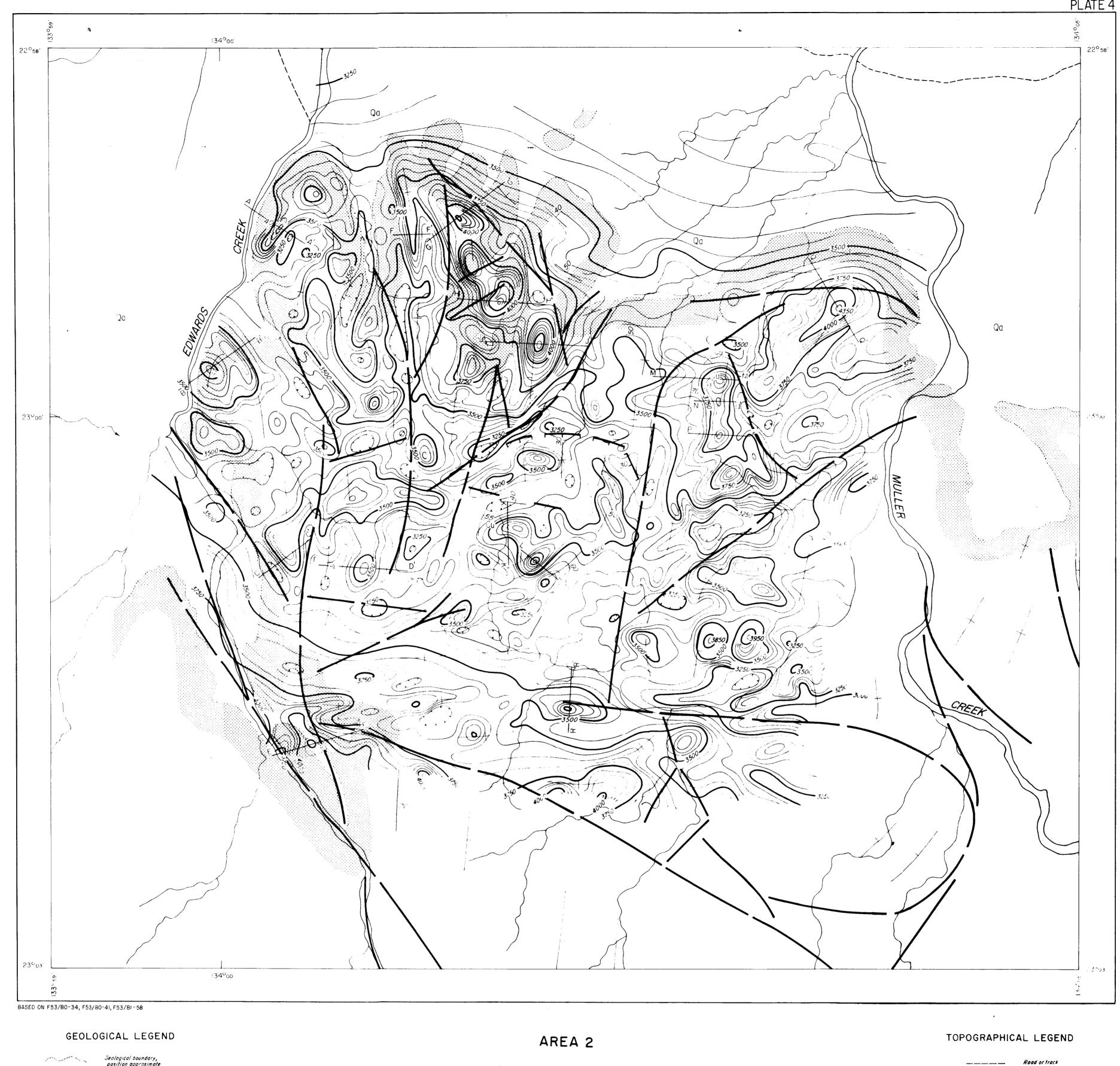
Susceptibility contrasts were calculated using standard formulae given by Reford (1964) and dip-angles were obtained using the techniques of Gay (1963) and Moo (1965).

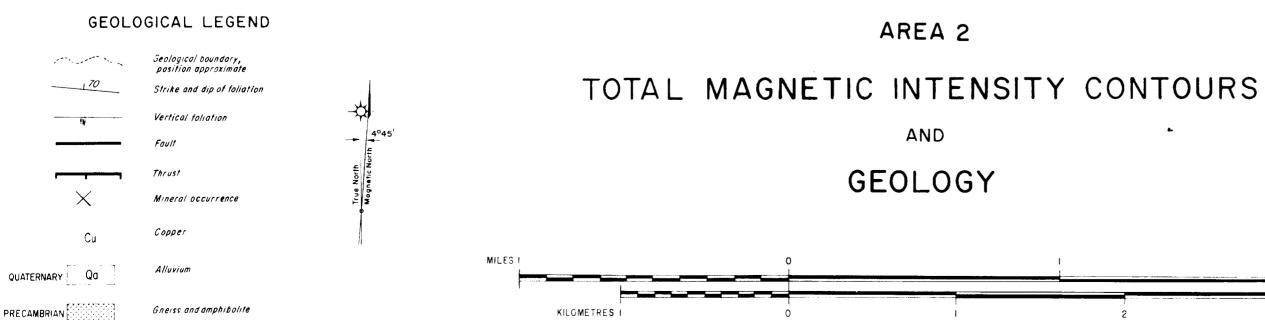






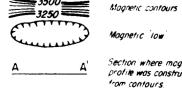






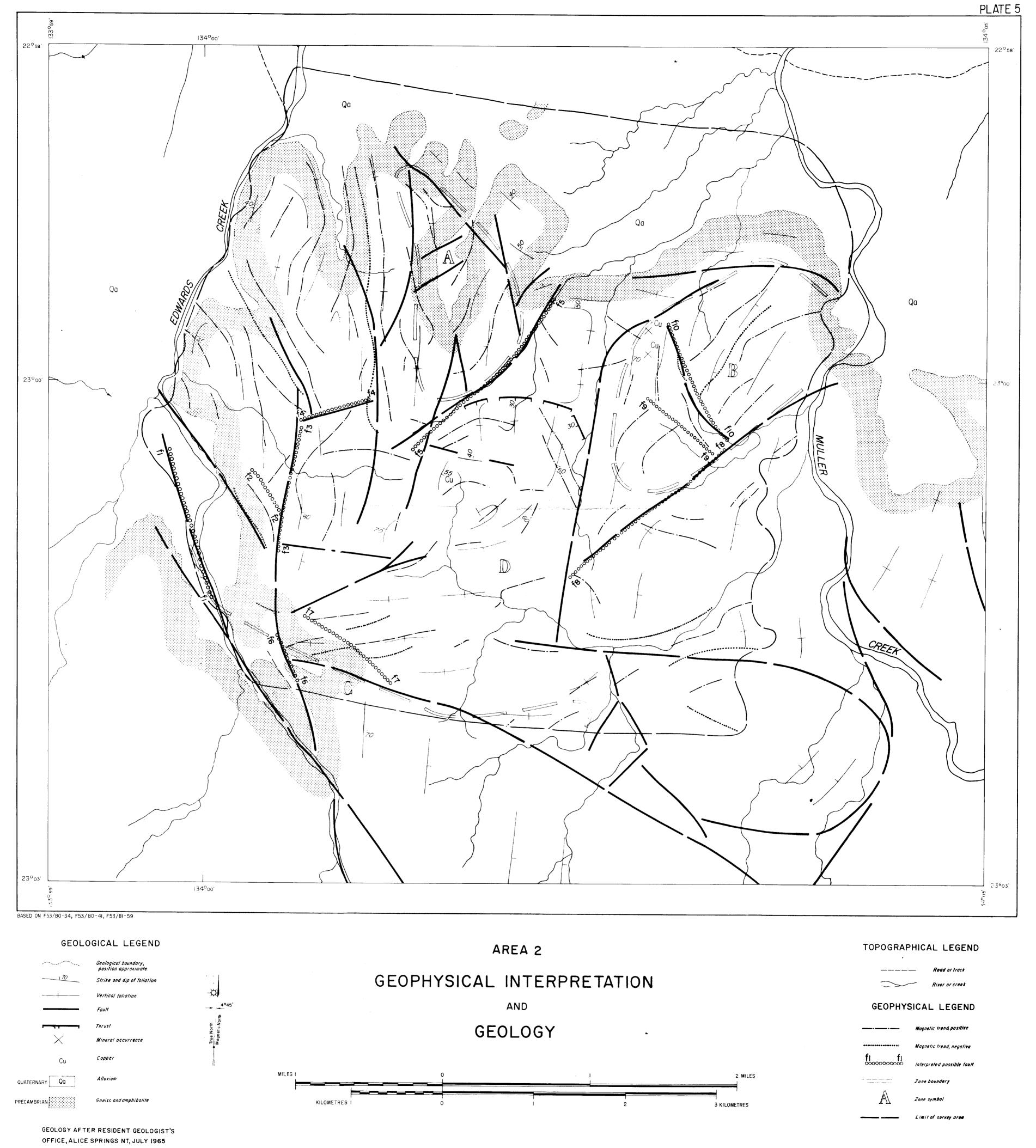
2 MILES KILOMETRES I 3 KILOMETRES

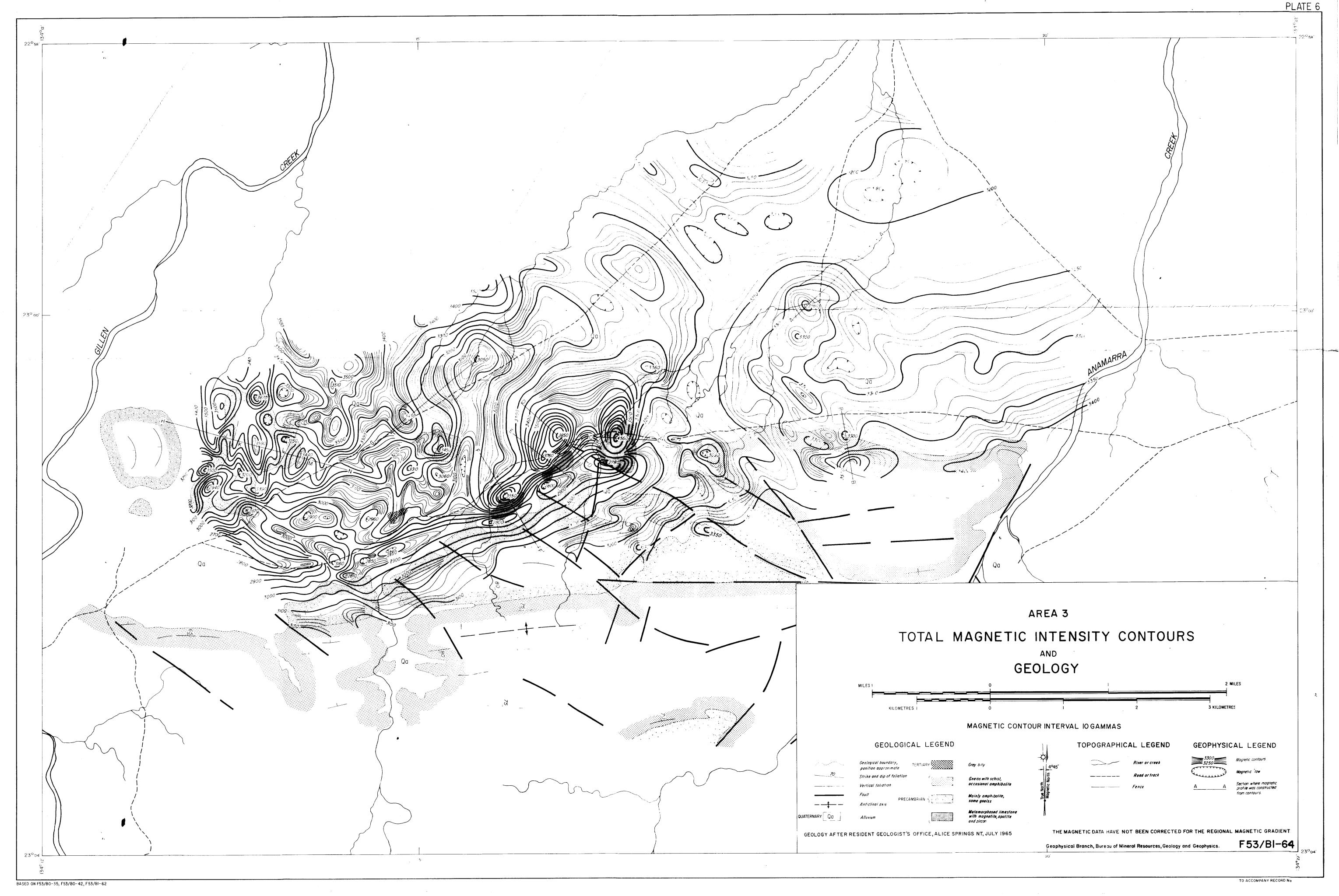
GEOPHYSICAL LEGEND

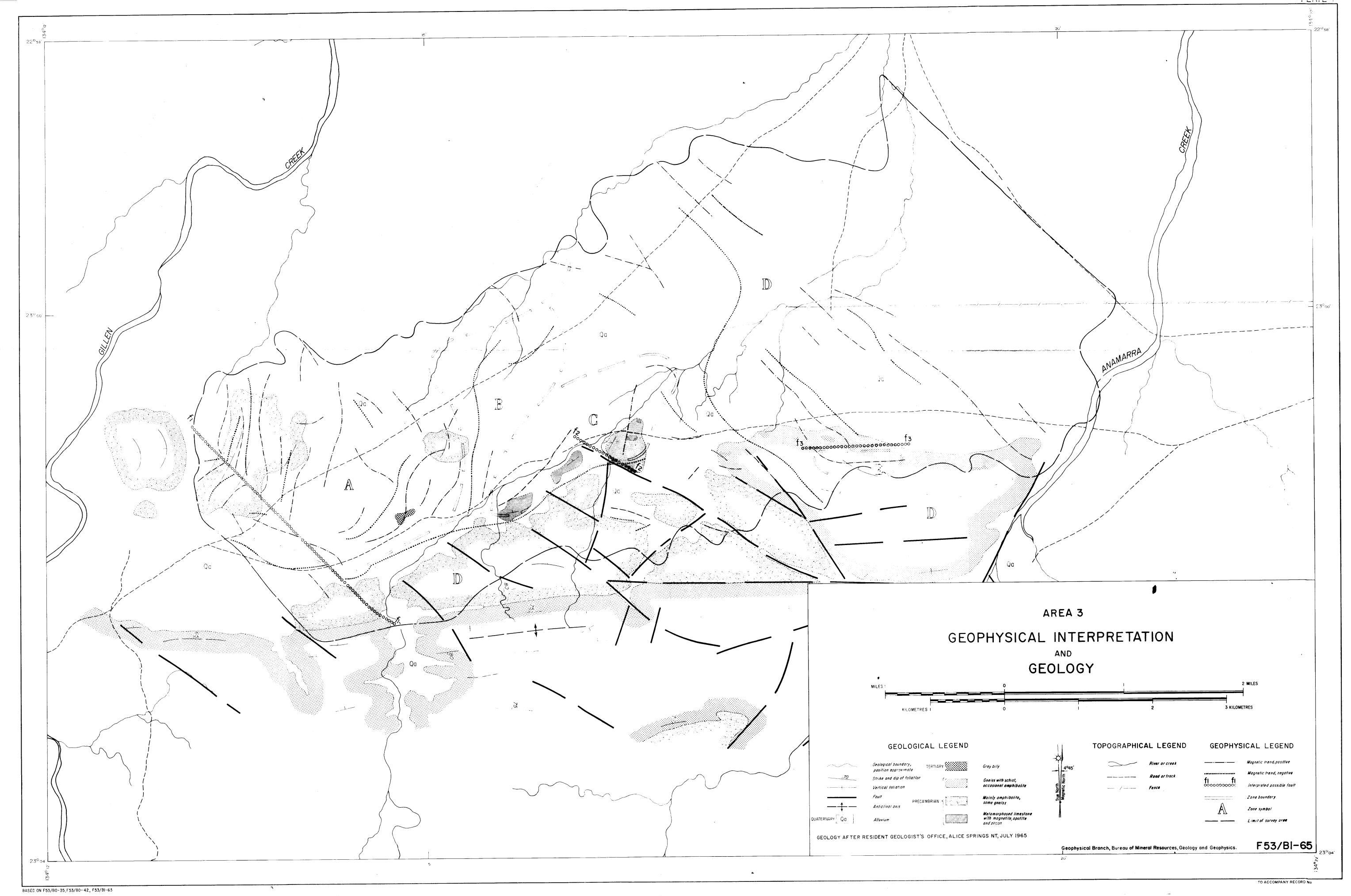


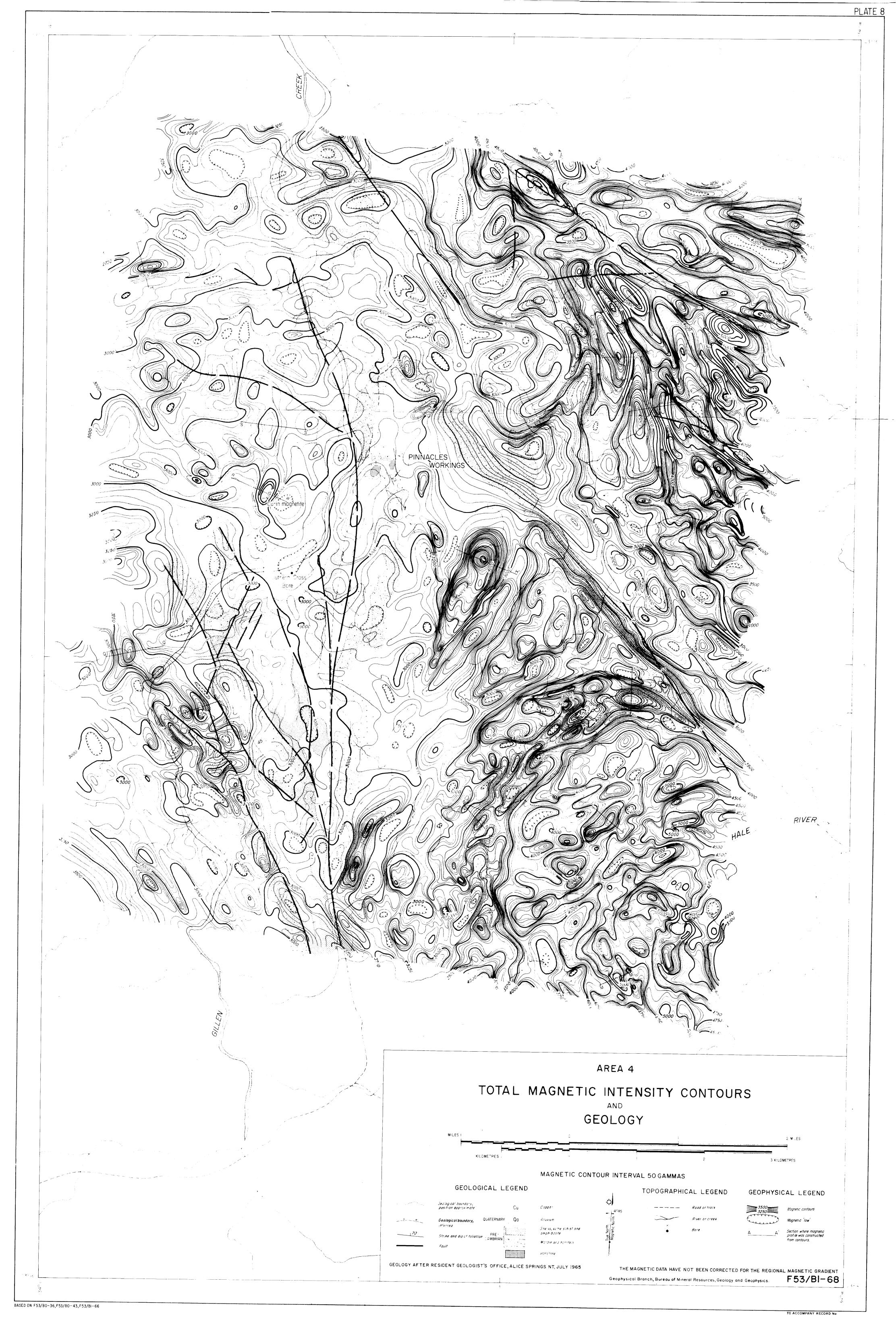
Section where magnetic profile was constructed from contours.

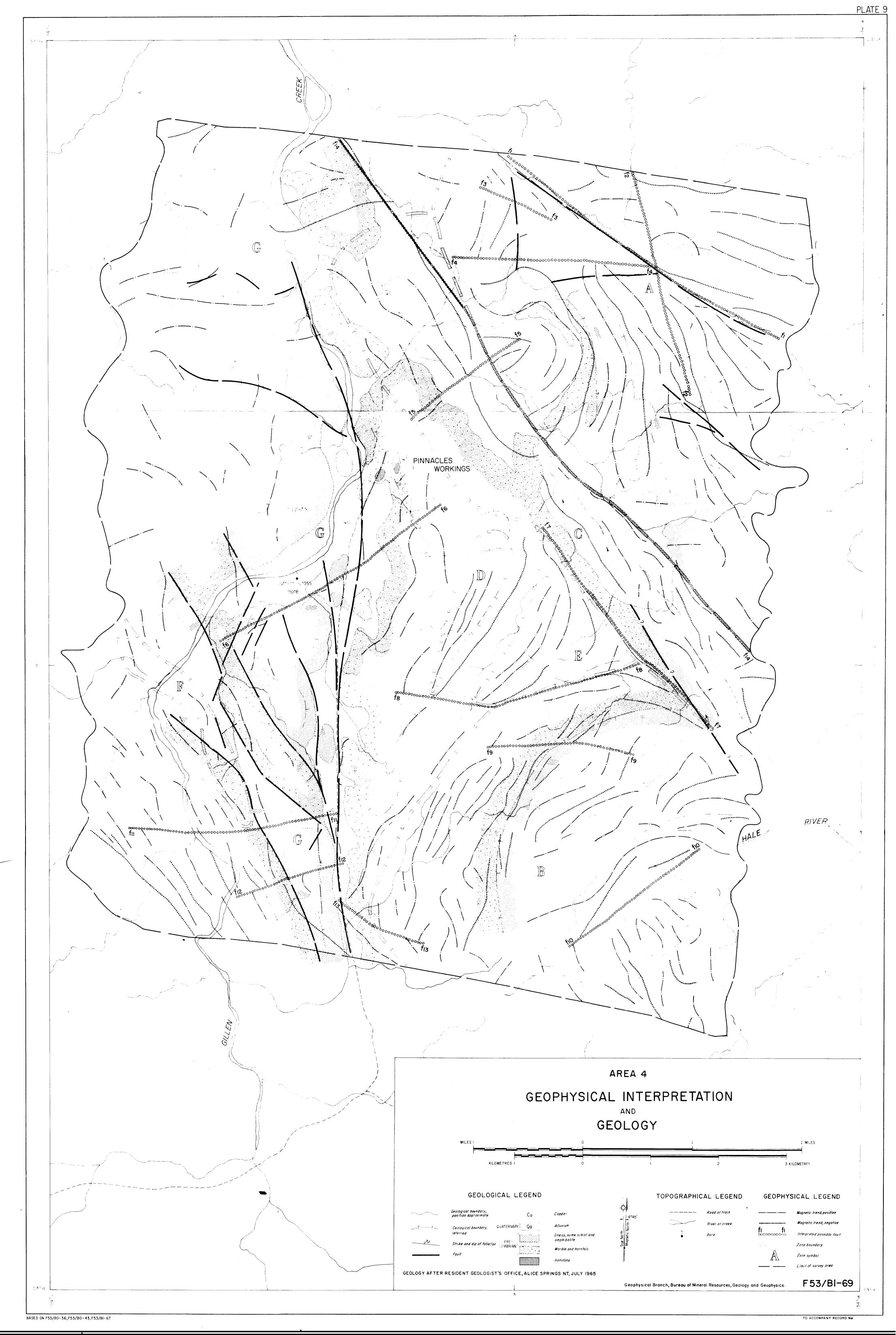
THE MAGNETIC DATA HAVE NOT BEEN CORRECTED FOR THE REGIONAL MAGNETIC GRADIENT

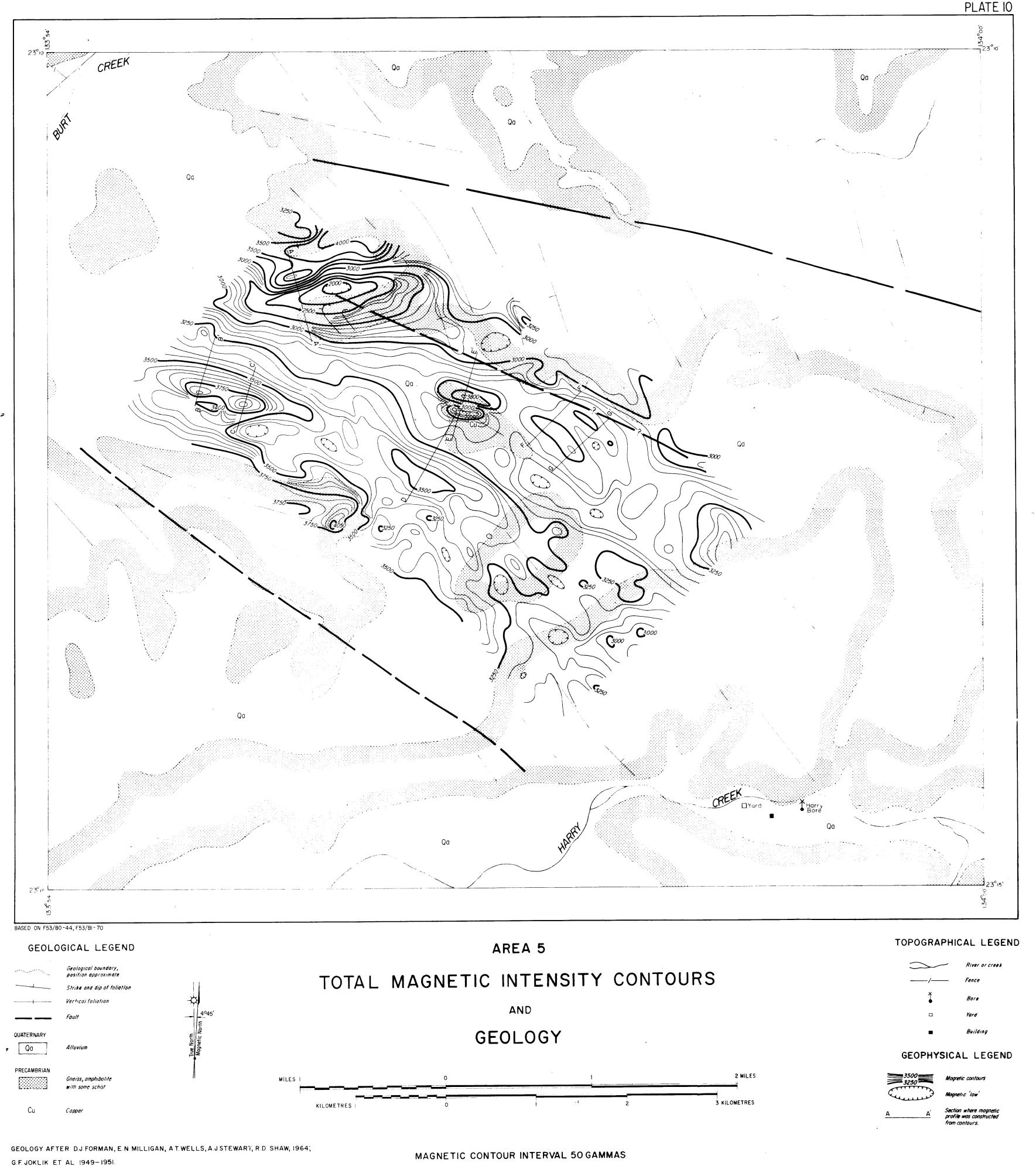












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THE MAGNETIC DATA HAVE NOT BEEN CORRECTED FOR THE REGIONAL MAGNETIC GRADIENT.

