



Collaborative MIM/AGSO Research Project:

BASE METAL MINERALISATION AT McARTHUR RIVER

**STRUCTURE AND KINEMATICS
OF THE HYC-COOLEY ZONE
AT McARTHUR RIVER**

1995/5

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SUMMARY AND CONCLUSIONS

This report summarises the results of a structural study on diamond drillcore (some oriented) and some surface and underground exposure at the HYC Pb-Zn(-Cu) mine at McArthur River, Northern Territory. The following main observations are relevant to the tectonic history of concurrent sedimentation, deformation and hydrothermal metal introduction at HYC-Cooley.

1. The Barney Creek Formation pyritic and mineralised shales at HYC contain distinctive drillcore-scale structures reflecting layer parallel compression and extension associated with predominantly "Top-to-NE" (and secondarily "Top-to-SW" layer-parallel movements). These structures, at the drillcore scale of observation, postdate diagenetic pyrite and concretions and finely laminated Pb-Zn ore and predate the Cooley inversion folding, thrusting and reverse faulting. They are interpreted to be syn-diagenetic; forming in a compacting and probably seismically active sediment pile composed of relatively brittle crusts, concretions, laminated pyrite and Pb-Zn ore interlayered with semi-consolidated (actively compacting), dolomitic-bituminous silts and muds which fail in a ductile fashion within a rapidly and inhomogeneously subsiding (NE tilting) sedimentary wedge. The kinematically "Top-to-SW" structures are probably produced during later SW-directed thrusting during the second kinematic phase of the "Cooley Inversion". The proposed NE tilting of the HYC portion of the Barney Creek shales is consistent with the known distribution of thickened sedimentary breccia accumulation (reflecting most rapid subsidence) in the NE.
2. The "Cooley Zone" at HYC is (half of) a transpressively inverted domain (positive flower structure) centred on the reactivated Emu Fault zone which contains the stratigraphically lowest units now in structurally highest positions. This inversion is believed to have culminated in upper HYC sedimentation time and probably produced a local source for the Upper Breccias. Two distinct kinematic phases characterise this inversion at Cooley: "Top-to-NW" followed by "Top-to-SW"; and are thought to be the local tectonic expressions at HYC of very broad scale regional transitions from an extended phase of extension controlling upper Tawallah Group to Barney Creek Formation deposition (Vic Wall, 1993) into the regionally extensive NS shortening event (Isan D1; ca 1600Ma). Onlap relationships north of HYC possibly involve condensed and atypical Upper Barney Creek Formation - Reward Dolomite sedimentation with high angle unconformity over reverse faulted, steeply dipping, brecciated and "Cooley mineralised" Mara Dolomite. If so, this would mark the end of the local "Cooley Zone" transpressive inversion.
3. The "Cooley Breccias" hosting copper (and lead) mineralisation are tectonic breccias lying along reverse faults within the overthrust block of Teena, Mitchell Yard and Mara Dolomites. The degree of fragmentation, fragment rotation and both dilation and stylolitic solution increases towards the reverse faults. It is unclear at present whether these breccias formed during both kinematic phases of transpressive inversion or only during the latter "Top-to-SW" phase (absence of oriented core in "Cooley Zone"). Pyrite and chalcopyrite are late in the "Cooley Breccia" infill paragenesis and within the footwall of the Basal Cooley Reverse Fault (at least) copper mineralisation is clearly associated with the latter "Top-to-SW" kinematic phase. Further detailed work on oriented core in the hangingwall breccias is planned.

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LEGEND (for all of the above)

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ACKNOWLEDGMENTS

Many past and present MIMEX staff geologists produced the excellent and voluminous geological data base at HYC in the context of which this study has been undertaken. In particular, Matt Williams and Dugi Wilson had initiated work on the internal structure of the Cooley Zone after the observation of Vic Wall and Bill Perkins that a major reverse structure would explain the gross lithological relationships in the HYC-Cooley zone. Ross Logan has been very generous with his intimate and detailed knowledge of HYC. I thank Vic, Matt, Ross and Dugi for frank and open discussions of all aspects of McArthur and HYC geology. I also salute Joseph Janacek whose insightful work in the mid-1970's was clearly well ahead of his time.

AIMS of the MIM-AGSO COLLABORATIVE PROJECT

To use available HYC-Cooley drillcore, underground exposure in the exploration decline and outcrop mapping, where available:

- *to establish the gross geometry of the "Cooley Zone" in relation to the HYC Barney Creek stratigraphy with special emphasis on the distinction of the sedimentary and "Cooley" breccias.*
- *to resolve the structural-kinematic details of the "HYC-Cooley" zone and within this framework establish the relative timing of the stratiform Pb-Zn and the "Cooley-breccia" copper mineralisation.*
- *to establish a paragenetic sequence for the "Cooley-breccia" copper mineralisation to permit subsequent well-constrained fluid inclusion and geochemical work.*
- *to examine the microstructures common throughout the Barney Creek Formation in order to further constrain the timing of Pb-Zn mineralisation relative to sedimentation, diagenesis and the structural evolution of the hosting basin.*

INTRODUCTION

HYC and the McArthur River region are located in northern Australia within the Northern Territory on the eastern edge of the "Batten Trough" adjacent to the Emu Fault Zone (Fig. 1). The McArthur Region Geology is presented on Figure 1 and shows that the HYC deposit lies to the east of the NE-plunging Barney Hill anticline and immediately west of the Western Fault Block which is bounded to the east by the Emu Fault Zone containing horsts of Masterton Sandstone. The "Cooley Breccias" lie along the western edge of the Western Fault Block juxtaposed with the HYC sequence with unknown relationships that were the focus of this study.

The HYC-Cooley drillholes logged in this study are listed in Table 1 and located on Figure 2. Table 1 also indicates the holes that were quantitatively structurally logged, the orientation method applied and the detailed Plans and Sections that accompany this report.

DRILLCORE STRUCTURE

Principles and Assumptions

Structural information can be extracted from drillcore from surveyed drillholes provided the recovered core can be rotationally reoriented (ie. correctly rotationally positioned around the drillcore's axis). Two groups of methods exist to achieve this. The first involves the periodical physical marking of core before it is broken from the bottom of the hole and then the "fitting together" of all the drillcore between oriented pieces to provide a continuous reference line. Structural features of interest are then measured relative to this reference. The second method exploits the known orientation of one structural element to orient core from surveyed drillholes allowing other elements of interest to be measured.

Both orientation methods allow drillcore to be physically reoriented in space (using a sand bucket or core orientometer jig) which subsequently allows the direct compass measurement of the other elements of interest. However, such physical reorientations and measurements are laborious and time consuming, preventing the generation of large data sets. An alternative method involves the measurement of the relative orientation of structures with respect to either the core orientation mark or the known structural element without physically reorienting the core. The spatial orientations of structures can then be calculated within a spreadsheet (Hinman, 1993). The present study uses a spreadsheet, which is a modification of the one published by myself.

This study uses both principle methods of drillcore orientation. Within the main area of HYC Pb-Zn mineralisation most of the existing drillholes are subvertical and the drillcore unoriented. However, the local orientation of bedding is well constrained from detailed correlation between drillholes within the grid-drilled HYC evaluation area. Therefore, in individual holes, provided the drillhole has not been drilled normal to bedding, drillcore can be reoriented using the interpolated local bedding orientation. However, due to the natural variability in the dip and strike of bedding and the rigidity of the bedding orientation assumption, this derived structural data will have a scatter (error) of the same order as the natural variability of the bedding orientation.

McArthur Geology

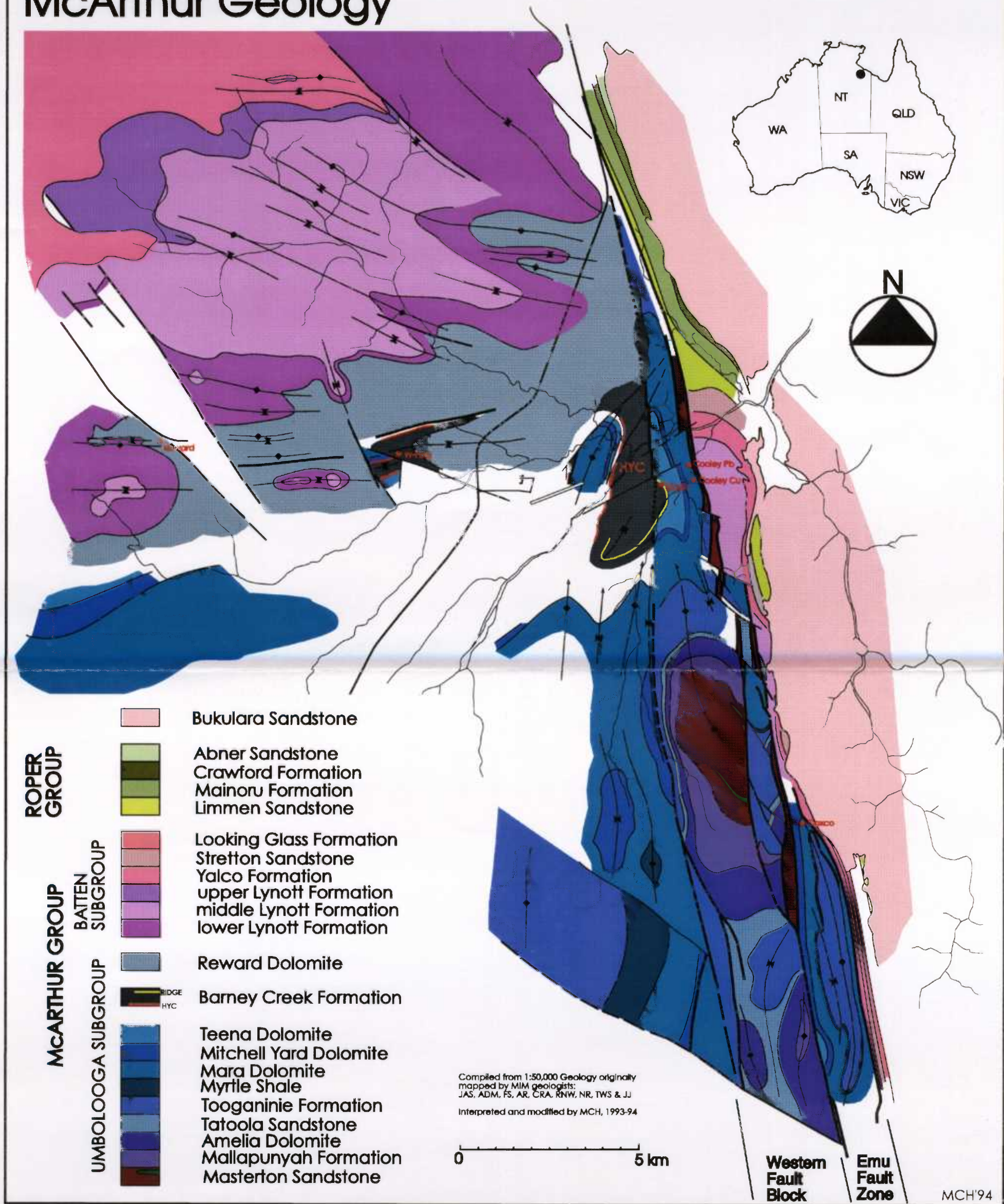


Figure 1. McArthur River Geology



* R 9 5 0 0 5 0 2 *

Drillhole and Section Locations

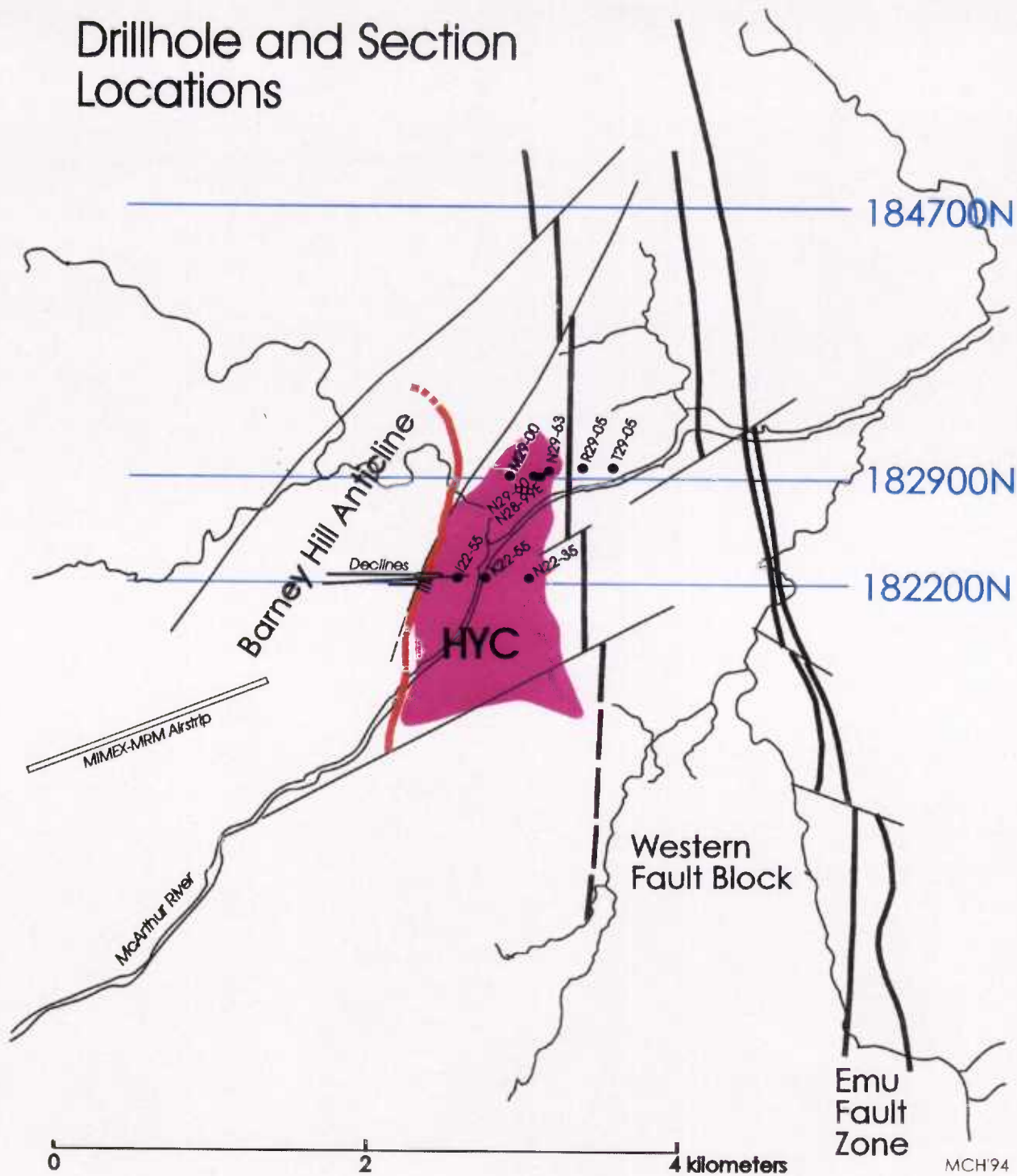


Figure 2. Location of Section and Drillholes at HYC



* R 9 5 0 0 5 0 3 *

Along the eastern edge of the HYC Pb-Zn mineralisation and within the "Cooley Zone", where the gross geometries were uncertain and detailed structure unknown, fully oriented drillcore was required to resolve the geology, structure and kinematics. Once some of the aspects of the gross geometry, detailed structure and kinematics were resolved from the fully oriented drilling, measurements from long lengths of fitted-together drillcore from other vertical and unoriented drillholes could be iteratively manipulated within the spreadsheet (ie. rotated around their own drillhole axis) to produce "best fit" solutions. Fully oriented drillcore is only as good as the technique employed to orient the core and the care taken in the technique's execution. There is little point in attempting to recover oriented core to solve critical problems if the confidence in the technique and its application is low (see Limitations below).

Table 1. Drillholes logged in this Study

STRUCTURAL				
Hole	Section	Orientation Method	Data Status	Plans/Section Plotted and Provided
I22-55	182200N	Bedding	Finalised	1:500
K22-55	182200N	Bedding	Finalised	1:500
N22-35	182200N	Bedding	1°Calc only	
N29-00	182900N			
N29-60	182900N			
N28-99E	182900N	Downhole Spear	Finalised	1:250,500,1000
O29-63	182900N			
R29-05	182900N	"Best Fit"	Finalised	1:20,100,500
T29-05	182900N		Max App Dip	
R27-05	182700N		Max App Dip	
Emu 16	184700N	Downhole Spear		
Mt Stubbs 3	184700N			
Mt Stubbs 4	184700N			

1. "Best Fit" of bedding, fault orientations and kinematic indicators with adjacent fully oriented drillhole and nearby outcrop.
2. MaxAppDip=Maximum Apparent Dip treatment only. Commonly core too broken to fit together large sections for iterative "Best Fitting".

Limitations

The resolution of structures and kinematics in drillcore has limitations. The major sources of these are, firstly functions of the drillcore orientation process and secondly, the relative scale of the observation (ie. core size) in comparison with the scale of the investigated structures.

The scatter in data derived from drillcore oriented according to rigid assumptions made about bedding orientations will reflect the natural variability in bedding dip and strike (see also above). Around fault zones where bedding is commonly rotated significantly the method completely fails (eg. I22-55). The only recourse is to attempt to fit all the core together from outside the disturbed zone into the fault zone....commonly a task made difficult, impossible and fruitless by the broken nature of core around the fault.



In drillholes where bottomhole orientations have been successfully done at regular intervals, the derived structural data should be excellent. Unfortunately, those methods that rely on gravity in inclined holes to mark the bottom side of recovered core are prone to considerable "bounce" off the back of the drill bit and/or the bottom of the core barrel at the bottom of the hole resulting in very significant misorientations. This becomes obvious when long sections of core through more than one core orientation mark can be successfully

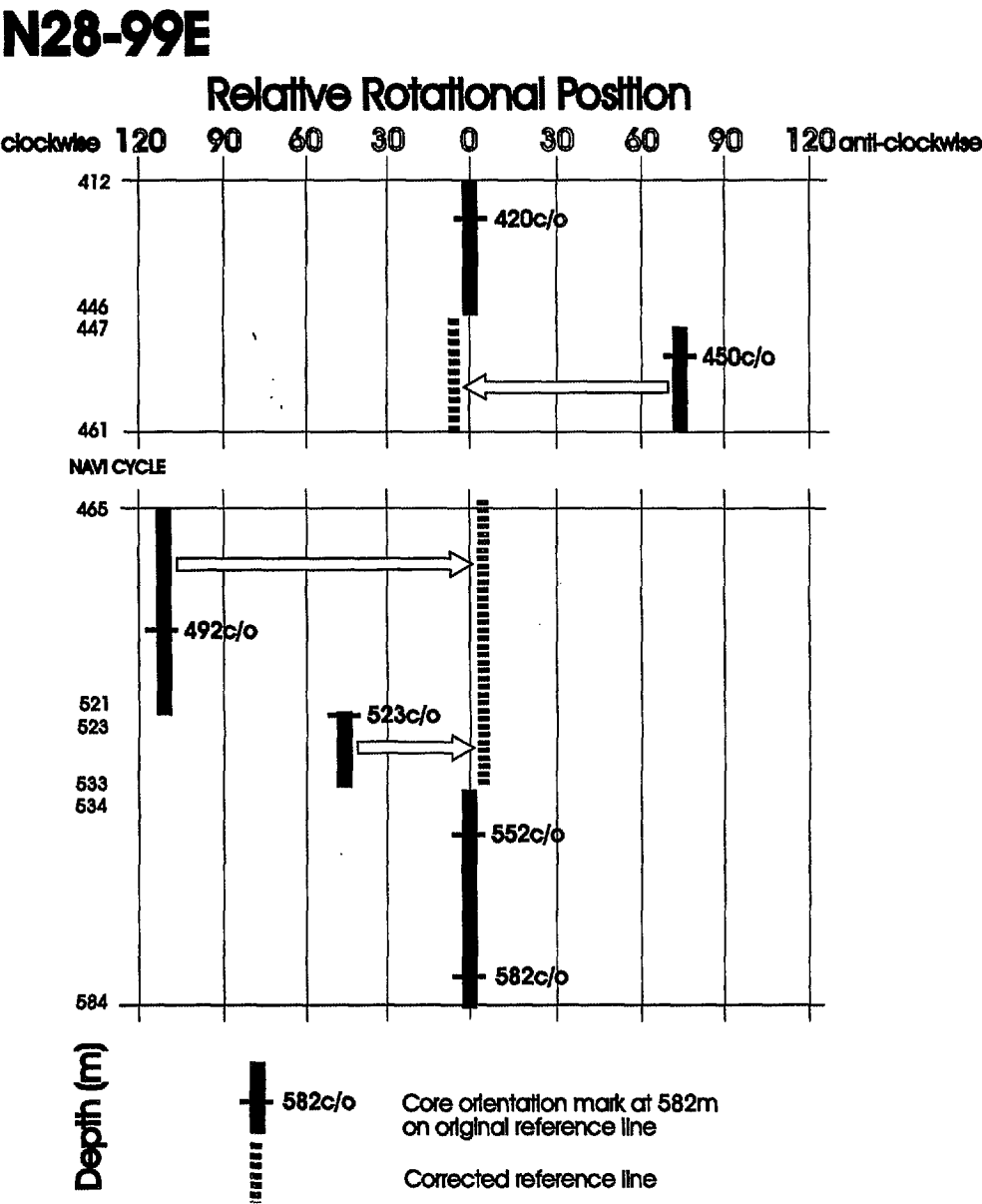


Figure 3. N28-99E Core Orientation - Relative Rational Positions

fitted together. Severe problems of this nature occurred in the critical hole, N28-99E between 412m and 584m (Fig. 3) with the result that confidence in the structure of this zone is considerably reduced. A little more care in the execution of the techniques and more regular orientation where ground is more broken can vastly improve the final structural product and improve the understanding of critical relationships and processes. Figure 3 shows how far successive core orientation marks were relatively rotated in two

sections of completely fitted-together drillcore (separated by a Navi drilling cycle) in N28-99E between 412m and 584m. The figure also shows the final (version 5!) adjustments that were made, based largely on bedding orientation correlations across the Navi drilling gap and the "agreeing" core orientation marks at 552m and 582m. The most doubtful zone remains between 412m and 461m....a section that contains an abundance of interesting thrust? microstructures in W-Fold Shale.

The second problematic aspect of the resolution of structures from drillcore is scale dependant. Structures smaller than the size of the core (eg. microfaults, low angle microthrusts....) can commonly be fully resolved with respect to orientation and direction, sense and magnitude of movement when offsets are smaller than the intersection of the structure with the core. However, at the other end of the scale spectrum, drillcore through very large structures (eg. the Basal Cooley Reverse Fault) are composed of broad zones (many to tens of meters) of mixed lithologies, spaced anastomosing fabrics and abundant secondary structures. Orientation of the main structure, in these cases, is commonly impossible to determine from oriented drillcore in an individual hole. Multiple intersections of the structure are required to accurately resolve its orientation. Oddly, however, the gross direction and sense of movement can commonly be determined, although the magnitude of displacement is generally well beyond the scale of the core. Between these end members of scale there are many structures that can be variably resolved with respect to orientation and kinematics. It should be born in mind that, without multiple intersections, it is the larger structures that are less well constrained in orientation, despite being generally well understood kinematically from associated smaller scale secondary structures.

DETAILED STRUCTURE AND GEOLOGY

The detailed geology and structure of selected HYC and HYC-Cooley drillholes are presented in a series of plans, sections and overlays listed at the start of this report and summarised in Table 1. The structural logs are presented in spreadsheet form in Appendix I. There are two distinctive groups of structures observed in drillcore with contrasting orientations and kinematics. The earliest group of structures are confined to the W-Fold, Lower Dolomitic, HYC mineralised Lower Pyritic, Upper Pyritic, Bituminous and Upper Dolomitic Shales (Barney Creek Formation Shales) and are thought to have formed within the Barney Creek sediment pile while the sediments were still relatively unlithified. The second group of structures clearly affect the entire sequence and have been logged from Mara Dolomite as stratigraphically high as the Upper Breccias. These reverse movement structures reflect a local inversion that formed the, so called, "Western Fault Block" and "Cooley Zone".

Detailed structural work has been completed on two sections: 182200N and 182900N. Work on 182200N includes mapping of Barney Hill, mapping of the Exploration Decline and detailed, quantitative structural logging of I22-55, K22-55 and N22-35. Work on 182900N includes mapping of Barney Creek, mapping of the Barney Creek Crossing outcrops and detailed, quantitative structural logging of N29-99E, R29-05 and T29-05 (see 1:10000 Geology Plan). Qualitative structural observations were made in the other logged HYC-Cooley drillholes listed in Table 1.

"EARLY" (MICRO)STRUCTURES

The earliest, drillcore-scale, structures include the extensional "Step" Microfaults, "soft sediment" intraformational Microfolds and Imbricate Stacks, the Low Angle (to bedding)

Microfaults and Crack Extensional Veins in concretions. Sketches of these structures are shown on Figure 4. They have been structurally logged in I22-55, K22-55, N22-35 and N28-99E (see accompanying plans and sections and Appendix I). The most remarkable feature of all these structures is their consistent strike and kinematic indication (bar one set of low angle structures; see below). Representative stereoplots of these structures are also shown in Figure 4 and show that, within the tolerances of the drillcore orientation method used (see Table 1), these structures strike NW-SE.

Microfaults and Extensional Veins....."Top-to-NE"

The extensional Microfaults are very common throughout the W-Fold Shale and the Pyritic and Bituminous Shales. The Crack Extensional Veins are confined to concretionary sections of the same units and are commonly continuous with moderate scale Microfaults. The Microfaults display a continuum of scales from very closely spaced faults displacing a small number of laminations and apparently confined to those set of laminations by bedding parallel slip on adjacent layers, up to larger, similarly oriented, faults cutting many centimetres to tens of centimetres of laminations including bands of the smaller scale Microfaults. They all show "SW-Block-Down" displacements and indicate layer-parallel extension. However, the bedding-confined, smallest scale Microfaults are difficult to rationalise with "SW-Block-Down" normal block faulting as they are "rootless". They are better rationalised by (and consistent with) "Top-to-NE" bedding parallel slip. The larger scale Microfaults are therefore likely to reflect similar bedding-parallel deformation but on thicker packages. The Extensional Veins in concretions are intimately associated with Microfaults and indicate NE-directed layer-parallel extension.

Imbrication and Intrafolial Folds....."Top-to-NE"

These moderately common structures occur within the Mineralised and more highly Pyritic Shales. Where they have been measured they too are consistent with "Top-to-NE" layer parallel deformation but reflect layer-parallel compression.

Low Angle Microstructures....."Top-to-NE & Top-to-SW"

The Low Angle Microstructures are extremely common within the Mineralised Shales and moderately so within highly Pyritic Shales. They usually lie at 20-30° to bedding regardless of the orientation of bedding and flatten out into bedding (Fig. 4). They clearly displace bedding laminations by fractions to a few millimetres, and they display both normal and reverse movements. Measurements in a number of holes (I22-55, K22-55, N22-35, N28-99E) suggest that NE-dipping NORMAL and SW-dipping REVERSE Low Angle Structures, indicating "Top-to-NE" layer-parallel deformation, are roughly equally common and together are considerably more common than SW-dipping NORMAL and NE-dipping REVERSE Low Angle Structures that indicate "Top-to-SW" layer-parallel deformation (see Fig. 4). Together they represent conjugate sets of layer parallel compressional and extensional structures reflecting dominantly "Top-to-NE" and secondarily "Top-to-SW" (or "Bottom-to-NE") kinematics (see extensional and compressional arrows on Fig. 4).

Relative Timing Relationships

Some of the relative timings of the above structures with respect to each other and with respect to syn-diagenetic features (eg. pyrite, concretions and ?Pb-Zn mineralization) are evident from overprinting relationships. Following are a number of timing observations, which are summarised in Figure 5:



- a) Medium to large scale Microfaults cut concretions, Intrafolial Folds, Imbricate Stacks, small scale Microfaults and rarely Low Angle Microstructures (Microfaults are rare within Mineralised Shales).
- b) Extensional Veins and associated Microfaults cut concretions.
- c) At the drillcore scale of observation, all the microstructures cut and/or deform early diagenetic pyritic laminae and Pb-Zn mineralised shale laminae.
- d) Low angle REVERSE structures cut and displace Low angle NORMAL structures as often as vice-versa.

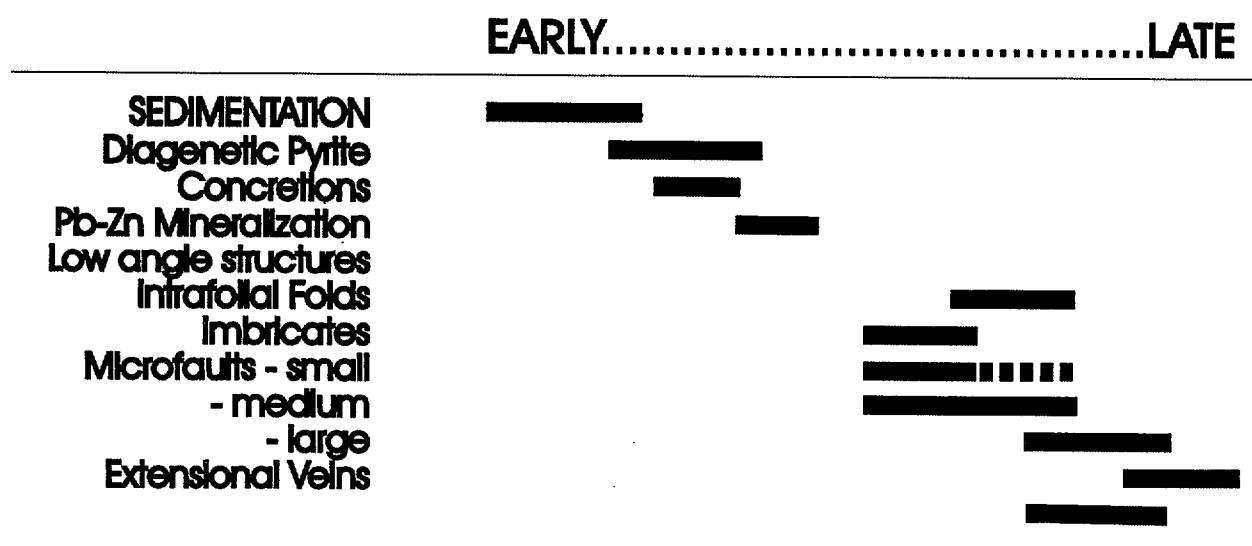


Figure 5. Relative Timing of Early Structures and Diagenetic Features within an individual horizon

"LATE" TRANSPRESSIVE STRUCTURES

A number of major fault zones whose gross orientations and kinematics are not easily resolved from drillcore (see Limitations in DRILLCORE STRUCTURE above) are present in drilling towards the eastern edge of the HYC Pb-Zn mineralised area and within the "Western Fault Block" or "Cooley Zone". Fortunately, on the major section logged, 182900N (Figs 6 & 7), the drillhole spacing is sufficiently close to constrain the gross geometry of these major structures. In addition, a host of secondary, parallel and conjugate structures associated with the major structures are fully resolvable with respect to orientation and kinematics. They can be used to further constrain the orientation of the major structures assuming a degree of parallelism between the different scale structures and fully resolve the large scale kinematics.

182900N

Barney Hill-HYC-Cooley Geology

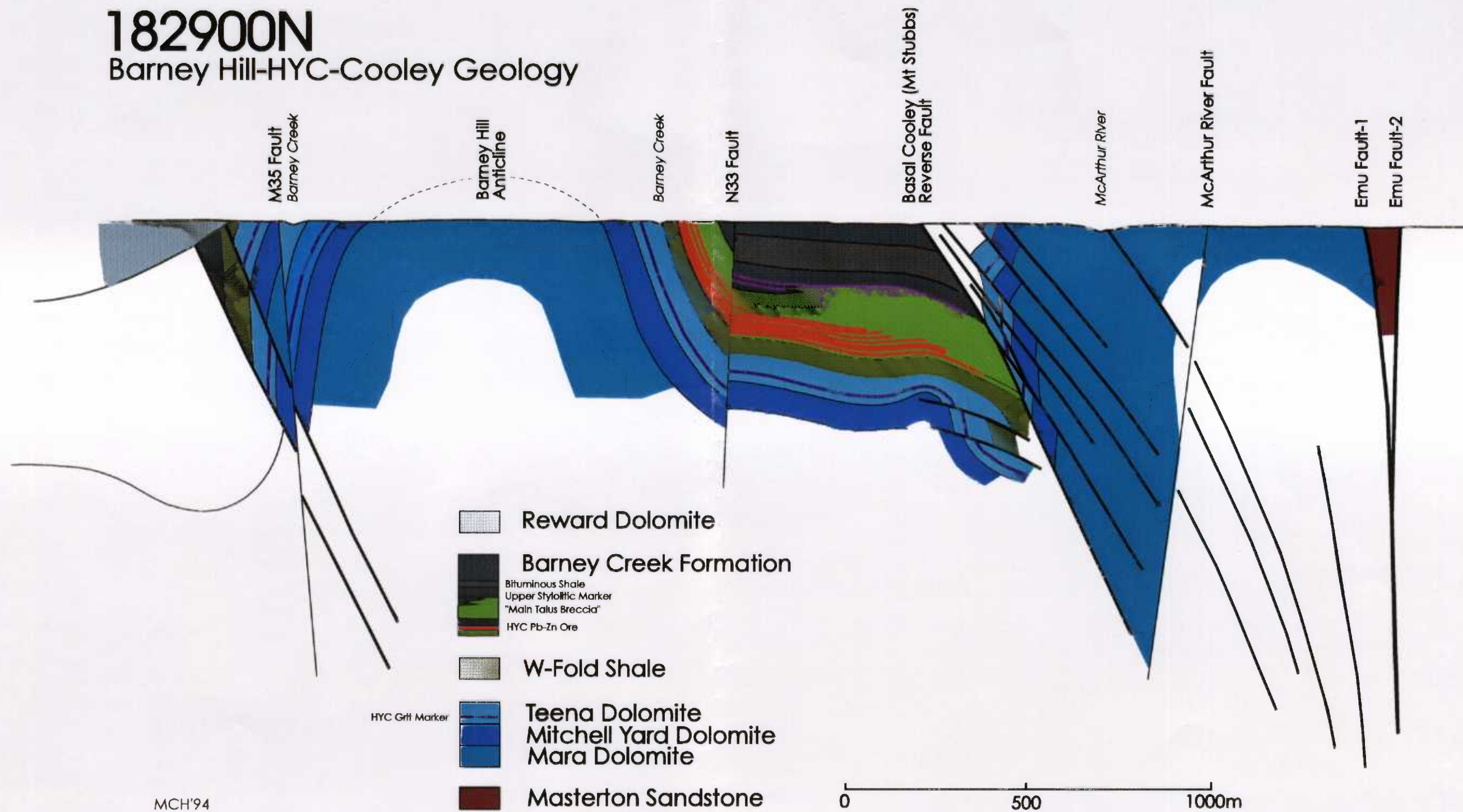


Figure 6. Section 182900N Geology



* R 9 5 0 0 5 0 7 *

Gross Geometry of Cross Section 182900N

The interpreted gross geometry across the HYC-Cooley zone at 182900N is presented in an accompanying Section (Fig. 7) and Structural Overlay (Fig. 8). The Cooley Zone is interpreted to be a transpressively reverse faulted, steeply dipping (to locally overturned), short limb block of folded Teena, Mitchell Yard and Mara Dolomite lithologies (on this section) thrust against and over much of the Barney Creek Formation.

The Basal Cooley Reverse Fault was intersected in R29-05 and is a 5 meter wide anastomosing zone of mixed rock types from above (Teena) and below (Shale \pm pyritic and sedimentary breccia-grits). The orientation of this structure is constrained by the geology of the drillholes adjacent to R29-05. The structure may be steeper than indicated on Figures 6 and 7 but must be a *reverse fault*. The structures within the hanging wall of the basal reverse fault are most commonly zones of intense tectonic brecciation containing rotated fragments of the adjacent dolomitic lithology. The zones of intense brecciation separate blocks of differently, but internally consistently oriented, coherent dolomitic stratigraphy. They have been represented on the section, for convenience, as reverse faults (F3; 182900N; Fig. 7) although only a few actually contain a discrete structural surface that can be measured (Fig. 8). Based on the bedding orientations (Maximum Apparent Dips; ie. rotationally unfixed core; see Table 1) and consistent facings (ie. uphole or downhole) within coherent blocks in R29-05 and T29-05, some attempt to schematically link the zones of intense brecciation has been made on Figures 6 and 7. Very steep to overturned (downhole facing) sections are present in both holes and have been tentatively correlated. However, it seems highly likely that these zones of intense brecciation and faulting anastomose more strongly than indicated and limited structural data, where fault surfaces have been measured, suggest that the structures within the hanging wall of the Basal Cooley Reverse Fault are not parallel to the Basal Fault and in fact strike at moderate angles to it.

In the footwall of the Basal Cooley Reverse Fault are a number of flat to gently NE dipping thrust surfaces (F2; 182900N; Fig. 7) that within the W-fold Shale, Lower Pyritic Shales (& sedimentary breccias) and Main Talus Breccias are probably bedding parallel (NE dipping) and exploit the shaly (weak) horizons within the HYC package.

HYC Ore-Sedimentary Breccia Facies Relationships

At the eastern end of Section 182900N west of the "Cooley Zone" the sedimentary breccia cycles of the HYC Mineralised Shale and Main Talus Breccia can be well correlated through a number of drillholes (N29-60, O29-63 and N28-99E; Figs 6 & 7). From west to east Orebodies 8 to 5 are progressively displaced in the sedimentary pile by breccia cycles. Orebodies 4 to 2 are significantly less mineralised in their easternmost intersections (N28-99E) where the Basal Breccia, the strongly pyritic units at 2 Orebody and above 3 Orebody, the 2/3 interore breccia cycle and the base of 4 breccia cycle are well correlated across the entire section.

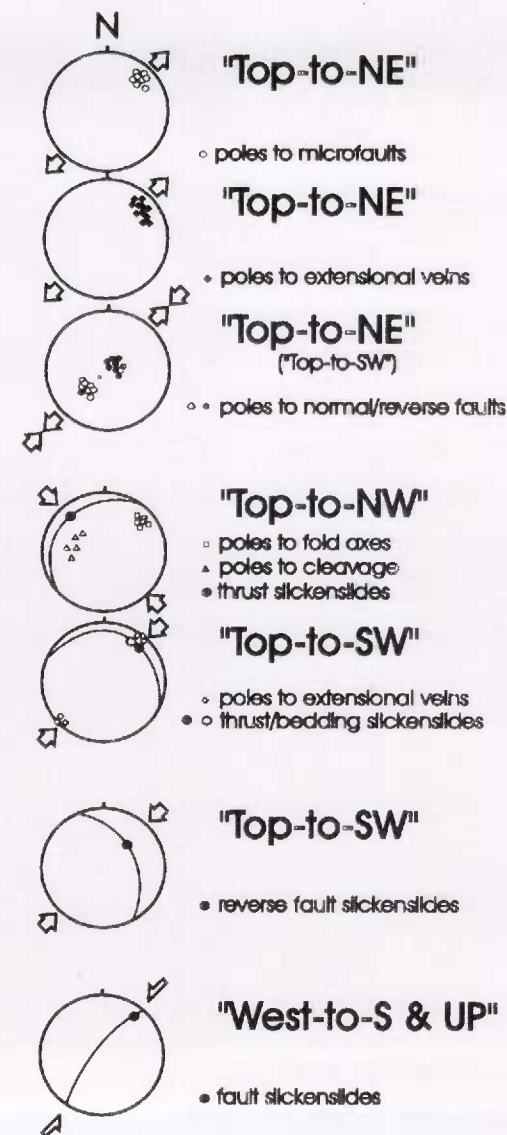
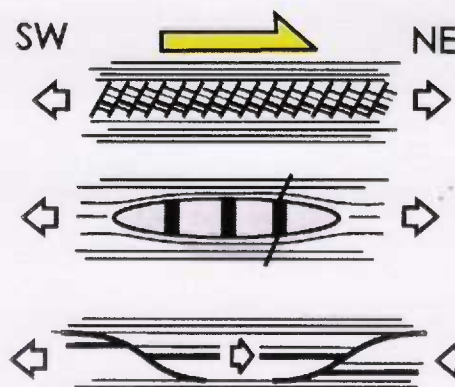
The breccias at 5 to 8 Orebody positions are (galena) mineralised. Both the sedimentary breccia matrix and the rims of dolomitic clasts commonly show partial galena replacement. This mineralisation is quite unlike the "Cooley Brecciation" style of mineralisation found further east on this section in the footwall of the Basal Cooley Reverse Fault (R29-05) where base metals are localised within fracture networks within the breccia clasts (Fig. 8). It seems probable that three mineralisation styles exist on this section. Temporally the first two involve ?syn-diagenetic replacement of sedimentary breccia clasts (and matrix) while, to the west, the main HYC Pb-Zn orebodies formed by ?replacement of fine grained sediments which were deposited in onlapping positions with respect to non interore





Structural Summary OVERLAY Key

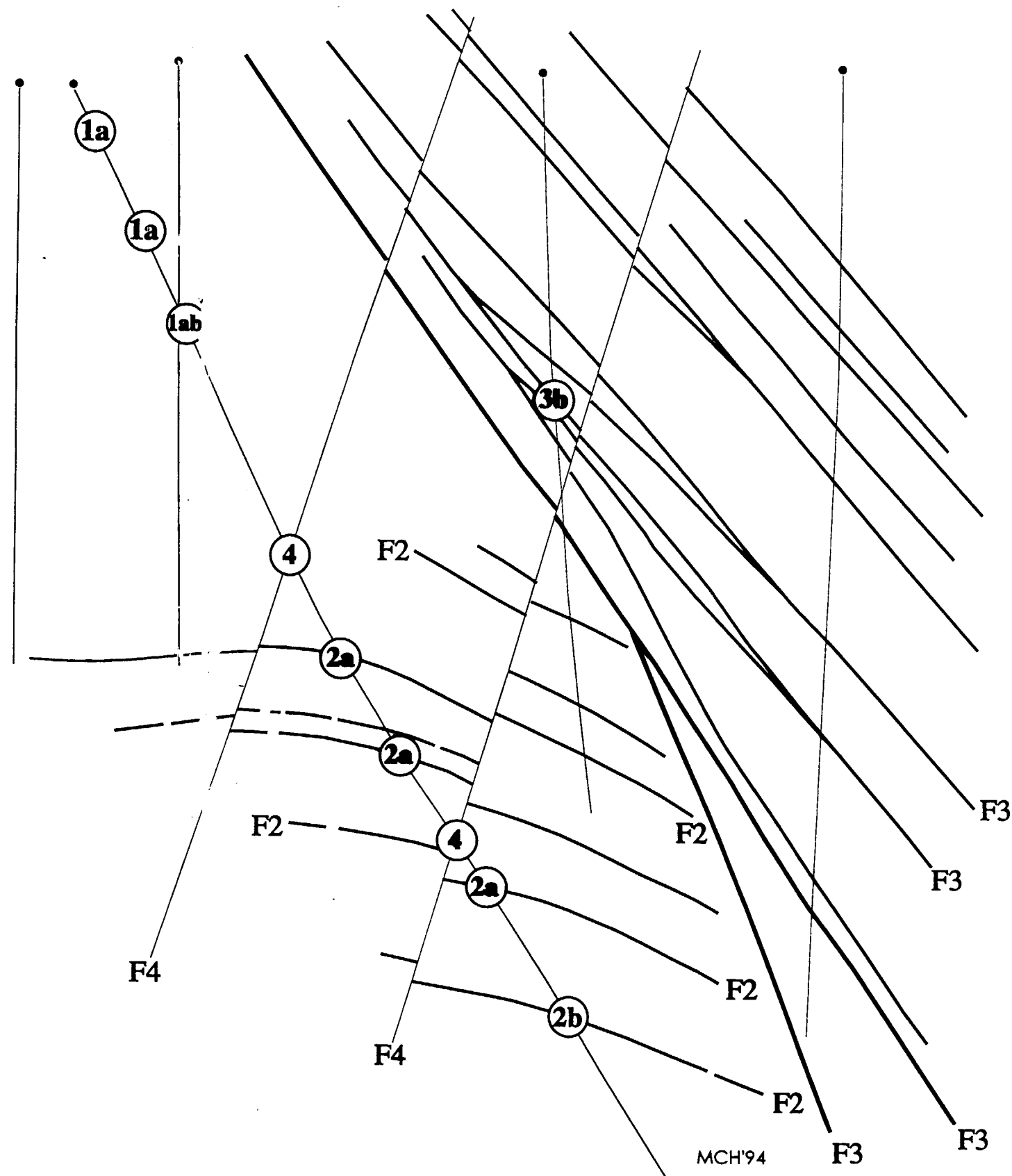
- 1a** Microfaults
- 1b** Extensional Veins
- 1c** Low angle Microstructures
- 2a** Shallow Thrusts
- 2b** Shallow Thrusts
- 3b** Reverse Faults
- 4** Late Steep Faults



182900N

Structural Summary
OVERLAY

(see accompanying
OVERLAY Key)



* R 9 5 0 0 5 1 1 *

Figure 8. Structure Summary Overlay

182900N Geology

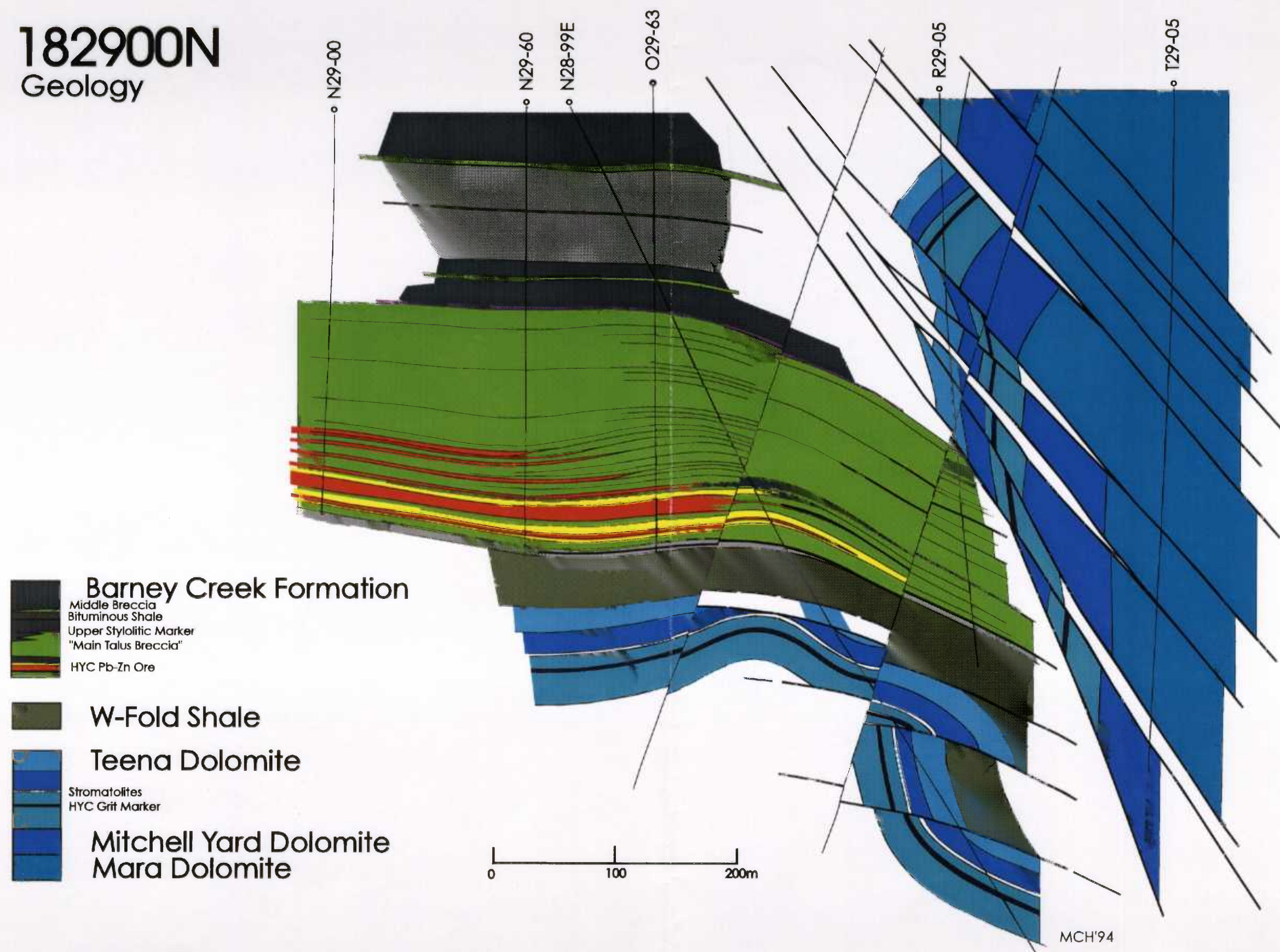


Figure 7. 182900N HYC-Cooley Geology



182900N Mineralization

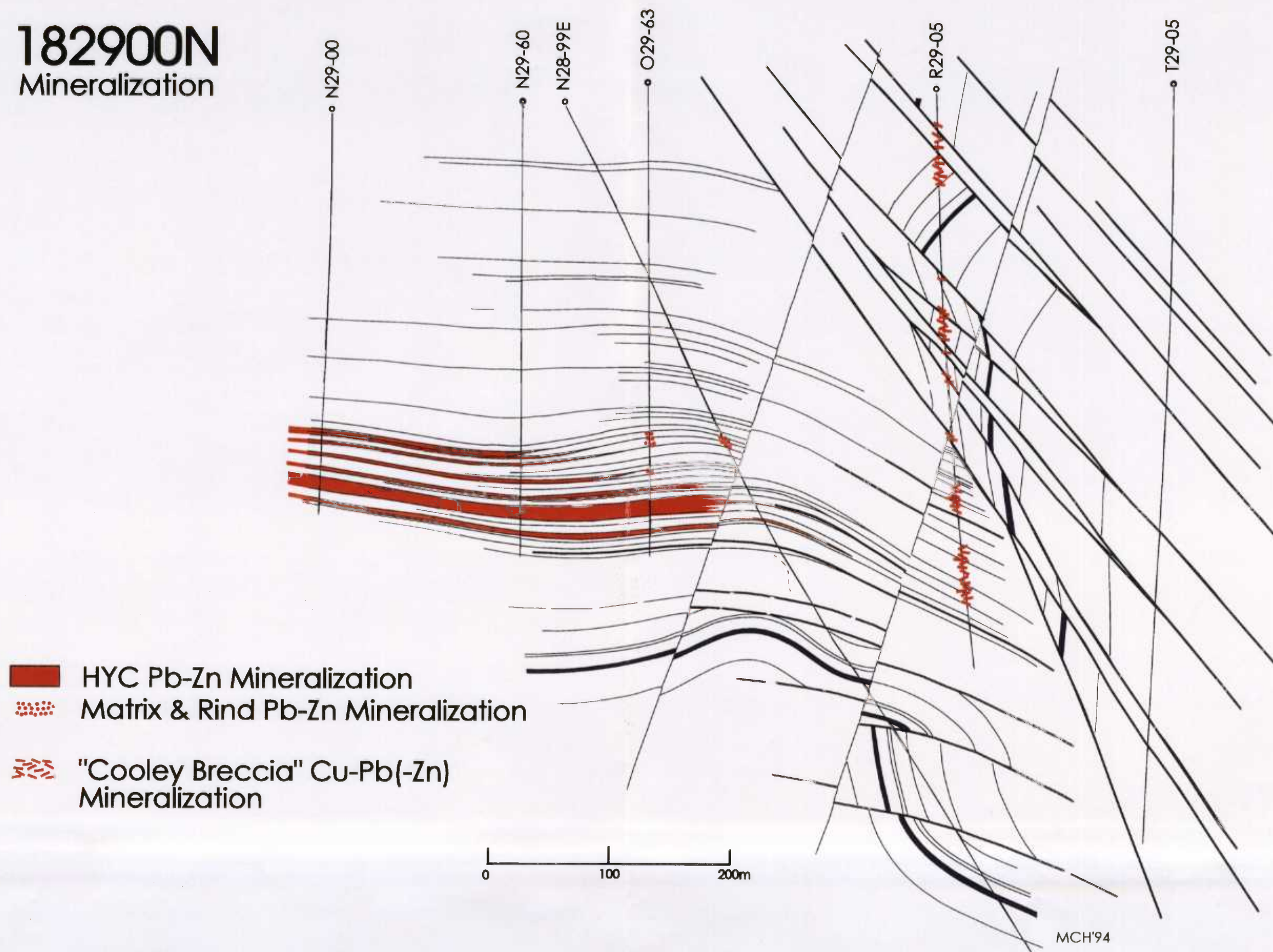


Figure 9. 182900N Mineralisation Styles



* R 9 5 0 0 5 1 2 *

sedimentary breccia cycles. The interore breccia cycles are continuous from west to east where they are sharply truncated by the Basal Cooley Reverse Fault. The third style is the "Cooley Breccia" mineralisation in both the footwall and hangingwall of the Basal Cooley Reverse Fault and is related to the later transpressive inversion event (see below).

Kinematics in the "Cooley Zone"....."Top-to-NW" and "Top-to-SW"

The kinematics of this complex zone between "normal" HYC Barney Creek stratigraphy and the overthrust "Cooley brecciated" Western Fault Block is best resolved from the oriented drillhole N28-99E. The detailed structural data and kinematic indicators are presented in the accompanying N28-99E 1:500 and 1:250 scale Structural Overlays and summarised on the 182900N Structural Overlay (Figs 7 & 8).

The major tectonic and kinematic indicators resolvable in drillcore are:

Cleavage	normal to a shortening direction
Fold Axes	normal to a plane containing a shortening direction
Slickensides	indicating direction of movement on fault and bedding planes
Extensional Veins	indicating direction and sense of movement when associated with planar failure

These indicators suggest two distinct shortening directions and failure movement directions on the shallowly dipping thrust planes (F2) and also (probably; see below) on the steeper reverse fault planes (F3).

Small scale folding (tens of centimetres to a few tens of meters wavelengths), associated with either (pre-failure) shortening or with planes of failure, and slickensides on shallow thrust surfaces at all scales suggest a phase of approximate *SE-NW shortening* with "*Top-to-NW*" movements where failure has occurred. Some slickensides on bedding within the Upper Pyritic and Bituminous Shales are consistent with this kinematic phase and are the only known (at present) expression of this inversion deformation in these units. Other structures that are probably associated with this phase of shortening are the Barney Hill anticline and a parasitic syncline-anticline pair at the base of the exploration decline. A recumbent syncline-anticline pair within the lower Pb-Zn orebodies along the eastern edge of the HYC ore block in the immediate footwall of the Basal Cooley Reverse Fault has a parallel trend in the northern half of the HYC ore block but has a more NNW-SSE trend to the south possibly reflecting the adjacent orientation of the Basal Cooley Reverse Fault there.

A crude macroscopic cleavage is visible within the dolomitic shales in N28-99E where a broad anticlinal arch exists within the lower Mineralised Shales in the footwall of the Basal Cooley Reverse Fault. The poles to cleavage here are fairly scattered, probably as a result of post-cleavage fault related rotation, but suggest a 50-70° east dipping fabric sub-parallel with the Basal Cooley Reverse Fault and consistent with a strong component of E-W shortening adjacent to the transpressively failed Basal Cooley Reverse Fault.

The second phase of this transpressive inversion is expressed by another distinct set of slickensides on shallow thrusts and bedding in the footwall of the Basal Cooley Reverse Fault and reverse faults in its hangingwall (a few that have been resolved). These are consistent with *NE-SW shortening*. Extensional veins at high angles to planes of failure associated with this shortening episode and measurable displacements indicate "*Top-to-SW*" senses of movement.



There are no diagnostic overprinting relationships observed in drillcore to fix the relative timing of these two kinematic phases of transpressive deformation responsible for the formation of the "Cooley Zone". The most logical sequence, however, would have initial SE-NW shortening folding the sequence and associated locally with and culminating in kinematically consistent ("Top-to-NW") failures followed later by the distinctively different phase of "Top-to-SW" failures. It is unclear (at present) whether the overthrusting of the lower dolomitic stratigraphy over the upper HYC stratigraphy initiated at the earlier SE-NW stage or was confined to the later NE-SW kinematic phase. However, it seems probable that the major macroscopic structures (Barney Hill anticline and associated decline folds, major folds centred within the Western Fault Block; Fig. 6) formed during the earlier SE-NW shortening episode. These progressively developed steep to overturned west limbs that then failed transpressively with "Top-to-NW" kinematics, resulting in the reverse fault juxtapositioning of the steeply-dipping, dolomitic Western Fault Block and the shallow-dipping, HYC pyritic shale sequence across the Basal Cooley Reverse Fault. Subsequent NE-SW shortening would have reactivated the Basal Cooley Reverse Fault and other existing footwall thrusts and faults but appears to have generated new structures in the hangingwall. "Top-to-SW" failures have been resolved by "best fit" methods in the overthrust block associated with "Cooley Brecciation" (see below) but the level of data does not preclude some effects of the earlier kinematic failure phase in this zone. Note, however, that SW bedding dips in outcropping Cooley Breccias are more compatible with reverse faulting and overthrusting of folded, NE-dipping lower Barney Creek Formation and underlying dolomitic stratigraphy during NE-SW shortening.

Cooley Brecciation

The "Cooley Breccias" are tectonic breccias formed along the reverse faults within the steep to overturned, brittle dolomitic lithologies of Teena, Mitchell Yard and Mara Dolomite which were overthrust against and over Barney Creek Formation lithologies. The breccias commonly show a zonal arrangement of increasing deformation approaching reverse faults, from fracture networks without any disaggregation, through an angular fragmental zone with minor separation and rotation, through zones of increasing disaggregation and rotation and fragment rounding to zones of significant fragment disaggregation, random orientations and sometimes significant rounding (see Accompanying Section Legend). The reverse faults are commonly marked by an early broad zone of redolomitization / recrystallization sometimes containing silica (chalcedony) blebs (Mitchell Yard) which can contain small scale, pre-brecciation, folding (alteration-recrystallization softening).

The tectonic breccia matrix comprises finer grained (grit-sand-silt size) fragments of the brecciated dolomites that commonly show grading (geopetal) within individual interstices and light to medium to dark grey "shale". This "shale" comprises fine grained dolomite, and pyrite and probably is a sedimented chemical precipitate rather than a true matrix sediment (further work required). The darkest grey matrix material is commonly pyritic and sometimes pyritic-chalcoppyritic.

In the footwall of the Basal Cooley Reverse Fault dolomitization, re-brecciation and copper(-lead) mineralisation occurs within the sedimentary breccias. Deformation probably focussed along shallowly to moderately NE dipping shaly horizons within the breccia pile resulting in re-brecciation, fluid flow and mineralisation in the surrounding breccias. The silty tops of breccia cycles are commonly dolomitized/recrystallised in the footwall of the Basal Cooley Reverse Fault and the "Cooley breccia" style (and generation) of copper mineralisation is most strongly developed within the sedimentary breccias of the Main Talus Breccia and Lower Pyritic Shale units.

Copper Paragenesis

A preliminary paragenesis of the copper mineralisation in the "Cooley Breccias" in the hangingwall of the Basal Cooley Reverse Fault based on core observations is presented in Figure 5.

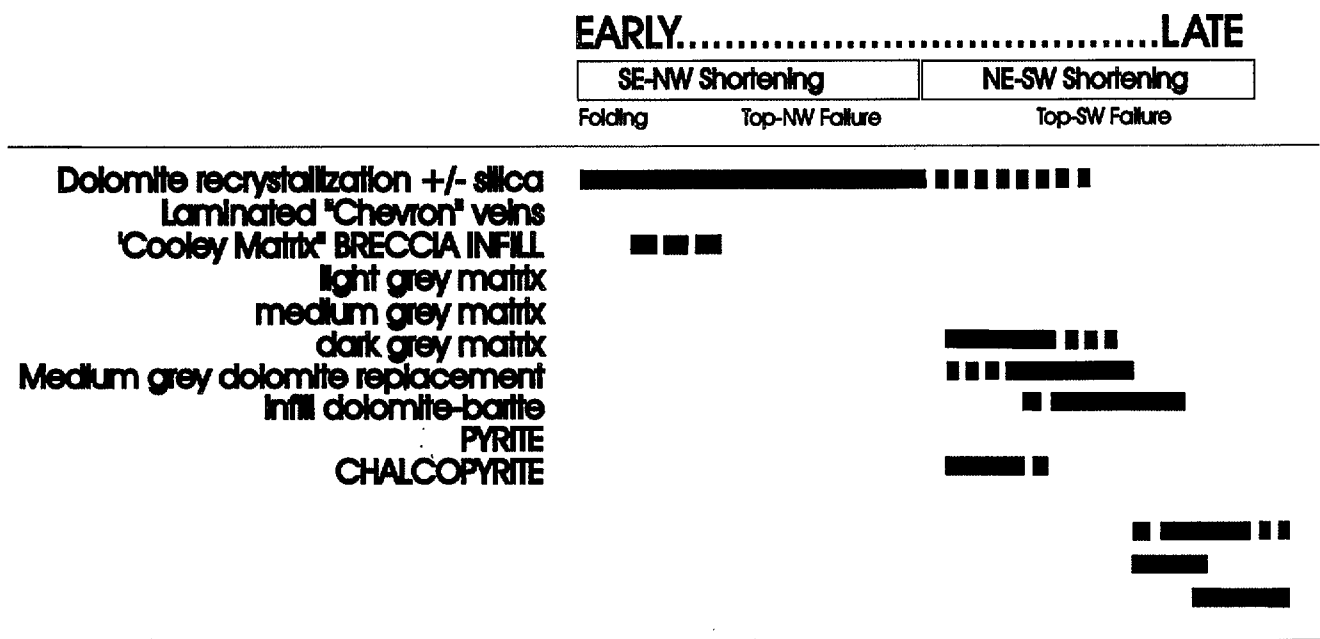


Figure 10. Paragenesis of the Cooley Breccia Copper Mineralization

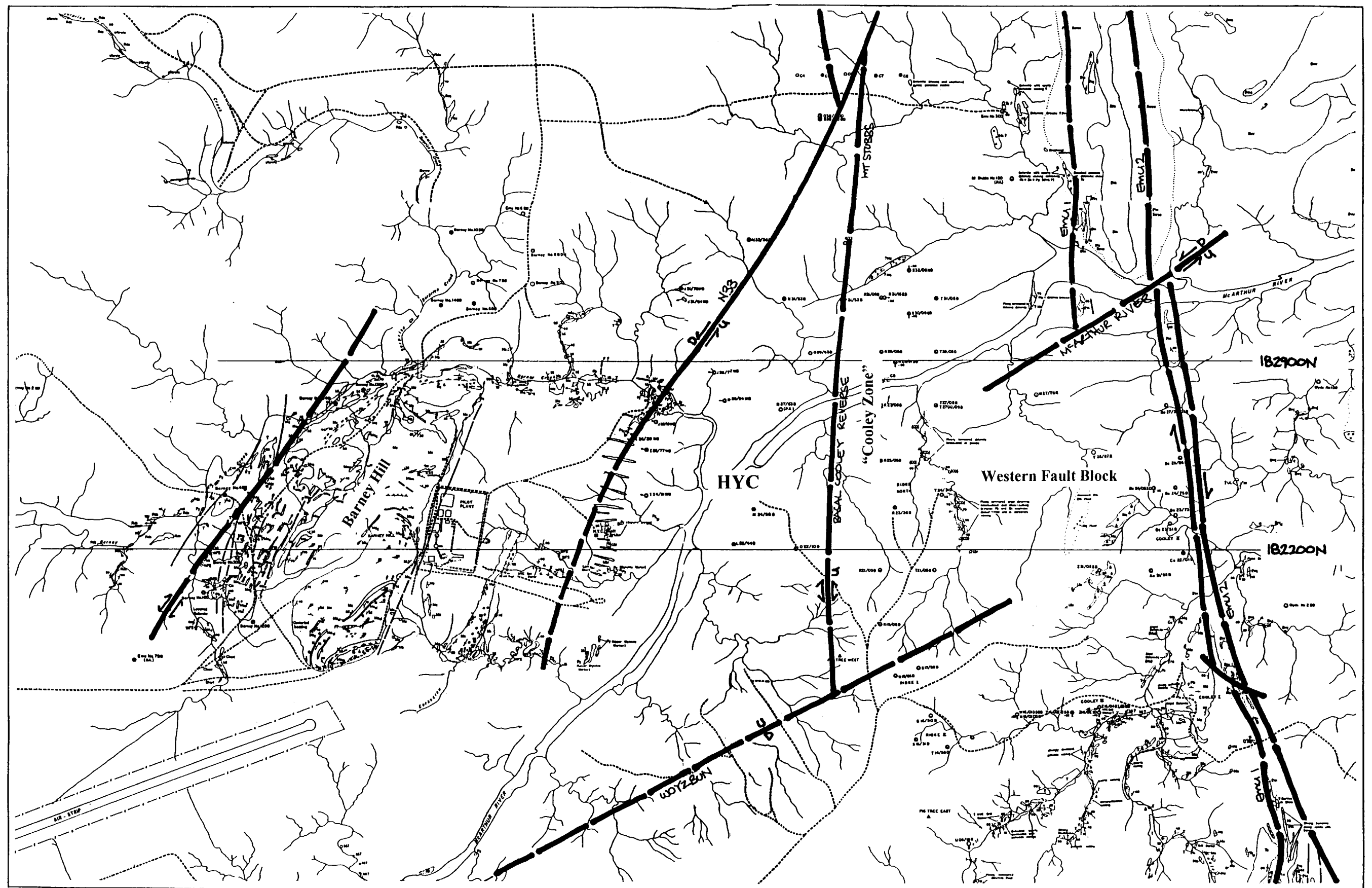
This paragenesis has not been fully fixed relative to the structural-kinematic evolution of the Cooley Zone, however, some constraints incorporated into Figure 10 are known:

- a) the dolomite recrystallization occurs along with the earliest phase of reverse fault failure (most probably NW transpressively directed)
- b) The Chevron Veins are commonly at a high angle to bedding and suggest layer-parallel extension in the early stages of macroscopic folding
- c) the terminal phase of pyrite±chalcopyrite mineralisation is synchronous with "Top-to-SW" kinematic failures. Copper mineralisation associated with extensional veins and ?clay/talc alteration on failed bedding is strongly associated with "Top-to-SW" kinematics on shallow dipping thrusts at the end of N28-99E in the footwall of the Basal Cooley Reverse Fault.

LATE FAULTING

Two late fault sets can be resolved within the drillcore data that are consistent with other mapped faulting at HYC-Cooley. The first set strike NE-SW and have reverse oblique-slip to strike-slip movements (F4 on Structural Overlay; Fig. 8). Faults of this generation include Woyzbun, McArthur River and possibly a reactivation phase of N33 (see Fig. 11). The second set have more NW-SE'ish trends with oblique-slip to strike-slip movements and are correlatives of the Cross Cut Fault from underground and the H20, I20 and I19 Faults from the structural drilling program.

BARNEY HILL-HYC-COOLEY STRUCTURE



Compiled from MIMEX 1:5000 Geology by MCH, June, 1993

Figure 11.



SYNTHESIS and GEOLOGICAL HISTORY

Figure 12 illustrates in cartoon form the structural-kinematic evolution of the Barney Hill-HYC-Cooley-Emu section from the initiation of Barney Creek sedimentation, through the "Cooley" inversion to the regional major N-S shortening episode (probably correlative of Mt Isa D1) to produce the final geometry depicted in Figure 6. The extensional/shortening directions and kinematics indicated on the Figure are appropriate for the local HYC-Cooley domain and do not necessarily reflect regional stress orientations.

Barney Creek Sedimentation

A thick sequence (up to ?800m+) of W-Fold, Lower Dolomitic, Lower Pyritic with sedimentary breccias, Main Talus Breccia, Bituminous, Upper Pyritic and Upper Dolomitic Shales deposited in an Emu-parallel elongate trough. This trough's eastern edge was somewhere east of the present trace of the Emu Fault (5-10km east based on Emu-parallel magnetic structure) and its western edge lay approximately through the present Barney Hill (magnetic edge).

Within the trough (Emu Trough), sedimentary breccias of the Lower Pyritic Shale and Main Talus Breccia flowed, from an entry point north of HYC, to the south and southwest (Logan, 1979). The sedimentary breccias within the Barney Creek Formation record the progressive exhumation of the pre-Barney Creek dolomitic McArthur Group stratigraphy. This implies that the province region for the sedimentary breccias (somewhere adjacent and close to this Emu Trough) was being actively uplifted and eroded during Barney Creek-time. Tectonically this is not possible in a purely extensional regime: it seems likely that this Emu Trough zone was transpressive-transensional throughout much of Barney Creek-time (Fig. 12) and that transpression ultimately inverted the HYC Emu Trough forming the Western Fault Block and the "Cooley Breccias".

The nature of the western margin of the Emu-parallel trough is unknown. It has not apparently reactivated during later deformations in the way the Emu Fault has, suggesting that this boundary of the trough may be a hinge line rather than a major fault or set of faults.

Subsidence within the HYC portion of the Emu Trough west of the present Emu Fault was most rapid in the NE. Breccias, which thin to the south and west in the HYC area, are envisaged to have filled the additional space created by the heterogeneous subsidence. As a result of this differential subsidence, buried sedimentation surfaces (bedding) within the diagenetically maturing pile adopted gentle and steepening NE dips. The zone was probably seismogenic which triggered failures on bedding producing the kinematically dominant "Top-to-NE" compressional and extensional structures (see below; Fig. 4). These structures post date a number of diagenetic features including crusts, concretions, biogenic pyrite and probably Pb-Zn mineralisation (Fig. 5) all of which formed at considerably shallower (some close to or at zero) depths of burial.

These observations are consistent with a deviation from the regional tectonic framework during Barney Creek-time (Etheridge & Wall, 1994) in which Emu parallel structures become transpressive-transensional. Local transtension at HYC within the active Emu zone resulted in more rapid subsidence and the accumulation of the anomalously thick Barney Creek HYC sequence within the Emu Fault Trough. Transtension across the HYC basin bounding faults would have made them dilational and have aided base metal bearing hydrothermal fluid access.



Pre-Barney Creek (McArthur Group) time
N-S Extensional Geometry (Etheridge & Wall)
early syn-diagenetic dolomitization

Start of Barney Creek time
W-Fold Shale
Transpressive-Transensional EMU ZONE

Lower HYC Pyritic Shale time
HYC Rapid Subsidence
Elsewhere uplift - NE sedimentary breccia source
Syn-diagenetic HYC Pb-Zn Mineralization

"Main Talus Breccia" time
NE Maximum Subsidence - TNE Bedding Failure
Warping - sedimentary breccia channelling

Upper Styrolitic Marker time ~1640Ma
NE Maximum Subsidence - TNE Bedding Failure
Folding - sedimentary breccia channelling

Upper Pyritic & Bituminous Shale time
NE Maximum Subsidence - TNE Bedding Failure
Overturned Short Limbs

Upper Breccia time
Short Limb Transpressive Failure - TNW Kinematics
Upper Breccias sourced from Western Fault Block
Cooley Brecciation - (re)dolomitization
(HYC North Cu-Pb Mineralization)

2nd Phase Transpressive Failure - TSW Kinematics
"Cooley Breccia" Cu-Pb Mineralization
in Western Fault Block & "Main Talus Breccia"

Upper Barney Creek ONLAP
Reward Dolomite Sedimentation
Batten Subgroup Sedimentation

Regional N-S Deformation (Isan D1 ~1600Ma)
Emu Fault Zone - DEXTRAL STRIKE SLIP
NE faults - oblique to strike slip
Vein MVT Cu-Pb Mineralization
Cooley Cu, Cooley Pb adjacent to EFZ

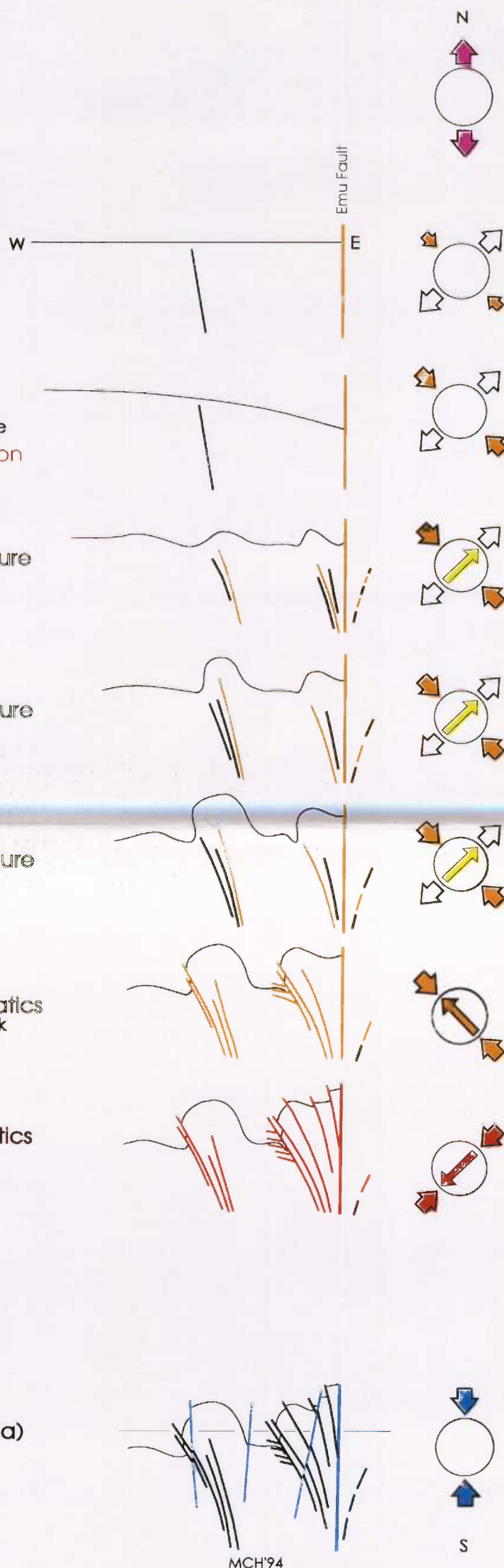


Figure 12. Structural-Kinematic Evolution of the Barney Hill - HYC-Cooley Section



Transpressive Inversion and Shortening

During continued deposition of the Barney Creek Formation, the HYC area began to invert. This transpressive inversion focussed (1) along the Emu Fault (reactivating it as a transpressive, positive flower structure), (2) along some NE-trending structure under Barney Hill and (3) possibly also along the western margin (?hinge line) of the Emu Trough. Shortening was approximately SE-NW directed and produced the NE trending Barney Hill Anticline and associated parasitic folds and more Emu-parallel folds closer to the Emu Fault (shear component of strain taken up along the Emu Fault resulting in local shortening being more normal to the structure). A crude cleavage nucleated in shaly units closer to the Emu Fault where shortening was more intense. The folds in the Lower HYC sequence and underlying dolomitic stratigraphy at Barney Hill and at Cooley progressively tightened, while Barney Creek sedimentation continued at the sediment-water interface above, until they developed overturned west limbs which subsequently failed.

"Top-to-NW" and "Top-to-SW" Failure

Ultimately the western short limbs fail. The gross geometry of the major structures is not sufficiently clear to resolve whether the shallow thrust failures, which are mostly "Top-to-NW" kinematically, pre-date or are synchronous with the steeper reverse transpressive faults of the overthrust "Cooley Zone". What is clear is that there are two kinematic phases within the shallow footwall thrusts with the "Top-to-NW" preceding "Top-to-SW" kinematics and well developed "Top-to-SW" kinematics in the steeper reverse faults within the "Cooley Zone". Dolomitization and recrystallization along the major reverse faults and thrusts is associated with the earlier phase of failure while copper mineralisation appears to be linked with the "Top-to-SW" phase of movement in the footwall zone (more detailed oriented core work required within hangingwall "Cooley Zone").

In a regional tectonic sense it is unclear what controls this rotation of the local shortening direction and kinematic transport direction, however, it can be rationalised in terms of the interaction of regional scale structures (Etheridge & Wall, 1993; unpubl. company report). The NE-SW phase of shortening could result from an initial pulse of NS shortening modified locally at HYC by synchronous right lateral reactivation of the Emu Fault. This shortening phase most probably produces the opposite-sense, "Top-to-SW" low angle structures within the HYC ore sequence (Fig. 4).

Upper Breccia sedimentation is hypothesised to be linked with these transpressive failure phases. The formation of a positive flower structure along the reactivated Emu Fault would provide a more local source (Logan, 1979) for the Teena to Masterton lithologies present within the Upper Breccias deposited west of the "Cooley Zone".



McArthur Sedimentary Architecture

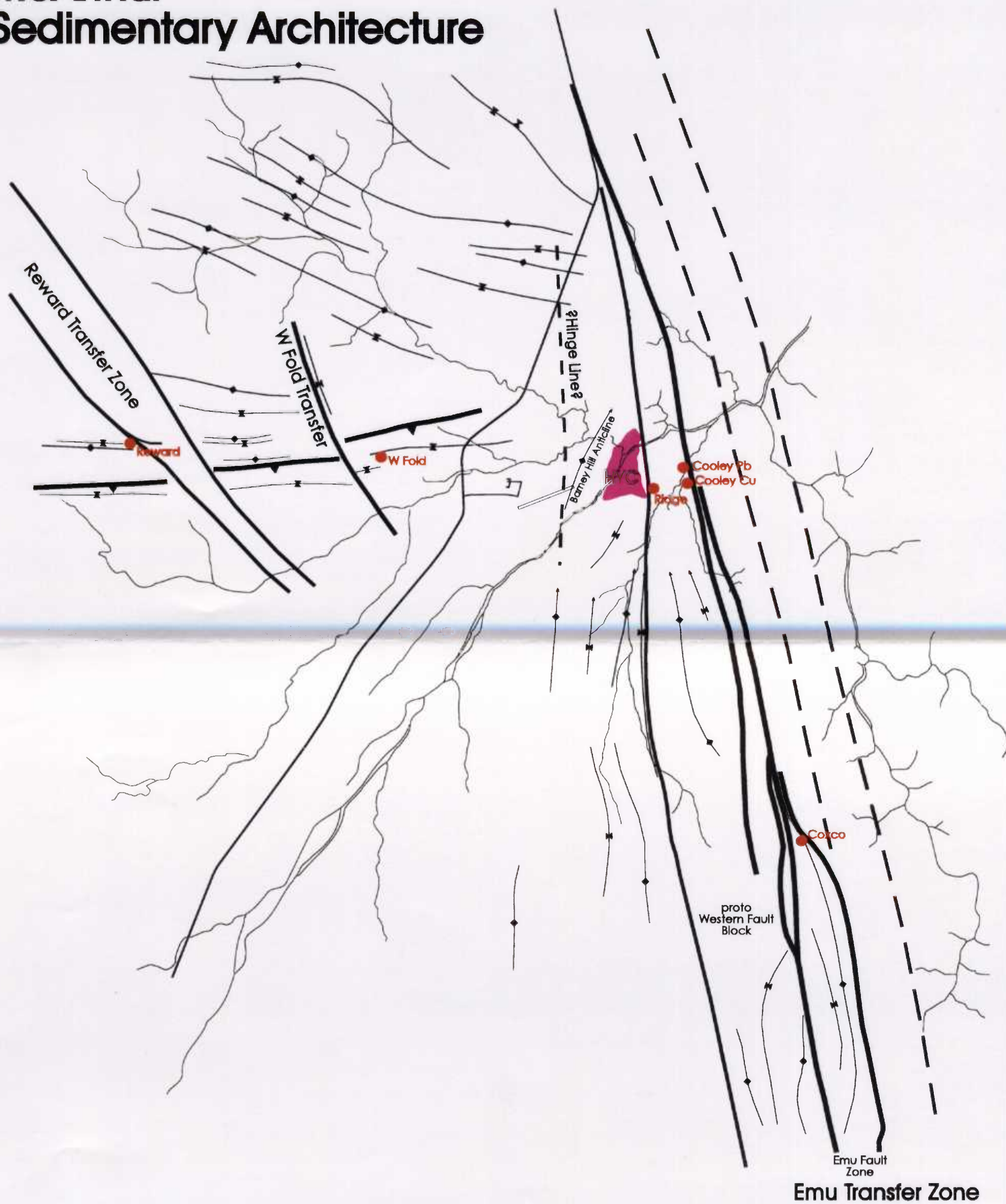


Figure 13. The Sedimentary Architecture at McArthur River

* R 9 5 0 0 5 1 8 *

184700N Geology

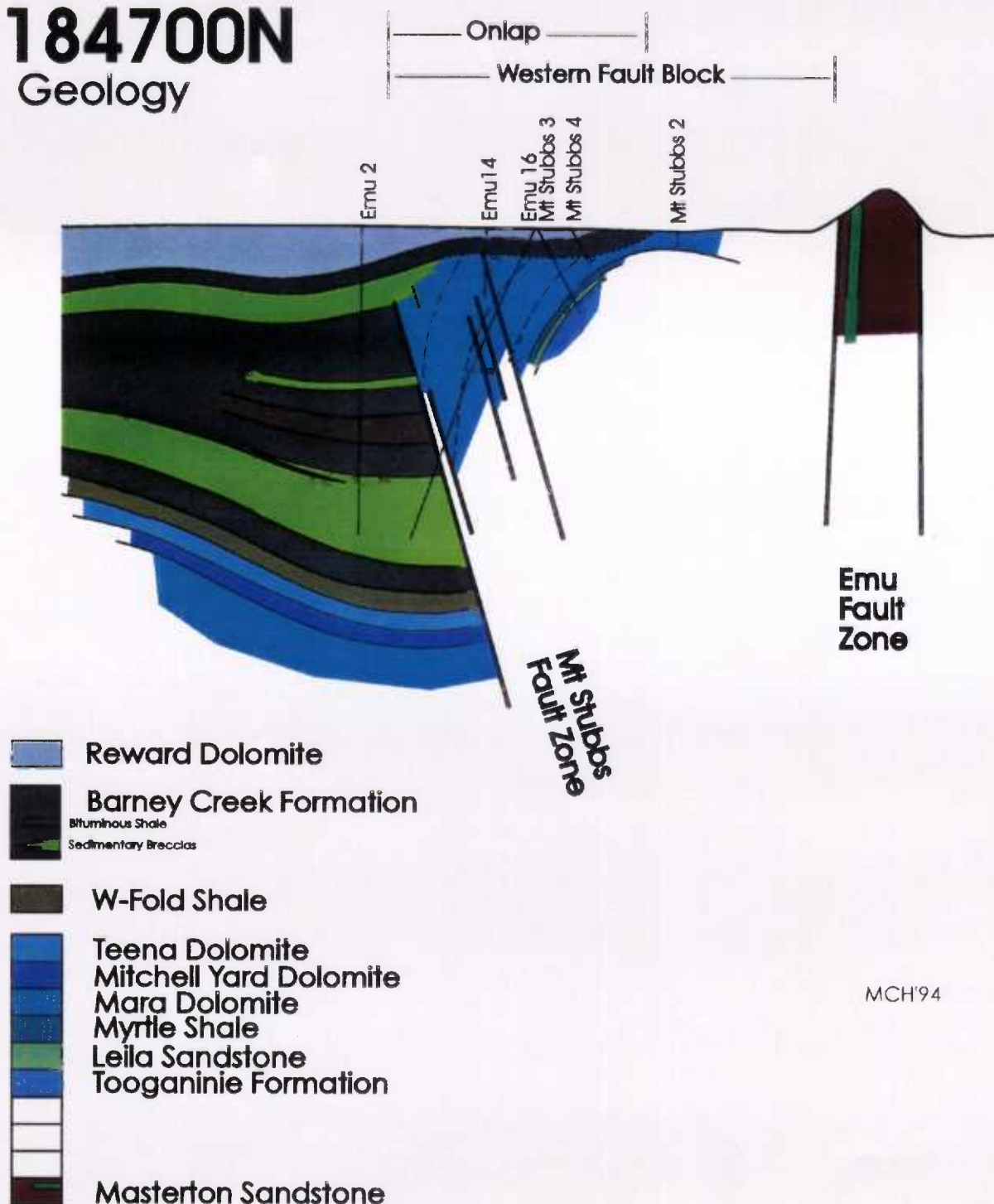


Figure 14. Section 184700N Geology



* R 9 5 0 0 5 1 9 *

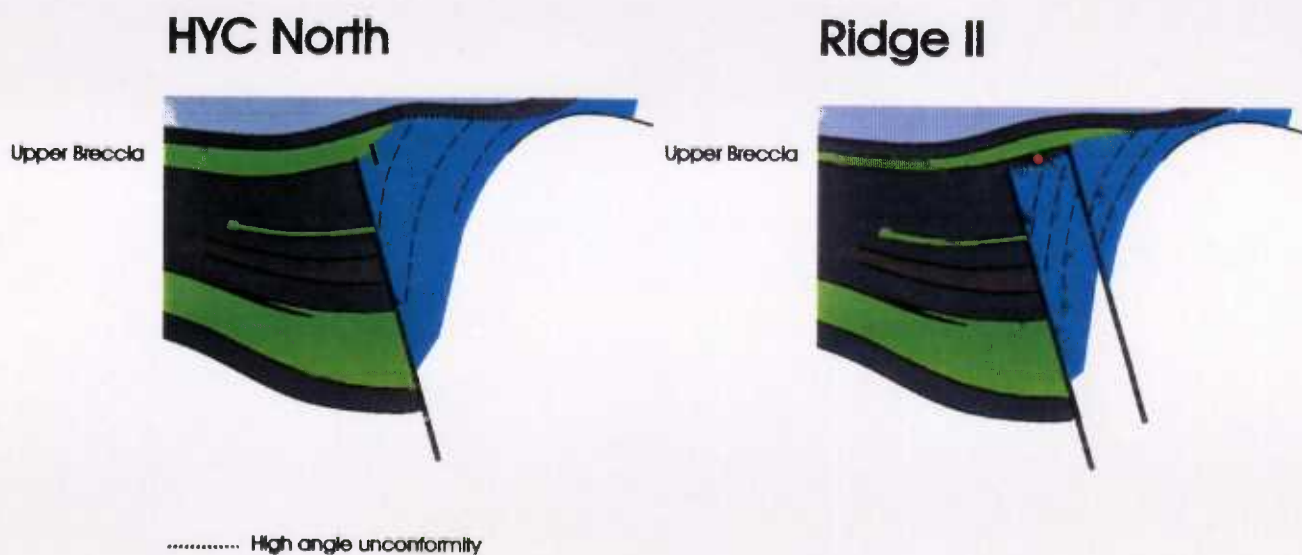


Figure 15. Schematic Onlap Geometries

Upper Barney Creek Onlap, Reward Dolomite and Batten Subgroup Sedimentation

As the transpressive deformation quietened, Upper Breccia sedimentation passed into Upper Barney Creek sedimentation at HYC. West of HYC breccias are not present and HYC shale sedimentation passes conformably into Reward Dolomite sedimentation. North of HYC, condensed? (atypical) Upper Barney Creek sediments were deposited with a high angle unconformity over the steeply-dipping, upthrust, redolomitized and mineralised Mara Dolomite lithologies of the "Cooley Zone" or Western Fault Block (Fig. 14). Present drilling geometries indicate that north of HYC, Cooley transpression-generation structures do not cut Upper Barney Creek and Reward Dolomite sedimentation (locally later structures and reactivations may do so; Fig. 14).

It is suggested that onlap onto the "Cooley Zone" may have developed in other localities *during* Upper Breccia deposition if the westernmost reverse fault became inactive while reverse structures further east continued to be transpressively active (Fig. 15). The relationships at Ridge, where strataform, shale-hosted Pb-Zn mineralisation at the base of the Upper Breccias ?unconformably overlies Cooley-brecciated and mineralised (probable) Teena Dolomite suggest this (Fig. 15).

Reward Dolomite and the Batten Subgroup was deposited over the Barney Creek Formation with transitional conformable relationships around McArthur-HYC. These unit probably completely overlapped the "Cooley Zone", Western Fault Block and Emu Fault Zone which was inactive at this time.



North-South Shortening ca.1600Ma

North-south (or probably more NNE-SSW) shortening folded the Reward and Batten Subgroup units and reactivated the Emu Fault as a right lateral strike-slip structure cutting these units probably during the regional Isan D1 deformation. Prominent east-west to WNW-ESE fold axes northwest of HYC (Figs 1 & 13) reflect this deformation. Closer to the Emu Fault more Emu-parallel folds formed due to the local strike-slip shear release along the Emu Fault. North-south thrusting also occurred locally on shallowly dipping contacts between units (north of HYC; Figs 1 & 13).

The Cooley Lead mineralisation immediately adjacent to the Emu Fault lies along a shallowly dipping, brecciated contact between Mara and Tooganinie and steeply north-dipping dilational veins in Mara above the brecciated zone. Bedding offsets across the veins are consistent with North-Block-Down and suggest "Top-to-South" movement of Mara over Tooganinie adjacent to the Emu Fault. This mineralisation and probably also Cooley Copper are therefore timed to be synchronous with the right-lateral reactivation of the Emu Fault during the regional (Isan D1) NS shortening event (ca. 1600Ma).

During this NS shortening at HYC-Cooley(-Emu) NE-trending faults showing oblique reverse to strike-slip movements (Woyzbun, N33 reactivation and McArthur River Faults; Figs 1 & 11) formed or reactivated. Some faults of this generation cut the Emu Fault (eg. McArthur River Fault).



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APPENDIX

Drillhole Structural Logs of N28-99E, I22-55, K22-55 and R29-05.

HOLE	SURVEY DATA			Dip/Azim		LOCAL SO		(C/O)		Sa		(C/O)		Sb		(C/O)		So		La		90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90	90
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[illegible]

[illegible]

745.0 58.7 90.0

765.5 58.7 90.0 47 170 17 64

770.0 58.8 90.0 50 60 21 60
775.6 58.8 90.0 57 60 28 60

777.9 58.8 90.0
781.3 58.9 90.0
781.8 58.9 90.0 40

783.3 58.9 90.0 65 125 51 17
785.4 58.9 90.0 72 160 43 62
785.8 58.9 90.0 70 150 45 48
786.0 58.9 90.0 72 160 43 62
787.0 58.9 90.0 62 135 44 25
788.0 58.9 90.0 60 145 37 35

789.2 58.9 90.0 60 175 29 81
789.4 58.9 90.0 60 180 29 90
790.7 59.0 90.0

791.0 59.0 90.0 55 190 25 110
792.4 59.0 90.0 55 165 26 61
793.3 59.0 90.0 57 175 26 80

795.0 59.0 90.0 55 165 26 61
796.8 59.0 90.0 60 155 34 49
799.0 59.0 90.0 60 180 29 90

801.0 59.1 90.0 60 175 29 81

795.0 59.0 90.0

cV1 75 -75 76 354 cV2 75 115 83 198

bseF? 55 53 207
F 55 53 207
F 66 53 207

F 75 135 55 33

cV 60 -90 64 197
pycV 75 -60 88 33
FLT 65 -70 78 210
FLT 50 145 29 25

Jf 65 85 71 342

SSlds -20 17 62

SSlds 25 14 107
SSlds 80 4 337

FSlds -55 1 289
FSlds -80 22 274

FSlds -85 25 323

FSlds 20 75 169
FSlds 27 12 93

JSlds 5 48 50

15 90

18 90 25 270 69 270
25 90

31 270
31 270
31 270

20 90
40 90 38 90
36 90
40 90
22 90

24 90 31 270
29 90 87 90
29 90 67 270
13 90

24 90
23 90
26 90 42 270
23 90
27 90
29 90
29 90

HOLE	SURVEY DATA			DPTH	PLNG	DIR	d	Dip		LOCAL S0		?	d	a	Sa		?	d	a	Sc		?	B	La		90	S0	90	Sa	90	Sh	90	Sc			
	Dpth	Plng	Dir					Azimuth	90	Dip	Dir				Dip	Dir				Dip	Dir			Plng	Dir	Dip	Dir	Dip	Dir	Dip	Dir	Dip	Dir			
122-55	0.0	90.0	270.0	24.0	89.5	274.5	20	90	20	90																										
				26.6	89.5	275.0	21	90	21	90	mf	60	140	60	230																					
				31.2	89.4	275.9	10	90	11	90	mf	58	105	58	194																					
				39.7	89.3	277.5	9	90	10	90	mf	70	110	70	199																					
				45.0	89.2	278.5	10	90	11	90	mf	65	85	65	174																					
				47.7	89.1	279.0	18	90	19	90																										
				52.0	89.0	279.8	14	90	15	90																										
	53.0	89.0	280.0	53.6	89.0	279.5	12	90	13	90																										
				58.5	88.6	275.3	13	90	14	90	V	87	-130	86	320																					
	60.0	88.5	274.0	63.0	88.5	273.0	17	90	19	90																										
				65.9	88.4	272.1	25	90	27	90																										
				70.3	88.4	270.7	18	90	20	90																										
				71.6	88.4	270.3	45	90	47	90																										
				72.5	88.4	270.0	35	90	37	90																										
				80.0	88.3	267.6	18	90	20	90																										
				82.4	88.2	266.9	18	90	20	90																										
				85.1	88.2	266.0	7	90	9	90																										
				86.3	88.2	265.6	23	90	25	90	V	75	-140	74	311																					
				94.3	88.1	263.1	13	90	15	90																										
				95.5	88.1	262.7	16	90	18	90																										
				96.1	88.1	262.5	5	90	7	90	V	80	120	79	213																					
				98.7	88.1	261.7	14	90	16	90	V	85	55	86	146																					
	104.0	88.0	260.0	114.9	88.0	262.3	15	90	17	90																										
				118.0	88.0	263.0	12	90	14	90	VF	75	-95	75	357																					
				119.7	88.0	263.4	14	90	16	90																										
				128.8	88.0	265.3	15	90	17	90																										
				134.5	88.0	266.5	14	90	16	90	CFV	76	-20	78	71																					
				135.4	88.0	266.7	8	90	10	90																										
				132.8	88.0	266.2	10	90	12	90	T	30	10	32	100																					
				136.6	88.0	267.0	6	90	8	90																										
				137.5	88.0	267.2	14	90	16	90																										
				137.9	88.0	267.3	5	90	7	90																										
				139.0	88.0	267.5	10	90	12	90	V1	90	65	89	336	V2	90	-20	88	250																
				140.1	88.0	267.7	11	90	13	90	T	40	-75	41	18	sgV	90	-130	89	320																
				141.1	88.0	268.0	12	90	14	90	F	27	-35	29	57																					
				141.7	88.0	268.1	5	90	7	90																										
				142.5	88.0	268.3	6	90	8	90	E	27	0	29	90																					
				142.2	88.0	268.2	11	90	13	90																										
				143.1	88.0	268.4	11	90	13	90	E	27	-20	29	71																					
				145.3	88.0	268.9	7	90	9	90																										
				145.6	88.0	268.9	12	90	14	90	T	10	-170	8	283																					
				146.5	88.0	269.1	8	90	10	90	E	28	-115	27	339	E	24	95	24	181																
				147.1	88.0	269.2	6	90	8	90	Enf	47	-5	49	85																					
				148.0	88.0	269.4	8	90	10	90	E	25	-140	24	313	T	40	-55	41	37																
				148.6	88.0	269.6	17	90	19	90	T	26	-135	25	318																					
				148.8	88.0	269.6	7	90	9	90	T	65	30	67	120	sgV	77	65	78	155																
				150.0	88.0	269.9	17	90	19	90	sgV	80	85	80	175																					
				151.4	88.0	270.2	5	90	7	90																										
				151.6	88.0	270.2	9	90	11	90	E	52	13	54	103																					
				152.0	88.0	270.3	7	90	9	90																										
				152.3	88.0	270.4	11	90	13	90	E	45	5	47	95																					
				152.7	88.0	270.4	9	90	11	90	TR	27	-45	28	48																					
				152.9	88.0																															

[illegible]

HOLE	SURVEY DATA			DPTH	PLNG	DIR	d	Dip	LOCAL S0	Dip	S0	Dir	?	d	a	Sa	Dir	?	d	a	Sb	Dir	?	d	a	Sa	Dir	?	B	Plng	Dir	90	S0	90	Sa	90	Sb	90	S0																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																
	Dpth	Plng	Dir					Azin=85								Dip					Dir					Dip														Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip	Dip

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