

Application of geochemical modeling to ore systems Lead and Zinc Solubility in skarn deposits

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INTRODUCTION

This project represents an attempt to clarify issues related to ore deposition using an approach that combines geochemical modeling and fluid inclusion data. The project is proposed as part of the pmd*CRc projects F1 (Fluid chemical paths in ore forming processes) and F3 (Micrometallurgy of hydrothermal fluids).

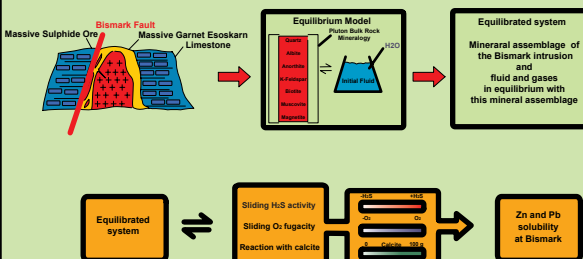
The first part of the project has focused on the study of the Bismark deposit (northern Chihuahua, Mexico). This deposit has been chosen considering its well understood geological setting and the availability of fluid inclusion data. The same approach will be applied to Iron oxide Cu-Au deposits in the Cloncurry district.

Background

The Bismark deposit (northern Chihuahua, Mexico) is an example of stock-contact Zinc-rich skarn. The deposit contains Zn (8%) with minor Pb (0.5%), Cu (0.2%) and Ag (50 g/t) in massive sulfides. The main aim of this study is to understand the mechanisms that control Pb and Zn solubilities at Bismark.

METHOD

Geochemical modeling-GWB



Zn and Pb solubilities were tested setting galena and sphalerite in the initial system in equilibrium with the Bismark stock. This approach was used to evaluate variation in Zn and Pb solubilities induced by oxygen fugacity and H₂S activity changes. For the reaction with the carbonate host rock the initial system contained Pb and Zn as aqueous species.

Sliding H₂S activity

Constant T (550°C)

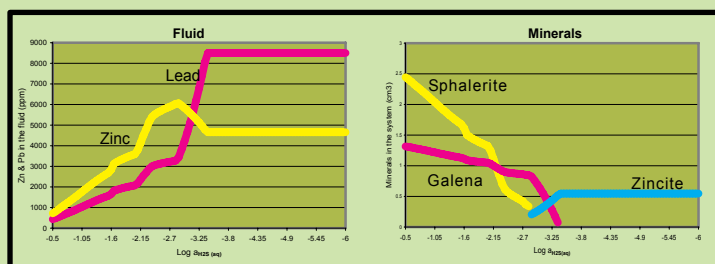


FIG.1 Zn and Pb solubilities generally increase as H₂S activity decreases. The higher Pb content in the fluid relative to Zn for low H₂S activities is caused by precipitation of zincite, and sphalerite is not present in the system.

Cooling (550-100°C)

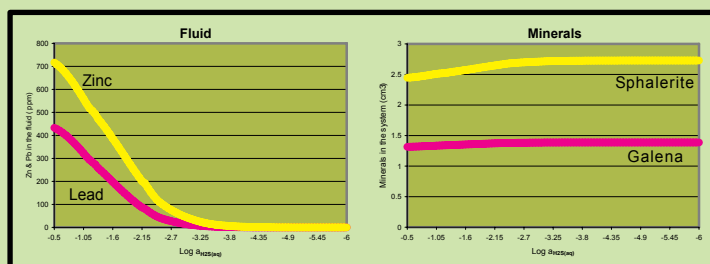


FIG.2 Lead and zinc solubilities generally decrease with decreasing H₂S activity. Lead predominates over zinc in the fluid, while sphalerite is more abundant than galena.

Reaction with calcite

Constant T (550°C)

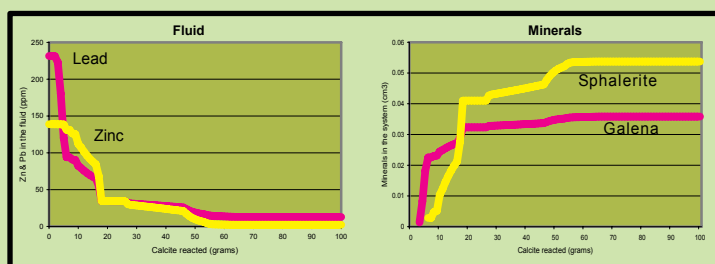


FIG.3 Initially the Pb fluid content drop rapidly, causing a relative Zn enrichment. As the reaction proceeds the fluid is depleted in Zn, and Pb becomes more abundant. Sphalerite predominates over galena.

Cooling (550 - 100°C)

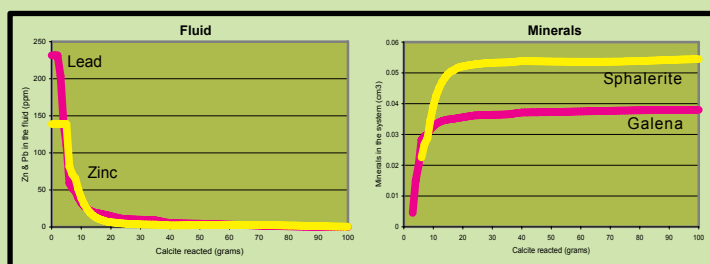


Fig.4 Initially the solubility trend for Zn and Pb is similar to that obtained for the reaction conducted at constant temperature. As the reaction proceeds the Zn and Pb content becomes similar. Sphalerite predominates over galena.

Sliding O₂ fugacity

Cooling (550 - 100°C)



FIG.5 This different behaviour displayed by Zn and Pb at low oxygen fugacity is a consequence of the precipitation of anglesite, without formation of any Zn-bearing mineral. This caused the fluid to be enriched in Zn relative to Pb.

PRELIMINARY CONCLUSIONS

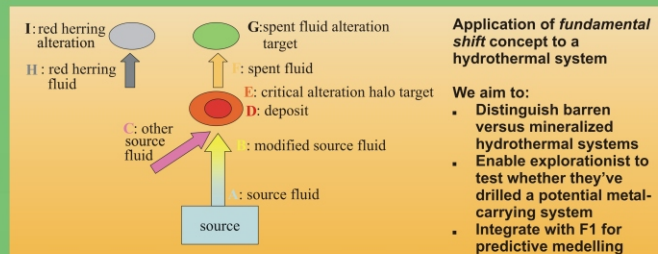
Preliminary results suggest that Zn and Pb solubilities at Bismark were mainly controlled by the reaction between the magmatic fluid and the carbonate host rock. The role of other factors (oxygen fugacity, H₂S activity) seems negligible.

Participants:

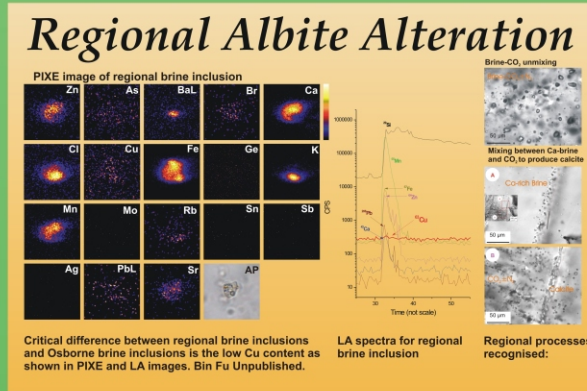
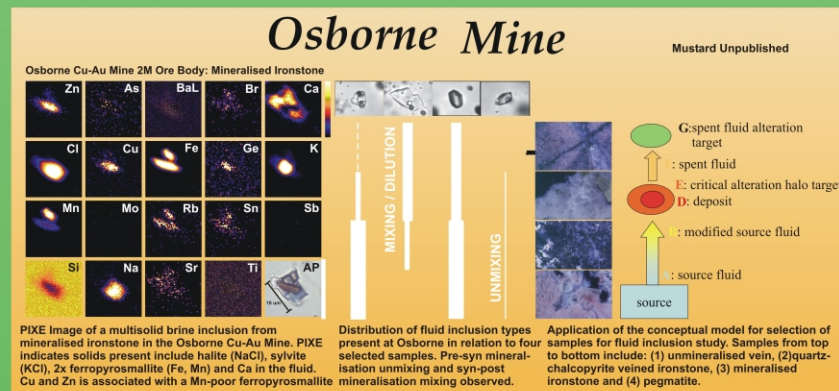
Tim Baker, Pat Williams, Roger Mustard, Bin Fu, Nick Oliver (JCU), Terry Mernagh (GA), Chris Ryan, Esme van Achterberg (CSIRO), Thomas Ulrich, John Mavrogenes (ANU), Joel Brugger (SAM); pmd*CRC Links: Georgie Mark, Andy Wilde (Monash), Steffen Hagemann, Peter Neumayr (UWA), Dave Phillips, John Miller (University of Melbourne) .

Project Aims:

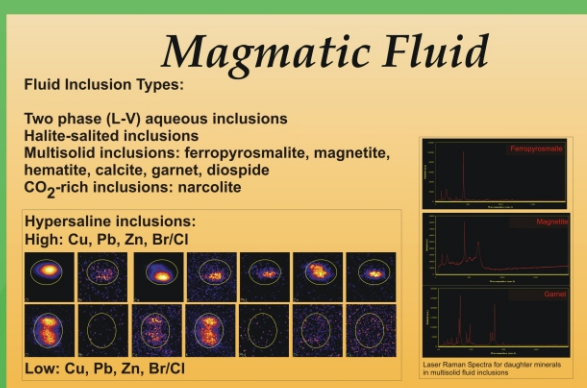
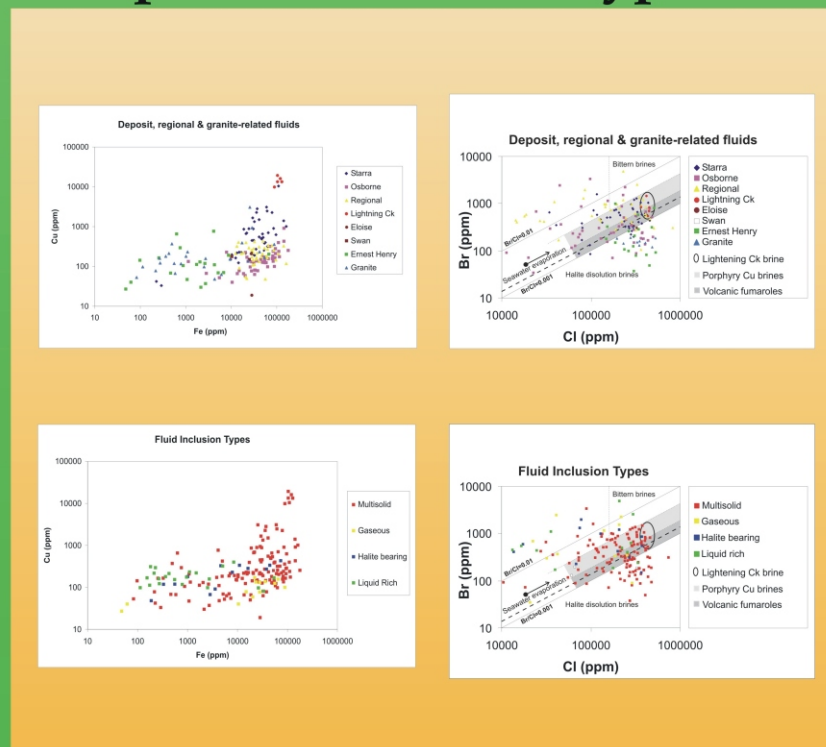
- To produce a fully integrated approach to characterising hydrothermal fluids (fluid inclusions)
- Driven by geology - terrane and ore system based studies (I2/I4/Y3)
- Through the development of advanced analytical methods (integration of enabling technologies such as PIXE, LAICPMS, Laser Raman)
- For use in exploration & fluid modelling (F1)



Regional, Magmatic and Deposit Fluid Types



Comparison of Fluid Types:



Discussion:

- Major and trace element compositions of fluid inclusions have been collected (using PIXE and LA-ICPMS) for a range of separate Fe-Oxide Cu-Au systems, regional alteration and intrusions including Lightning Creek & Mt Angelay (n= 245).
- Comparing fluids from a range of different systems will assist us in: (1) Discriminating between potential carrier and spent fluids, (2) Understanding the role of key fluid processes such as mixing and unmixing in the mineralisation process, and (3) Determining potential fluid sources.

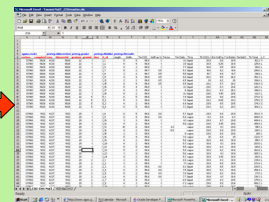
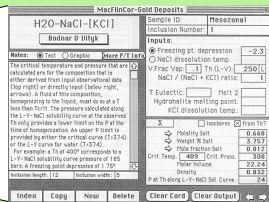
AN INTEROPERABLE FLUID INCLUSION DATABASE

TO RECORD OBSERVATIONS AND CONSTRAIN CHEMICAL MODELLING

Terry Mernagh, Stuart Girvan, and Lesley Wyborn
Geoscience Australia, GPO Box 378, Canberra, ACT 2601 Australia

The Way We Used To Do Things

Fluid inclusion studies were carried out in isolation by individual researchers. The data was kept on their own computer. Often this data was not backed up and was lost when they left university or got another job.

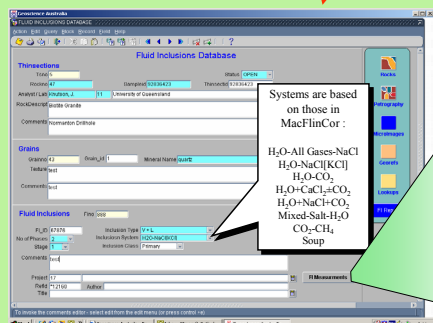
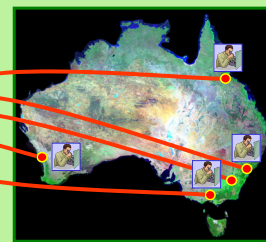


Observations could be recorded in MacFlinCor but you had to have a Macintosh computer to use this program.

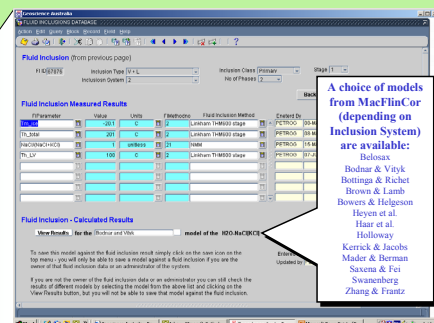
The data had to be exported into Excel for further massaging and for plotting isochores.

The Way We Now Do Things

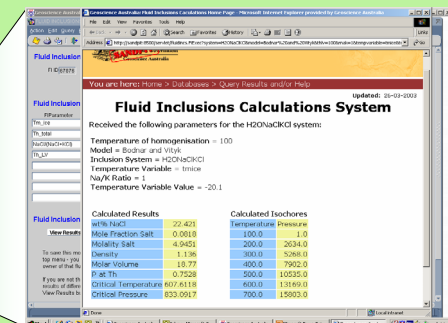
pmd*CRC researchers across Australia can now enter their data into the Fluid Inclusion Database at Geoscience Australia to build a library of ore fluids that can be used to Constrain geochemical modelling.



This screen allows you to record descriptions of the thin section and the individual fluid inclusions.



This screen allows you to enter your observations and choose a model for calculating the fluid properties.

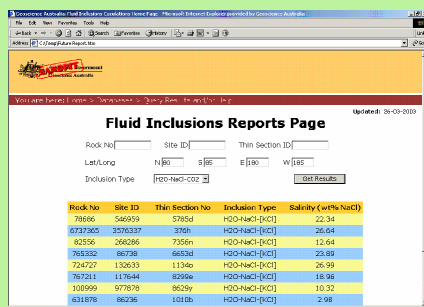


This screen displays the calculated results for the chosen model in a web browser. You can simply change models and recalculate the results without changing the observed fluid inclusion parameters.

The Way We Plan To Do Things In The Future

pmd*CRC researchers will be able to create reports of fluid inclusion data over the internet. The data set reported will be defined by a number of criteria (for example spatial boundaries) set by the user and will include both primary and derived data.

Applications which are written to handle web feature service data will be able to read data in real time.



There will be a report screen where you will be able to report both primary results and derived/modelled data as whole data sets

Fluid Inclusions Database

Web Reports Application

Core Modelling Routines

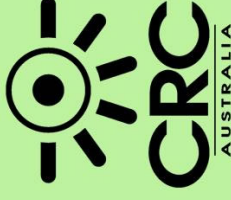
Web Feature Service



Applications will access both primary and modelled data in real time



predictive mineral discovery **COOPERATIVE RESEARCH CENTRE**



Exploring Fluids in Fe oxide-Cu-Au deposits

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Figure 1: Map showing the location of the Wernecke Mountains, Canada and the distribution of breccia sites and associated mineral prospects.

Wernecke Geology

Dominant rocks

- Early Proterozoic Wernecke Supergroup(1840-1725 Ma) Siltstone, shale, dolostone, sandstone
- numerous breccia zones (1595 Ma)
- widespread hydrothermal Cu-Co-Au-U mineralisation related to the breccias
- no known economic deposits

Fluid source?

Intrusions

- dikes @ 1720 and 1270 Ma
- no intrusions @ ~1600 (assoc with breccia and mineralisation)

- ### **Introduction**
- #### **Fe oxide-Cu-Au deposits**
- Very large masses of iron oxide (hematite/magnetite)
 - Economic components: copper, gold (uranium, REEs, cobalt, nickel)
 - Very large scale metasomatism

Fluid source debate

- Major theories
- Magmatic
- Evaporitic
- Mixture (incl meteoric)

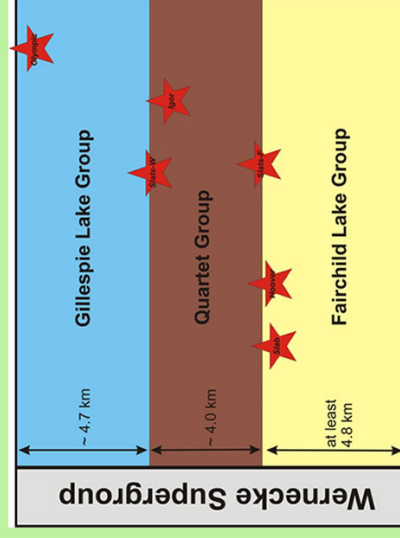


Figure 2: Relative stratigraphic positions of the study sites/prospects. The relative depths of the prospects and differences in the rock types in which each prospect lies, may have an influence on fluid properties.

Wernecke vs Cloncurry

Similar types of deposits

- Large hydrothermal systems
- Brecciation
- Age

Key differences

- Granitic intrusions
- Fluids?
- Grade of metamorphism/ P-T conditions
- Economic deposits

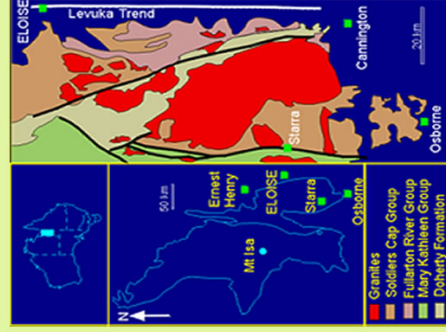


Figure 3: Map of the Cloncurry district, showing deposit locations. Note the dominance of granite intrusions (red) in the area. Such intrusions are notably absent in the Wernecke region.

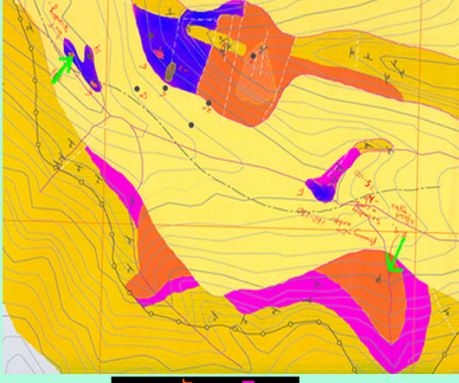
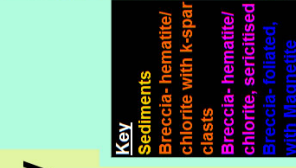


Figure 4: Basic map of the Igor prospect showing the occurrence of breccia zones within the Wernecke Supergroup sediments.

-possible intrusion at depth?

Evaporites

- not obvious
- scapolite may be indicative of evaporitic influence

Previously no fluid inclusion data to indicate fluid source

Important Questions with implications for exploration

- Which of these "key differences" are responsible for the lack of economic deposits in Wernecke Mountains
- eg. What is the importance of magmatic intrusions in terranes?
- What are the crucial factors for ore deposit formation?

Sample	Type	Tfm	System	Salinity (wt% NaCl equiv)	Th vap (degrees C)	Th hal
				(wt% NaCl equiv)		
			(H ₂ O-NaCl +)	*[NaCl +CaCl ₂]		
Slab SW	L+V+S	-61 to -54	+CaCl ₂	~40*	167-230	202-261
Hoover	L+V+S	-55 to -49	+CaCl ₂	~50?*	152-200	???
Igor	L+V+S	-50 to -47	+CaCl ₂	~30*	150-174	175-184
Igor	L+V	-56 to -38	+CaCl ₂		70-136	
Igor	L+V	-32 to -21	+K/Fe?	3-23	85-138	

Table 1: Summary of thermometric data on fluid inclusions in Wernecke samples

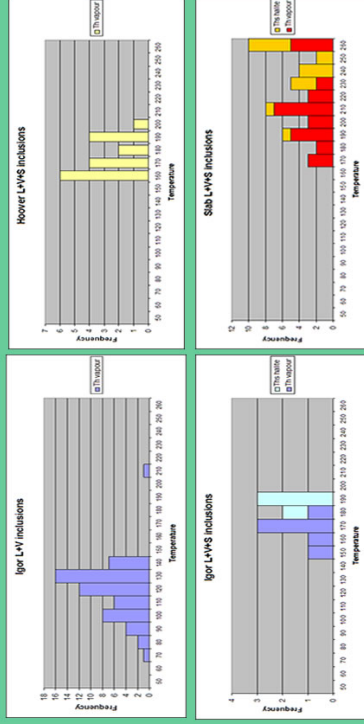


Figure 7: Histograms displaying the homogenisation temperatures observed in fluid inclusions from the Igor, Slab and Hoover prospects. The L+V inclusions homogenised by vapour disappearance, while the L+V+S inclusions homogenised via halite dissolution.

Summary of Data

- 2 main fluid types
 - H₂O-NaCl-K/Fe?
 - H₂O-NaCl-CaCl₂
- Salinities
 - relatively high for L+V+S inclusions
 - lower for L+V inclusions
- Homogenisation temperatures
 - relatively low compared to usual temp of magmatic fluids
- 2 phase inclusions homogenise by vapour disappearance
- 3 phase homogenise by halite dissolution

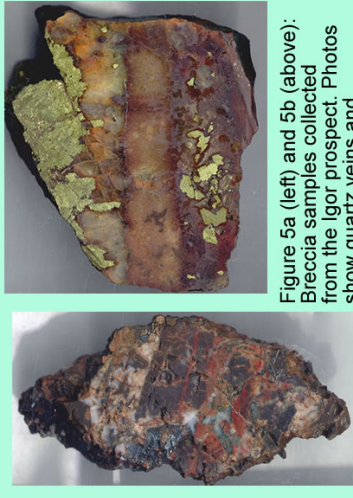


Figure 5a (left) and 5b (above): Breccia samples collected from the Igor prospect. Photos show quartz veins and associated mineralisation. The quartz contains fluids (in inclusions) which may have been associated with this mineralisation



Figure 6a (above) and 6b (right): The two types of fluid inclusions observed in quartz from the Wernecke region. a) 3 phase, L+V+S with halite crystal as the solid phase and b) simple 2 phase L+V.

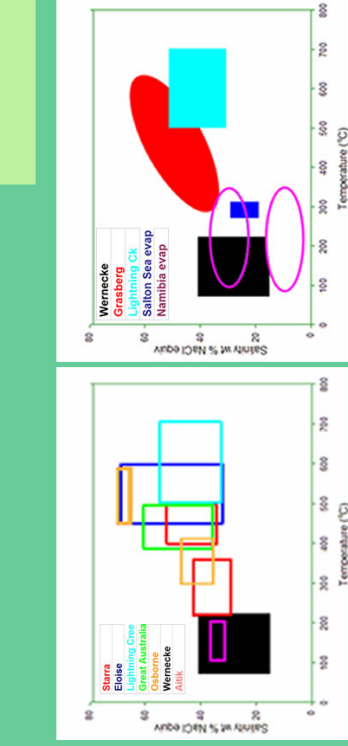


Figure 8a (left): Plots fluid salinity vs temperature for Australian Fe-oxide-Cu-Au deposits. Wernecke and Aitik (Sweden) fluid data has been added for comparison. Fluids from the Australian deposits are thought to be predominantly magmatic, while the Aitik fluids are magmatic with a strong evaporitic influence.

Figure 8b (right): Similar plot comparing Wernecke fluids with endmember magmatic and evaporitic fluids. They seem to lie at the evaporitic end of the spectrum.

Results/conclusions to date

- Wernecke fluids initially seem more consistent with evaporitic type fluids
- But using temperature and salinity alone is very broad and limited
- So further techniques are required to enable better characterisation of the fluids
- In particular- detailed compositional data
 - PIXE, LAICPMS, Laser raman, Bulk/leach techniques

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Predictive Mineral Discovery



Fluid Flow in Extensional Environments: Numerical Modelling with an Application to Giant Microplaty Hematite Ores

¹ John G McLellan, ¹ Nicholas H.S. Oliver, ² Peter M. Schaubs

¹ James Cook University
² CSIRO Exploration & Mining

AIM

To test the effects of extensional deformation and topography on fluid flow and apply numerical modelling outcomes as a predictive tool for Giant Microplaty Hematite Ores.

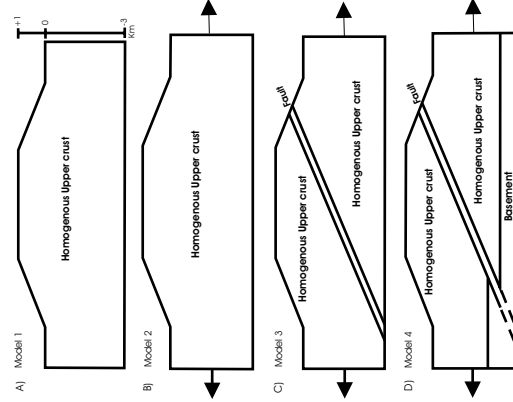


Figure 1. Starting conditions for four basic conceptual models 1) mountain range without deformation, 2) mountain range undergoing extension, 3) normal sense reactivation of former thrust faults, and 4) basement impermeable block and collapse of basement thrust. Note: vertical scale = horizontal scale.

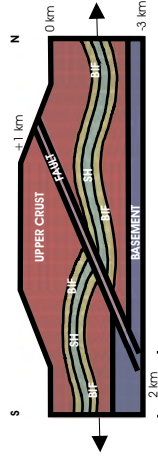


Figure 2. Conceptual Model 5 consists of a 12 km x 3 km two dimensional block of upper crustal material with additional 1 km of topographic relief, which represents simplified topographical and geological elements of an inferred Ophiolite mountain range. Additional structures include a fault, a granite basement block, and folded stratigraphy (BIF- banded iron formation, SH- shale).

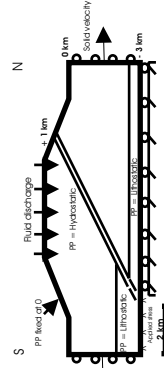
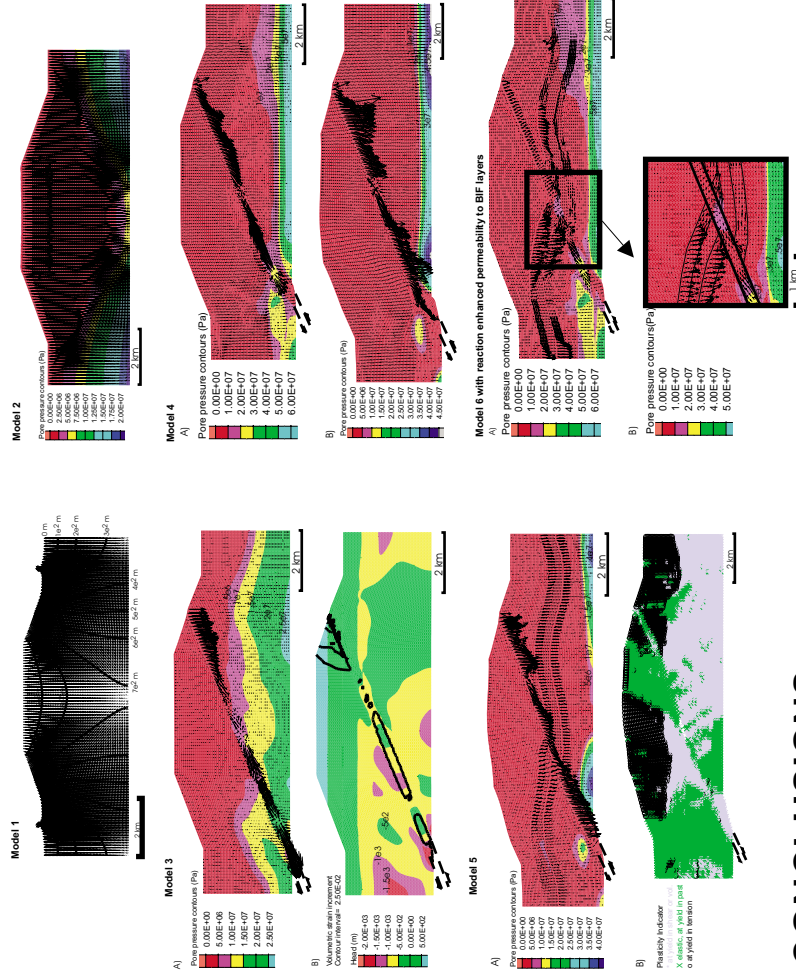


Figure 3. Mechanical boundary conditions for all models. The geometry is progressively altered or contraction but edges remain planar. Initial pore pressure conditions are hydrostatic but allowed to change on all boundaries during deformation. A stress equal to the overburden is applied at the base of the hangingwall, which allows normal sense movement on the fault during extension but no lateral movement. The model is divided into three regions: upper crust, middle crust, and lower crust. The upper crust is shown with a reaction-enhanced permeability layer at the base. Fluid is applied to the top of the model at a rate of 1e-10 ms⁻¹ simulating rainfall.

CONCEPTUAL MODELS

- Model 1 - Mountain range without deformation
- Model 2 - Mountain range undergoing extension
- Model 3 - Mountain range extension with a permeable shear zone
- Model 4 - As Model 3 with a low permeability basement material
- Model 5 - (Fig.2) As Model 4 with folded BIF and Shale layers
- Model 6 - As Model 5 with a reaction enhanced permeability function

Results



CONCLUSIONS

- The presence of topography demonstrates the strong influence of hydraulic head in a homogenous material.
- Extension drives fluid flow down into the centre of the model.
- Addition of a permeable structure allows focussing of fluids within a shear zone and the competition between upward and downward fluid migration is evident.
- Lateral fluid migration can be noted within BIF layering and this is magnified by permeability enhancement during silica dissolution.