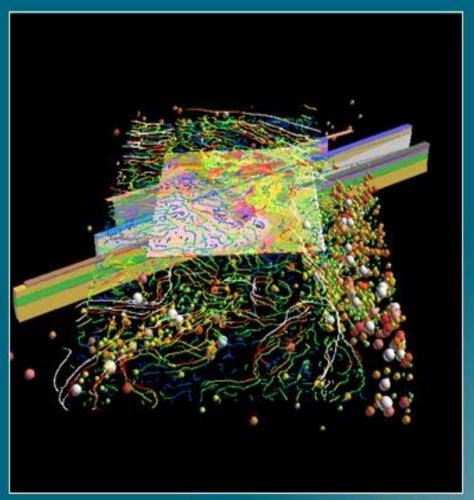
Observable crustal controls on hydrothermal activity in the Olympic Cu-Au province

Nick Direen (University of Adelaide)

Patrick Lyons, Elizabeth Jagodzinski, Peter Milligan, Roger Skirrow, Richard Lane, Dylan Collins (Geoscience Australia)









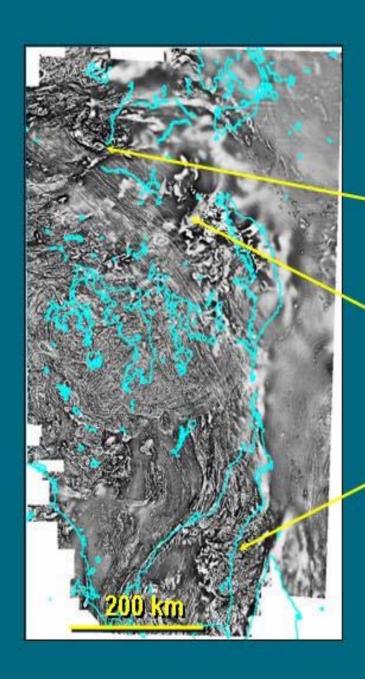


Presentation Outline

- Development of a crustal architectural model for the OCGP
- What was happening at 1590 Ma?
- Implications for fluid pathways
- Implications for fluid buffering







Under cover in the magnetoscape

Prominent Hill

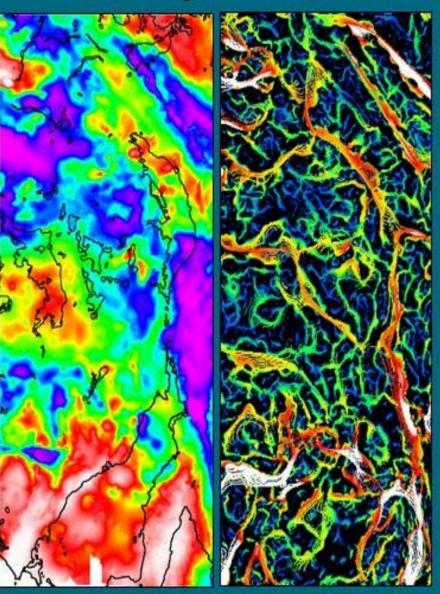
Olympic Dam

Moonta-Wallaroo

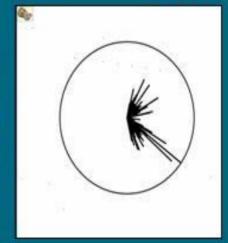
New stitch of residual total magnetic intensity data, 1VD RTP



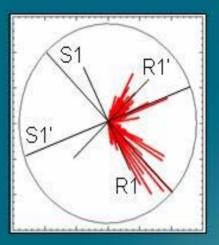
OCGP Gravity Gravity Worms



Structure



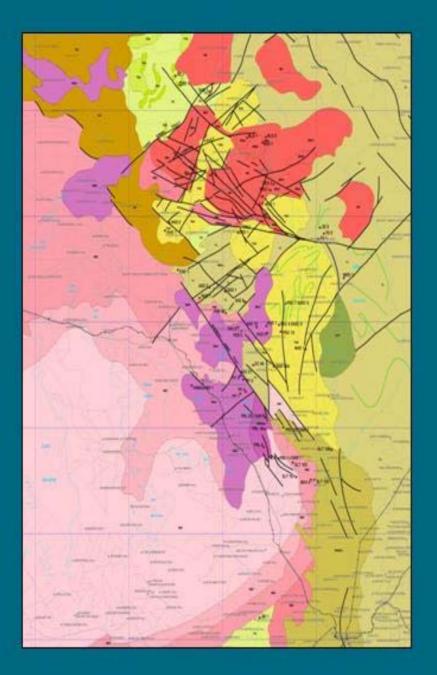
Edge histogram depth slicer 500 m to 41 km UC



Edge azimuth histogram







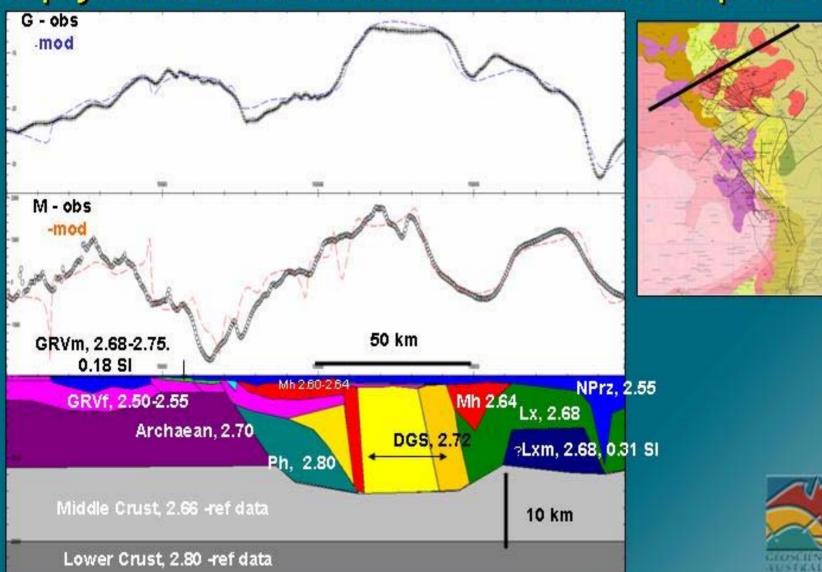
Geophysical Interpretation of the NE Gawler Craton

- Pandurra Fm
- GRV / Roopena Volcanics
- · Hiltaba Suite (incl. mafics)
- · Wallaroo Gp
- Donington Suite
- Hutchison Gp
- Archaean basement



1

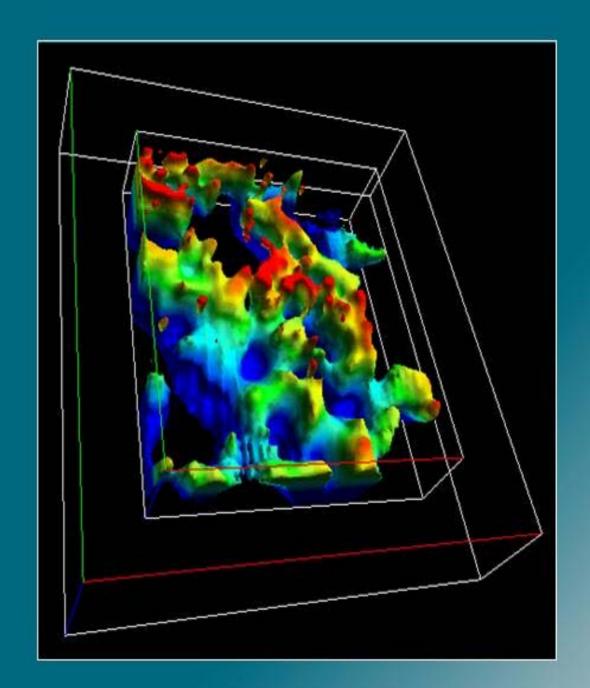
Geophysical models test basement structure and composition





Inversion of magnetic data

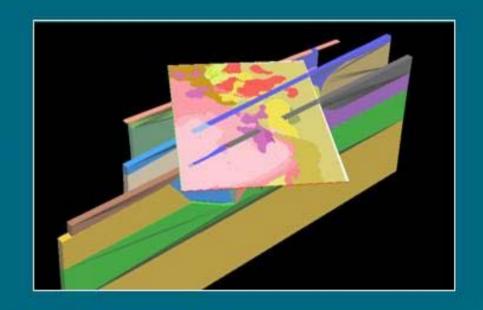
- Susceptibility isosurface
- Uses UBC / MIRA code
- Preliminary
- Builds 3D volume
- Confirmation of structural model





Insights from crustal architecture

 Basement in OCGP forms continuous, eastwardyounging belts with "outcropping" areas to south

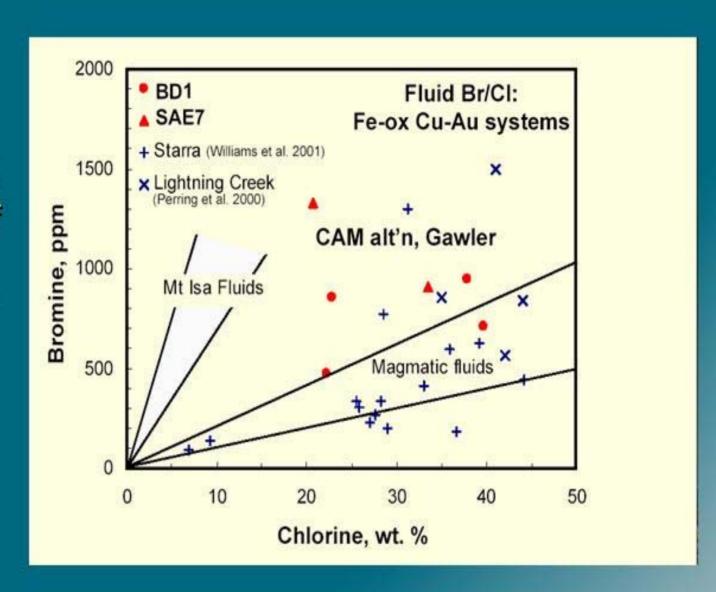


- Structural style typical of regional brittle dextral transtension in the OCGP at ~1590 Ma
- Basin formation processes accommodate GRV + sediment deposition
- Hiltaba suite intrusion appears to be controlled by NW faults opening NNE accommodation spaces
- Basement architecture and evolution is 1st order control on fluid composition, buffering & conduits at 1590 Ma



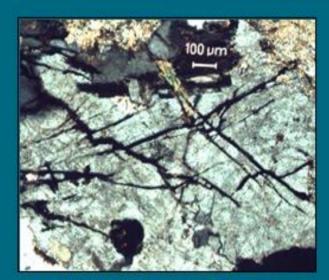
Cu-Au fluids were buffered by a variety of rock-types in the basement

(Bastrakov et al., 2002)

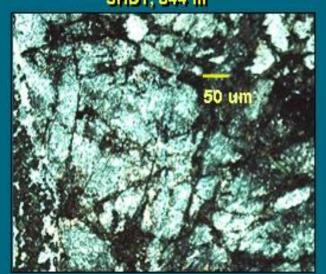


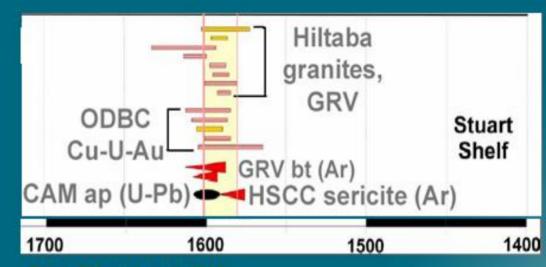


'Tight' rocks of the OCGP were opened to fluid flow, depositing ser-hm ~1590 Ma



Palaeoproterozoic paragneiss, SHD1, 844 m





From Skirrow et al., 2002

Fractured and altered feldspars in BLD1 gabbro, 588 & 615 m



Scale of the Mineral System

 Depending on the efficiency of the system and the availability of metal,
10¹ - 10² km³ of Cu source rock

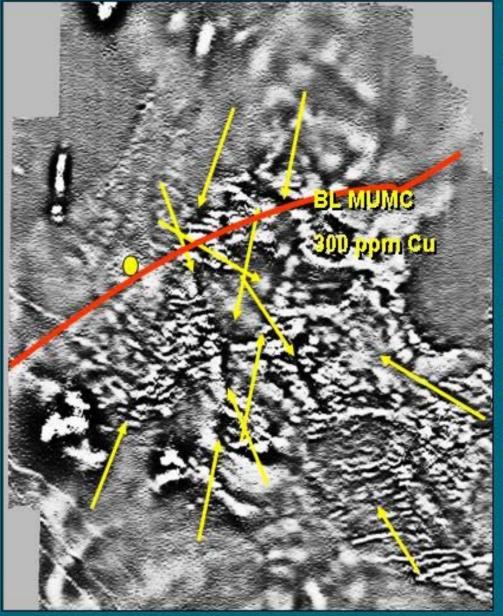
are required to generate 32 Mt in the Olympic Dam fluid system

A fault N km × 5 km will have a fluid reservoir

~ 20N km³ (Cox et al., 2001)





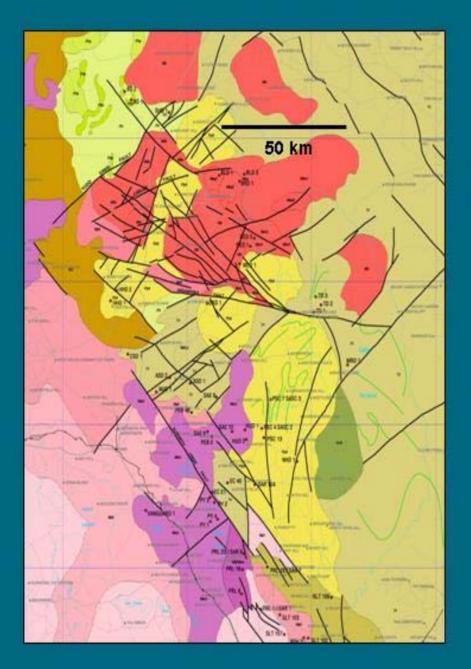


The Olympic Dam region is criss-crossed by a network of faults

These faults are apparent in the mag data -- 1979 data! -- because they carried oxidising fluids

The demagnetisation is very subtle





Implications

•1590 Ma faults of these dimensions tap a variety of potential Cu, Au,U & Ferich source rocks in the OCGP

 The total length dimensions of the 1590 Ma architecture suggests potential for more Cu-Au fluid systems



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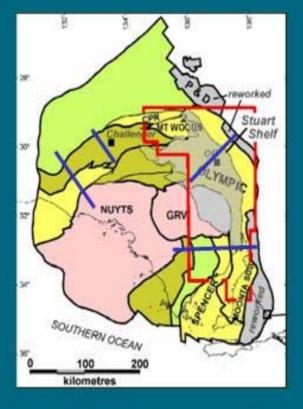
Olympic Cu-Au -

a 1590 Ma structurally controlled fluid system

- Basement architecture continuous with belts to south, forming an apparently outward-stepping concentric system
- This system was put under transtension and high heat flow at 1590 Ma: half-graben, granites, volcanism and faulting
- Basement rocks at 1590 Ma were fertile sources Cu, Au, Fe, U, and were pervaded by altering fluids of mixed compositions
- Faults acted as fluid conduits and fluid drivers

Future Work

MT Geotransect across
 OCGP at Olympic Dam
 and south of Gawler
 Ranges will test
 contiguity of structural
 belts –PhD project with
 PIRSA



- Constrained G & M inversions to refine crustal model, possibly with a view to full fluid FEM modelling
- Geotransects of the Fowler Domain and Challenger area for Ni-PGE and Au system architecture –Hons & PhD projects with CRC-LEME

