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THE CORONATION HILL URANIUM PROSPECT

KATHERINE-DARWIN AREA, NORTHERN TERRITORY.

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THE CORONATION HILL URANIUM PROSPECT KATHERINE-DARWIN

AREA. NORTHERN TERRITORY.

SUMMARY

The uranium prospect at Coronation Hill is situated on the south side of the valley of the South Alligator River, 75 miles from the town of Pine Creek, from which it is accesible by trucks in the dry season. The surface showings consist of discontinuous exposures of autunite over an area approximately 400 feet square upon the north-eastern corner of Coronation Hill, an isolated feature within the river-valley. The rocks present are of Proterozoic age. Mineralization is found within two deeply-weathered members of a sedimentary-volcanic complex which forms the basal member of the Upper Proterozoic: the rocks of the prospect may be in part of Lower Proterozoic age. Sulphide mineralization accompanied by a secondary uranium mineral has been found at one point by diamond drilling: the sulphides are principally pyrite and the nickel sulphide bravoite. Minoralization is associated with a partly crushed and highly altered zone adjacent to a mass of Upper Proterozoic sendstone which overlaps the older and lower rocks. The structure is imperfectly understood and the control over mineralization has not been established. Development work, including costemning, pitting, diamond drilling (1006 feet in two holes) and geophysical surveys (Rediometric, magnetic and electrical) has found no mineralization of economic grade, although it has shown the presence of alteration and sulphide mineralization which indicate that the full potentialities of the prospect have not yet been determined. Further work sufficient to complete preliminary exploration would require a minimum of 2500 feet of diamond drilling on underground operations at least equivalent to 1000 feet of drifting and crosscutting. Diamond drilling is somewhat preferable. since it can reach more deeply below the weathered zone. Only after completion of this additional work can it be determined whether or not the prospect is of economic value.

INTRODUCTION

Location and Access: The uranium prospect at Coronation Hill, Northern Territory, lies approximately in Latitude 13035'S and longitude 132036'E., upon the southern margin of the velley of the South Alligator River (See Plate 1). Access is easy only from the town of Pine Creek, 156 miles south of Darwin on the Stuart Highway. From Pine Creek a truck-road leads to the prospect, 75 miles distant, by way of the homesteads of Goodparla and Cimbat cattle-stations. The road is suitable for general traffic only in the dry season, although trucks of conventional type may travel as far as Gimbat homestead at favourable times during the wet season. Beyond Gimbat homestead, the track is suited only to vehicles with four-wheel drive, and is passable only during the dry season. The prospect may easily be found, since it lies closely adjacent to the river upon an isolated hill: the workings are distant 1000 feet from the river and lie 150 feet above it. At this point, the South Alligator River continues to flow throughout the dry season in quantity sufficient for the purposes of prospecting or small-scale mining operations.
2. Description of the Prospect: The workings of the prospect are situated upon the lower and middle slopes at the north-eastern corner of Coronation Hill, a mesa-like outlier of the plateau which bounds the river-valley on its southern and western sides. The slopes are steep and in great part covered by rubble spalled from the red sandstone bluffs which crown the hill. At the site of the prospect, the foot of the hill is formed by the cliff-like outcrop of a grey and white quartz reef. The reef, and associated showings of secondary copper minorals, is held under Minoral lease 3A by Mr. J. Cellanan, owner of Gimbat Station. The uranium showings are not directly associated with the reef, but lie above it on the middle slopes, in great part beyond the southern boundary of M.L. 3A. The showings are principally exposures of autunite minoralization within a brown-weathering slabby siliceous rock (See Plate 2, chlorite schist) which forms low ridge-outerops, and within an adjacent soft-weathering white sedimentary breccia. Plate 2 shows the outcrops of the mineralized area, and in addition, radiometric "contours" or isorads which define the area of anomalous radioactivity. Plate 2 shows as well the location of the workings, which are individually referred to later in the report. No mining plant and no permanent camp has been established at the prospect: the only buildings constructed consist of two small wooden storage huts and an open kitchen built of screening and galvanized iron. The buildings lie upon a terrace 25 feet above the river, almost due north of the workings of the prospect: from them a rough truck-road leads to the collar-sites of the diamond drill holes. A permanent water-line of 2-inch galvanized pipe leads from the river to the site of D.P.H. No.1 (See Plate 2), but no pumping equipment is installed.

REGIONAL GEOLOGY

The geology of the region surrounding Coronation Hill is the subject of a separate report (Walpole, 1953). Plate 1 shows the major features of the area, and the position of the prospect within it. The rocks present range in age from the Lower to the Upper Proterozoic, with thin remnants of Cretaceous sandstone locally forming hillcappings. In the immediate vicinity of the prospect the upper-most and youngest rocks are thick sandstone of Upper Proterozoic age (Mount Callenan Group, correlable in part with the older-named Buldiva Quartzite), which form extensive plateaus bordering the valley of the South Alligator River. The sandstones are underlain by a volcanic-sedimentary complex, which forms the base of the Upper Proterozoic sequence and unconformably overlies folded sediments of Lower Proterozoic age. Within the river valley, the older rocks are exposed as hill or ridge-outcrops often of large size and relief. The valley-floor, at the lower elevations, is a wide flood-plain with a nearly continuous cover of alluvium, diversified by small outcrops of Mount Callanan sandstone, Lower Proterozoic schist, and rhyolite flows and tuffs, together forming a complex pattern of inliers and outliers. Structural details are largely hidden by the cover of alluvium; and it is debatable whether the structure is a mosaic of fault-blocks or whether the heterogeneity of outcrop is due to partial re-exposure of a rugged topography developed upon folded Lower Proterozoic meta-sediments, later partially buried by lenticular rhyolitie flows and pyroclastics, and finally completely buried before peneplanation beneath the thick sandstone of the Upper Proterosoic: the two views are not incompatible, since both structures may be present. Outside the immediate area of the prospect, the Lower Proterozoic rocks are intruded by granite and diorite. The prospect lies adjacent to a zone of regional faulting (South Alligator Fault, See Plate 1) which trends with the regional strike of the Lower Proterozoic: the fault-line is marked by a complex belt of silica-replaced and injected rocks. Siliceous rocks belonging to the belt form a great part of the western edge of Coronation Hill, and of its north-western and south-western corners. The mineral showings of the prospect lie within the sedimentary-volcanic complex interposed between the Lower Proterozoic folded slates and the Upper Proterozoic sandstone, at a point where the volcanic complex is thin and where the irregular surface of the oldest group is close to and partly interlocked with the irregular underface of the andstone. The uranium mineralization is considered to be of late Upper Proterozoic or post-Upper Proterozoic age, since it apparently post-dates the folding and faulting of the Upper Proterozoic rocks of the Mount Callanan Group.

The occurrence cannot be related directly to any intrusive igneous rock; a genetic relation to the rhyolite extrusives is possible, but has not been demonstrated.

LOCAL GEOLOGY.

Rock-Types and Stratigraphy: The local sequence within the area of the prospect is normal and conforms to the regional dip:it comprises a lower series of altered shales and siltstones, in part silicified, an intermediate member of lenticular habit composed of angular fragments of slate, schist, sandstone, rhyolite and volcanic detritus, and an upper member of thick cross-bedded sandstone. Essentially the structure is that of a steepened southwest-dipping monocline, with the sandstone member dipping more flatly than the remainder. The surface exposure of the sandstone overlaps and transgresses the lower members as if separated from them by an irregular and tilted plane of unconformity: an unconformable relation between upper and lower members is further indicated by the fact that the average strike of the bedding of the sandstone is N60°W (300°) in contrast to an average strike of $N35^{o}W$ (325°) within the lower shales, and by the presence of the intermediate conglomerate member, differing in origin and composition from both sandstone and shale. Local alteration is intense: it is therefore difficult to separate age-groups on the criteria of regional metamorphism, and the prevalence of overburden makes lateral correlation impossible: this must qualify the writer's opinion that the lower series represents an inlier of Lower Proterozoic sediments within volcanics and sandstone of the Upper Proterozoic. The irregular interrelation of the three members makes for a lenticular and discontinuous distribution. Much of the information on the stratigraphy and rock types has been gathered from the study of diamond drill core, which provides an uninterrupted sequence free from the grosser effects of weathering: descriptive logs of the two holes drilled are given as Appendices I and II. So far as is known at present, only the silicified uppermost members of the lower series and the angular conglomerate are important host-rocks to uranium, and they alone will be described at length.

The rocks exposed lowest upon the north and north-east flanks of the hill are altered shales and siltstones, dipping south-west-ward and striking N350W. The upper exposures are silicified and form the rock described immediately below. At a small distance eastward and south-eastward from the hill, the shales are observed to pass beneath a rhyolite flow. To the north and north-east of the hill, the shales are overlain by sandstone or by alluvium.

The chlorite schist marked upon Plate 2 is a complex metasomatic differentiate from the shale and siltstone of the lower series: its attitude and persistence are therefore no predictable as sedimentary structures. The rock is predominantly silica with varying but subordinate quantities of chlorite and clay minerals and with a minor but universal con ent of iron. Remnants of bedding are preserved in the less siliceous portions: the more completely silicified rock is massive and quartzite except for a mottling with segregated iron oxide/silicate (See Appendix II) It is possible that the more siliceous representatives include portions of the base of the upper sandstone, since field evidence has not been decisive in fixing the age of the silicification relative to the Upper Proterozoic. The rock as seen upon surface is brown-coated with limonite, and shows a course irregular foliation of slabby structure: locally breccia structure is visible, with angular fragments weathered in slight relief. Autunite is locally present in the form of thin flakes or films upon the surface of planes of bedding, foliation or jointing, usually in company with limonite. In its present form, the autunite post-dates the period of silicification.

The contact between the chlorite schist and the angular

or iron

conglomerate, as seen upon surface, stands vertically and is faulted. The fault strikes N350W, and is apparently of minor intensity. Within the angular conglomerate, a vertical jointing parallel to the fault is observable: the fault, and the associated jointing, have locally but not generally concentrated autunite.

The angular conglomerate is scarcely recognizable in outcrop since it tends to weather recessively, and to loosen to a talus-like mass closely resembling rubble of the current erosion cycle. In the weathered state, it is conspicuous by its whiteness and kaolinization. Unweathered material is known only from the core of Diamond Frill Hole No. 1 (See Appendix I). The lower portion of the conglomerate contains slabs and angular fragments of soft, possibly pre-weathered grey or black slates and phyllites, together with sub-angular to sub-rounded fragments of grey micaceous sandstone and of red sandstone, set in a friable grey siliceous matrix which weathers or alters readily to kaolin or sericite. Rarely, well-rounded quartzite pebbles occur. The rock absorbs water fairly readily: core from below the water-table cracks and exfoliates slightly after exposure upon surface. In higher portions of the conglomerate, blocks or slabs of vesicular rhyolite appear accompanied by a grey siliceous mass suggesting sltered pumice. The upper portions of the conglomerate become increasingly arkosic and eventually difficult to distinguish from the brecciated portion of the overlying sandstone. The rock is probably composed of the weathered rubble and pebbles of an old hand-surface, heterogeneously mixed with detritus from the softer products of vulcanism. Chemical weathering could have affected the rock only slightly before deposition and burial.

The Upper Proterozoic sandstone surrounds the mineralized rocks on the South, west and North sides, limiting their exposure, as if by overlap. The upper portion of the andstone is a dense red ferruginous quartz-sandstone with imp rfect bedding, locally conglomeratic and locally current-bedded, ith a dip constantly to the west of south at angles averaging 350 to 400, and with little observable alteration or mineralization.

The lower portion of the sandstone is known only from diamond drill core: where it occurs upon surface it apparently weathers recessively and becomes covered by the rubble of the upper layers, so that it cannot be determined whether the diamond drill intersections are typical of the base of the sandstone, or whether local agencies have been respondible for its local condition. The lowest known portion of the sandstone, which forms the northern margin and is tentatively considered its base, consists of interbedded sandstone and shale, intensely altered and brecciated, with local patches of earthy red hematite and with veinlets of specularite. It is improbable that the alteration can be due to meteoric water, since it occurs well below the water-table. The brecciation is not accompanied by observable retation or stringing-out (mylonitization), and the altered shale shows few shear-structures. The dip of the lower margin of the sandstone, as fixed by diamond drilling, is steeper than any needing dip observed upon surface. This margin is noted because it forms the southern border of the mineralized zone and to some extent controls its attitude.

2. Alteration and Mineralization: Weathering of the present erosion cycle penetrates to depths which vary with the elevation of the surface. The water tables lies 110 feet below the collar of D.D.H. No.1 situated on the 350-foot topographic contour (See Plate 2). Below the 500-foot contour nest the top of the hill, the depth of weathering probably exceeds 200 feet. Throughout the mineralized zone, the level of the water-table may be taken to be approximately the elevation of the 250-foot

topographic contour. Within the more siliceous rocks, the effects of weathering are not great, and are scarcely noticeable at a depth below 80 feet from the surface: it is to be expected that sulphide minerals would be more severely weathered than the rock but that they would leave clear traces of their former presence within it. The effects of weathering upon the angular conglomerate are pronounced and reach fully to the water-table: the rock is soften ed and bleached, and portions of it are rotted to a fine incoherent sand which may hinder diamond drilling. The intensity of the effect indicates an inherent porosity in the rock and suggests that the components, as deposited, were not fully weathered.

The silicification present within the chlorite schist has not been completely studied. It is earlier than a period of quartz injection which is visible in the rock as stringers and veinlets roughly following the foliation, but occasionally transgressing the rock-structure. So far as is known the quartz of neither period is genetically connected with the uranium mineralization which is probably later than both. The silica replacement is assigned on general grounds to the period or periods of silicification within the zone of influence of the South Alligator Fault.

Alteration of the angular conglomerate, other than weathering, has taken the form of sericitization and the formation of clay minerals, particularly within the matrix. The alteration is described by Dallwitz (1953) from microscopic study as probably of hydrothermal origin, and this conclusion is supported by the field evidence.

Within the base of the upper sandstone, diamond drilling has found an altered condition at depths too great for the action of ground-water. The alteration has not been studied microscopically. It has resulted in a leached and rotted appearance within both sandstone and shale: the latter is completely altered to serpentine, chlorite and tale. The redistribution of iron originally present is marked by the frequent irregular patches of red hematite; specularite, present in distinct veinlets, appears to be additive to the rock. The nature of the alteration, and the depth at which it is observed, make it probable that the rock has shared in the hydrothermal alteration of the adjacent angular conglomerate.

Secondary mineralization with autumite is invariably in the form of thin flakes or films upon joint or parting-planes, irregularly distributed throughout the mass of the rock in a manner for which no consistent control can be seen. Torbernite occurs infrequently but in the same manner as does autumite. Because of the ease with which these minerals may be transported, their distribution cannot be relied upon to reflect the distribution of the primary minerals from which they may have been derived.

Only one instance of sulphide mineralization is known, from the core of Diamond Frill Hole No. 1. The mineralization occurs within the angular conglomerate, below the zone of weathering (See Appendix I). Pyrite alone is visible megascopically, in small lenses or small rounded aggregates. Mineragraphic examination (Roberts, 1953) discloses the presence additional to pyrite of a nickel sulphide (Bravoite, NiFeS2) frequently associated with marasite, together with small quantities of chalcopyrite, sphalerite and galena. Associated with the sulphides, but independently disseminated throughout the rock in a very fine dispersion, is a uranium mineral, as yet unidentified but presumably of secondary origin. The Association suggests mineralization by hydrothermal agencies, probably recurrent or serial. This is in accordance with the field evidence, and with the microscopic study of the unweathered core (See above.

Deliwitz 1953). The writer considers it probable that the granium mineral present has been derived from an early-deposited primary mineral by paulopost alteration accompanying a later phase of a long mineralizing sequence, and that the pyrite owes its rounded shape to correction by a similar agency. The Mount Callanan sandstone has shown only a few small radioactive patches, which have been found to be of negligible significance. Fluorescence occurs upon joint-planes in the sendstone, apparently due to the presence of wavellite. Thin veinlets composed of hematite and apatite are also present. (Dallwitz, 1953)

Further information on mineralization is given below under Samples and Assays, and in Appendices I and II.

3. Structure: The gross structure of the rocks, and the factors which tend to obscure the details, have been described briefly under Stratigraphy. In general terms, the prospect consists of a hill or hamp of relatively old steeply-dipping slates and siltstones together with their silica-chlorite derivatives, enclosed within an embayment in a thick mass of sandstone, which lies upon the south, west and north of the shales at topographic levels both above and below the shale outcrop. To the eastward the shales decrease in topographic elevation and pass beneath a cover of rhyolite, part of an extensive flow. Interposed between sandstone and altered shale, and occupying the south-western and western portions of the prospect, is a thick lense of angular conglomerate, composed of fragments of sedimentary and volcanic detritus. As previously stated under Stratigraphy, the shales are considered an irregular inlying peak of Lower Proterozoic rocks, the flanks of which are partly covered by volcanic and associated sedimentary volcanic detritus, of extremely variable thickness. The irregular resultant surface is partly exposed through its sandstone cover. The mineralized portions of the prospect, and the rocks enclosing them, dip steeply south-westward at an angle approaching the vertical: the overlying and enclosing sandstone shows bedding dips ranging between 25° and 40°, but its margin dips at an angle approaching 80°. Within the area known from diamond drilling, a crushed and hydrothermally altered zone includes the angular conglomerate and the adjacent margin of the sandstone: within and adjacent to the crushed zone secondary uranium almoralis and procedure. The cause of the crushing, and the linear extent of the zone, have not been determined. The mineralized rocks presumably pass beneath the sandstone at the western side of the prospect, but this point has not been investigated and remains unknown.

The massive sandstone on the south and west of the mineralised rocks is interposed between the prospect and the ridge-outcrop of silicified rocks which marks the South Alligator Fault zone. The relationship of the prospect to the Fault is thereby obscured, although it is probable that the silicification of the older rocks is related to the silicification of the Fault zone. At present, no direct connection is demonstrable between the Fault and the local mineralization.

Sithin the area of the prospect, localized folding is apparently absent: although the general structure approximates to an anticlinal nose or to a hemi-dome, this is considered to be essentially an original structure present at the time of deposition of the sandstone, subsequently modified by a regional tilt towards the south or south-sest and probably by local faulting. The altered condition of the rocks, and the extensive cover, has concealed the effects of minor faults within the area of the prospect: considerable local faulting is probably present, but the only fault which can be described is that noted above under Stratigraphy as marking the contact between angular conglomerate and silica-chlorite schist. This fault appears to be of minor magnitude and its persistence is known for only a short way: Its presence and its possible effect upon mineralization below the weathered some have not been established. With this possible exception, no discovered fault can be stated to have had a critical effect upon mineralization.

Knowledge of the structure is indefinite, as the foregoing description implies. The presence of an irregular original structure, embodied in the roughly unconformable ancient surface and the lenticular character of the intermediate volcanic-sedimentary series, obscures the presence of later structures, which can be delineated only by investigation in detail. The structure of greatest importance is the crush-zone bordering the sandstone and including the mineralized portions of the angular conglomerate, though not observably including the silica-chlorite rocks. This zone is known only at two points, through diamond drilling, and its size, attitude and ultimate cause are unknown. It may safely be said, therefore, that the structure governing the mineralization is yet undetermined.

GEOPHYSICAL INVESTIGATIONS

Geophysical study of the prospect included radiometric, magnetic and electrical (Self-potential) surveys controlled by a closely-spaced grid. Study of the results of the magnetic and self-potential measurements are not complete and are not included in this report. Radiometric surveying involved gridding of the surface exposures in order to establish the distribution of radioactivity, measurement of the intensity of radiation within the pits and costeans, and radiometric logging of the bores, sludges and core of the diamond drill holes. The results are shown graphically on Plates 2, 3, 6 and 7. The measurements upon surface and in the costeans are of relative value only since they afford a comparison of intensity of radiation between different parts of the surface without establishing the quantity of uranium present at any one place. The measurements of radioactivity in the horewalls and core are, however, related to absolute values established by laboratory assays (See Plates 6 and 7).

DEVELOPAGINT FORK

In addition to geological mapping, and to the geophysical investigations described above, the radiometric anomalies of the surface were tested by seven costeans and several pits, and by two diamond drill holes. The distribution of the costeans appears upon Plate 2: in general they are shallow in depth and are not sunken far into the bedrock, since their purpose was to remove the rubble cover of the hillside and to establish the nature and condition of the bedrock at selected points. Of the seven costeans shown upon Plate 2, only Costeans Nos. 2 to 5 inclusive were able to penetrate to bedrock through the rubble: they were all sited within the outcrop area of the angular conglomerate and within the most uniform and consistent portion of the surface radiometric anomaly. All four exposed autumite in place within the bedrock: in addition, showings of torbernite were present in Nos. 2 and 3. Depth-profiles of the costeans appear upon Plate 3. A vertical pit was subsequently sunk in the floor of No.5 Costean to a depth of 14 feet below the surface, exposing weathered bedrock within which flakes of autumite were sporadically visible. Of several smaller pits, only two (A and B, Plate 2) were in bedrock: both exposed small showings of autumite.

which lies approximately 150 feet below the surface at No.5 Costean. Diamond drilling was relied upon for deep exploration, and for sampling. Two diamond drill holes, each 503 feet in length and inclined at 35° below the horizontal, were drilled on parallel bearings from collar-sites 150 feet apart (See Plates 2, 4 and 5). The holes pass across and below the principal anomaly: because of the limited choice of drilling-sites, the holes were drilled in the direction of the structural dip, though at a flatter inclination, rather than in the preferred direction normal to the dip. Because of the prevailing steepness, reasonably satisfactory transection of the rock-sequence was obtained, in succession from lower beds to higher. Diamond Brill Hole No.1 tests below the pit in No. 5 Costean: Diamond Brill Hole No.2 tests the zone to the south-east of No.1. Geological logs of the holes are appended (See Appendices I and II); and vertical sections are shown on Plates 4 and 5. The results of the drilling is further discussed under Geology and under Sampling and Assays. No great difficulties were encountered in drilling, which progressed at an average rate of one foot per working-hour. Core recovery in D.D.H. No.1 averaged 80%: D.D.H.

No.2 recovered 95% of core in more siliceous rocks. Water and sludge recovery was excellent from both holes.

SANTATIO AND ASSAYS

With the exception of a series of channel-esuples out within the deep pit in No.5 Costean, no systematic sampling was undertaken upon the surface exposures: a measure of the relative distribution of radioactive minerals, given by close-spaced radiometric surveys, was considered sufficient to guide exploration. The relative distribution of mineralisation within the costeans is known only from the radiometric profiles measured along the length of each costean: these profiles appear upon Flato 3.

The percentages of wranium oxide quoted in this report are based upon radiometric assays, and are stated as the radiometric equivalent of the standard oxide, UjOg (abbreviated e UjOg). The presence of uranium has been confirmed chemically in a few cases, and the visible presence of autumite and torbernite has been accepted as confirmation that at least the greater part of the radioactivity is produced by granium.

Channel sampling of the deep pit tested only weathered material, and the results are impossible to assess at present as an indication of conditions likely to be found below the weathered sone. Twelve samples average 0.129% a U;08 in grade: the values are creatically distributed and the average is made of ten samples averaging 0.033% and two samples averaging 0.607% a U;08.

The greater number of systematic semples, and the only eachles including tests of unweathered material, were taken from the sludges and cores of the two dismond Crill holes. Sludge samples were taken continuously, each sample representing 5 feet of borelength. Core samples were split at selected lengths to afford checks upon the sludge assays. The walls of the bore were scanned by a geiger-maller probe, and a continuous log of the radioactivity present was measured and plotted graphically. The process was repeated with the core, which was scanned throughout its length by a ratemetered g.m. counter. The results of the various assays and measurements were plotted graphically and appear upon Plates 6 and 7. Lists of assays of dismond drill hele samples are given in Appendices III and IV: particulars of the mineralization are given in the descriptive logs of the core (Appendix I and II).

Dismond Brill Hole No. 1 found traces of sutunite in weathered material beneath the deep pit, 75 to 100 feet deeper within the zone of weathering. Sludge and core essays over a length of 35 feet, however, average little more than 0.010% e Va0g. (Appendix III, 70-105, and Flates 4 and 6). Sulphide mineralization, recovered from below the zone of weathering, averages 0.047% e Ua0g over a bore length of 45 feet in sludge samples (Appendix III, 245-290). A representative section of core 29 6 in length from the same part of the bore averages 0.051% e Ua0g (Appendix III, 246-277).

Diamond Drill Hole No.2 found minerelisation only in the somes of weathering. Two somes were cut, of which the lower and larger averaged O.O.I., e U308 over a bore length of 20 feet. Core samples, over a more accurately determined length of 17 feet, averaged O.O.77% e U308 in the same zone. (Appendix IV, 95-115, and Flates 5 and 7).

The relation between bore-length and the true horisontal width of a mineralized some is not accurately known. The somes appear to dip nearly vertically: if this is generally true, horizontal widths are approximately 80%, of the bore-length. In view of the imperfect knowledge of the structure, no attempt is made to estimate the size of the mineralized lenses, or to limit them together in a single defined probody.

CONCLUSIONS AND DECLES REPAIRIONS

The prospect at Coronation Rill cannot be evaluated unless further work is done upon it. The investigation completed up to

the present has shown an increasing promise without disclosing the presence of an expleitable orebody, or fixing the limits and source of the mineralization. Further work, if undertaken, must have as its object the search for continued improvement at greater doubt or laterally beyond the limits of the known area. It is considered improbable that an orebody exists within the limits marked by the two dismond drill beles, despite the presence of small lenses or pockets of good-grade secondary minerals. The number of such lenses, and their aggregate value, can only be determined by an underground operation or its equivalent in close-spaced deep pitting or drilling: a programme of work of this nature is not recommended at the present time, since it would require concentration of effort within an area of doubtful value which may represent a detatched fringe of a larger mineralized zone. The value of the prospect remains speculatively dependent upon the existence of such a zone, may yet undetermined but adumbrated by the degree of hydrothermal alteration observable within the rocks of the known area.

What is required, therefore, is a continuation of the employatory progresse beyond the limits of the completed work, rather than intensive development within them. As its primary object, the work should aim at finding and testing possible extensions of the sulphide zone encountered in Diemond Drill Hole No.1. at greater depth beneath the present hole and within the area to the north-west of it. An effort should be made to test the unweathered zone beneath the chlorite schist outcrop and beneath the deep pit, for the possible presence of sulphides.

It is considered that five more diamond drill holes, each between 500 feet and 600 feet in length, would suffice to disclose the potentialities of the prospect. It would be most convenient, and probably most effective, to drill these holes parallel in bearing and inclination to the two holes already completed. It would be necessary to collar them as low on the hillside as practicable in order to get below the weathered some as quickly as possible. The point of most immediate interest, as noted above, is the area lying below D.D.H. No.1 and extending some 200 feet to the northwest of it: three diamond drill holes may be required here, with the remaining two holes testing the area below D.D.H. No.2 and to the south-east of it. In all a minimum of 2500 feet of diamond drilling would be necessary, or tix menths work for one diamond drill working two shifts per day.

The work may also be accomplished by mining operations. A shuft, if collared just below the surface showings, would reach below the weathered some at a depth of 100 to 125 feet. Exploration by adit is feasible, but an adit level could not be established conveniently at a greater depth than 90 feet below the collar of D.D.M. Ho.1. The level would therefore be alightly above the water-teble throughout the greater part of its length. The zer crosscut leg of the adit would be approximately 500 feet long if carried through to the sandstone. Drifting length would very from a minimum of 500 feet to a maximum of 900 feet, depending on the requirements of an initial test.

of the two methods, the underground test would be the more certain within the scope of the workings, but it would not reach as deeply as the diamond drill: it has been noted that an adit level would scarcely test the unweathered zone at all, unless winzing were resorted to. It would seem profemable to continue exploration with the diamond drill, rather than to sink a deep shaft or develop an adit level which may be too shallow to be effective. Because the closely-adjacent private holdings upon AL JA form the only convenient base for further operations, future exploration and development can be carried out most expeditiously by a non-governmental organization capable of entering into negotiated agreements without the restraints necessarily imposed upon the public service.

In conclusion, it may be said that while the work so far completed at Coronation Hill has not proved the prospect to be of economic calue, certain features are present which indicate the possibility of improved conditions in adjacent parts of the mineralized zono. The ferourable features are the widespread

occurrence of secondary arenium minerals and the relative concentration of intense hydrothermal alteration, coupled with the occurrence of sulphide mineralization of a varied and complex type. Their presence must qualify the otherwise unfavourable conclusions suggested by the low grade and by the lack of any easily identifiable controlling structure.

APPENDIX I

CORONATION HILL, NORTHERN TERRITORY.

DIAMOND DRILL HOLE NO. 1.

LOCATION: COORDINATES 02+40S 01+65W

ELFVATION OF COLLAR:

BEARING: 850 degrees T.

DEPRESSION : 35 degrees.

LENGTH: 508'4" BEOUN 15 SEPT. 1953. FINISHED 20 CCT. 1953.

CORR STING : NXO-10, NM 10-S1, DX 81-244, AX 244-508.

0 -10: Casing (westbered chloritic Schist, as below)

10'0" - 48'3"

Siliceous chloritie schiet: approximate composition silica 75%, chlorite 20%, iron uxides 5%. Weathered, somewhat limonitie. Prevalently brecciated with angular fragments: Locally shows faint bedding crossing the core at 30 degrees. Many thin quarts stringers:

Pronounced weathering is localized 36'0" to 40'0" and 45'0" to 48'0".

No radioactivity was detected in the core, sludge or bore-walls.

The contact at 48'3" is steeply across the core and its vicinity is marked by broken core, closely jointed or slightly sheared. Fault material was not noted, but the contact is probably faulted.

48'5" - 174'0"

Angular conglomerate.

Groundmass composed of grey-green porous sandstone (porosity judged by gain in weight after immersion in water): Abundant sub-angular fragments of grey micaceous sandstone and red sandstone. Sparse well rounded quartaite pebbles. Abundant sharply angular fragments of black slate or phyllite to 105'0"; thereafter slate fragments sparse. Abundant fragments of white or grey slate to 75'0" only. Core weathered, kaolinitic and limonitic to 104'0", weathering particularly marked 48'8" to 80'0" and 97'0" to 104'0". Below 104'0" the core appears fresh and solid. Traces of limonite persist to 134'0" : The mater table probably lies et 140 feet. Redioactivity detectable in core, sludge and borewalls between 58'C" and 105'O". Generally of low order, 12 times background. Autumite fluorescence (specks) noted at 62'5", 69'9" intermittently to 71'8", 98'5" intermittently to 100'6", and 103'0" to 103'8". The autumite could not be detected without ultra-violet light, and could not be found below 105'8". The mineralized zone, 69'-104' is vertically below the deep pit in No. 5 costean, and corresponds in appearance to the material exposed in the pit. Core losses were marked between 80'0" and 95'0", where only 5'3" of core was recovered. Considerable loose sand was encountered at this point, nearly directly below the deep pit. No evidence of sulphide mineralization was noted. Coorse limonitic rugs (after carbonate?) noted 80'-911.

174'0"-289'0"

Interbedded grey porous sandstone (as in the groundmass of the foregoing section) with gritty chlorite schist (altered shele?) The rock is brecciated, and is prevailingly rotten and crumbly. Red earthy hematite noted 221'0" to 239'0". The core, though from below the voter table, appears leached and weathered, in port due to water wesh from the drill. Shele predominates 179'0" - 183'6" and 196'0"-807'0".

A single isolated rise in radioactive count was noted at 216'0". The effect is slight and no reason could be seen.

239'0" - 277'8" Anguler conglomerate: Mearly identical with the conglomerate 48'0" - 174'0" but greater proportion of red sandatone, and groundmass more nearly normal sandstone or arkose: Core has prevalent reddish colour. Black slate fragments present, not abundantly. No white or gray slate. Green chloritic shale fragments present, not abundant, Fragments prevalently small, up to 1" maximum dimension. Irregular soft flaky green streaks and patches (sericitic?) frequent as if alteration products.

> The core shows little evidence of jointing or shearing. It is fairly soft and cracks slightly on exposure to sun and sir.

Small subedrel masses of pyrite are found sporedically, in aggregate mass less than 5% of the total core, throughout the length 245'-278'. The pyrite tends to occur in the block slate fragments. as if replacing them. Radioactivity is noted in close association with the pyrite. Ratemeter reading up to 3 times background noted 250'-270'.

277'8" - 287'9" Dark red amygdeleidel volcanic (rhyolite) in section up to 24" in length, with grey porous sandstone interstitial. It appears that these are blocks or boulders in the conglomerate, but the evidence is not conclusive. The care is broken in short pieces. and has a slightly water-weathered appearance.

> Traces of redioactivity exist at 287'. No reason could be seen.

287'9" - 305'0" Angular conglomerates: As 239'0" - 277'8". Very few small fragments of bleet shale could be noted, accompanied by sparse pyrite mineralization. Core soft, somewhat kaolinitic, with patches and streaks of red earthy hematite: Few bresks in core. No definite boundary between this and following rock-type.

> Radioactivity was noted throughout, of low order (12 times background). The source appears associated with the pyrite.

- 305'0" 835'0" Breccieted porous sandstone, or breccieted fine conglomerate. Much patchy and streaky red bematite. The core has an eltered appearance, and is broken in short lengths. There are no vugs. This horizon probably correlates with No. 2 D.D.H. at 249'0" -282'0".
- 335'0" 367'0" Green chloritic shale, variegated with streaks of yellow and clive tale or serpentine. The core is soft and broken into short lengths. Brecciation has occurred after formation of the serpentine. Some short sections of breceiated sandstone stained with red earthy hematite are present.
- 367'0" 436'0" Sandstone with partings of green chloritic shale and thin shale beds. Breceisted throughout - Sandstone and shale are mottled with red earthy hematite,

Thin veinlets of specular hematite cut the shale beds frequently but are rare in the sandstone.

Brecciation is intense 367 - 290, with much red hematite.

A trace of radioactivity was noted 370 - 375.

436'0"-505'4"

Sandstone, reddish-brown with sparse fine bedding and a few green shale partings which cross the core at an angle of 45 degrees. Pebbles, sometimes angular, sparsely present but no defined conglowerate beds. This is the first intersection of sendstone typical of the Buldiva (Mt. Cellenan) outcrop. Pissuring and slight breediation with red-brown hematite cement are noted throughout.

D.D.H. No. 1 ended at 503'4".

(E.B. Allen)

AT MORE II

CORONALION HILL. MORTHURE THER I'VAY

DIAMORD DRILL ROLL NO. 2

LOCATION : COORDINATES 03-608, CO- 908

MANYATION OF COLLAR

BEARING : 224 degrees T

DEFENDER TON 35 despess

LENGTH : 503 3"

BROUN : 19th Sept. 1953. Finished Lith October, 1953.

CORE SIZES : NX 0-121'8". BX 121'-8" TO 503'3".

0 - 162'6" Siliceous chloritic schiet:

As for the chloritic schist of D.D.H. No.1. but with less brecciation or fragmental structure and fewer quarts stringer-veinlets. Generally green in colour, weathering brown and patineted or mottled. Principal constituents silica (75%), chlorite (20%), iron oxides (5%). Locally shows bedding, when it resembles a cherty siltatone. The texture is often finely areneceous. Possibly a reworked and water-sorted tuff probably a hydrothermally silicified argillitic schist.

O - 80: Core deeply weathered but not crumbly. Much limonite in joint-cracks.

So - 162'6": Weathering largely confined to limonite deposition in joints, and formation of irregular vugs lined with red hematite, probably by leaching of carbonate. Core is frequently dark red due to hematite staining; limonite is not observed below 146'. Hematite most notable 126-130, 141-145

Sharply defined seem of clay gonge with angular fragments at 121, with broken core 120-122: probably fault. Radioactivity was detected in the core 65'-72' and 96'-112' which checked closely with the bore probing. Autunite was noted in thin films on joint planes or bedding planes at 19'6" and again intermittently between 64'6" and 71'6", again between 66'5" and 97'10", at 101'3", at 106'3", between 110'6" and 111'10", and at 145'3". The latter is near the water table (approx. at 150') and was the deepest point at which autunite could be detected. The autunite is transported and not residual. At 98'1" a vag was noted, suggesting a weathered iron minoral. Autunite and possibly torbernite is present in the vag. Very thin quarts veinlets with red crystalline hematite occur sparsely between 100' and 106'.

The core was split for essay in 12" lengths, continuously 64'6" to 72'6", and continuously 89'6" to 112'6". Spot samples were split at 49' and at 145'.

The contact at 162'-6" is bleached in colour; its attitude was indeterminate. No faulting is indicated.

- 162'6"-170'0" Siliceous Metamorphic: Grey, with a rough pumiceous feel. Fine groundmass of brown hematite. Spotted with white quarts aggregates. Solid and quartzitic.
- 170'0"-176'6" Siliceous chloritic schist as before: Golour pale and of altered appearance. Core somewhat broken.

 Trace of radioactivity at 174'.

176'6" - 247'0"

Siliceous metemorphic: Grey to brown. Feels rough. Somewhat granular. Brown hematite generally present as if filling fine pore-spaces. Characterized by presence of finely granular aggregates of quartz, often surrounded by reddish-brown halo of iron orten surrounded by reddish-brown halo of iron oxides. These are derived from a green translucent porphyroblastic mineral, essentially iron and silica, present in grains, laths and rarely as fine seems bordering quarts veinlets. The rock is spotted and flecked with white, pale orange or dark green. Silica (90%) and iron are the only recognisable constituents: The parent rock and manner of origin cannot be identified.

176'6' - 179'0': fine angular breecia, with brown hematite comment.

hematite cament.

179'0" - 200'0": Dense, highly siliceous, very fine in texture: Probably added silice. Resembles a phyolite superficially. 208'0" - 213'0": Red colour pronounced in hematite.

235'0" - 236'0": Bright red, very fine clay gouge possibly marking a fault. Sharply defined.

A trace of radioactivity was noted at 203' No reason could be observed.

247'0" - 249'0" Green chlorite-serpenting, probably altered shale.

Quartzitic sandstone: As 176'6"-247'8" but faintly streaked locally at 40 degrees to the core: Bedding 2L9'0" - 282'0" or alteration. Locally brecciated, with hematite cement. There is no sharp difference from the rock 176'6"-247'0".

282'0" - 286'6" Gritty chlorite-serpentine schist, probably altered shale with rotten sandstone. Irregular streaks of red hematite abundant. No regular structures.

286'6" - 295'6" Brecoiated hematized sendstone.

Gritty chlorite-serpentine schist, contorted. 295'6" - 303'6" impregnated with red hematite. Numerous veinlets specular hematite Probably a sheared, rotted shalesendstone bed.

305'6" - 385'0" Brecolated and altered sandstone. 303'6" - 309'0": Brecoin impregnated with red earthy hematite. 309'0" - 348'0". 309'0" - 348'0". Sandstone, resembling Mt. Callanan: Colour variegated, bleached. Hard, quartzitic. Brecciated throughout with patches and seems red hematite. Patches green shale locally present: Probably disrupted beds. Few vags filled with creas-coloured bladed carbonate, unidentified (Non-

fluorescent).

548'0'-556'0': Rotten sandstone, hematite-impregnated and out by many veinlets of specular hematite. Some crumpled chloritic material present. Bounderies of this zone are generally at 50 degrees to the core.

556'0' - 364'6': Sandstone as 309' - 348' above.

564'6' - 368'6': Rotten sandstone with green chloritic material. impregnated with red hamatite. Many irregular veinlets of specular hometite. Fair Pairly regular structure at 30 degrees to core: bedding or

shearing. 368'6" - 385'0": Brecciated sandstone as 309 - 348, but brecciation more intense. Red hematite prominent to 576'.

Chlorite schist with broken sandstone fragments. 385'0" - 392'0" Schistosity 45 degrees to core. Probably sheared

Brecciated sandstone, quartzitic. Patchy red hematite. 392'0" - L1L'0"

414'0" - 423'0" Green chlorite schist. Solid, structureless except for schistosity.

424'0" - 426'0" Red vuggy breccia: fragments green shale and sandatone. Highly ferruginous, red earthy hematite. Gause of vugs not determinable, probably carbonate leaching.

428'0" - 444'0" Sandstone, quartzitic. Coarsely brecciated, cemented brown hematite. 4" wide fault breccia at 434.

ции'O" - 453'O" Rotten earlistone with chloritic material. Core crusbly, moderately ferruginous with hematite. Traces of radioactivity noted.

453'0" - 503'3" Breclated sandstone: breccia not intense, but fissuring prevalent throughout. Veined and recemented with brown hematite. Locally finely bedded at 40 degrees to core, locally shows pebbles, angular to sub-angular. Prevalent colour reddish brown.

468'0" - 469'0": Rotten, leached and crumbly. Trace radioactivity.

D.D.H. No. 2 ended at 503'3".

M.B. Allen.

AFRENDIX III

CORONACION NILL. N.T.
SLUDGE GARTER ASSAYS (RADIOSEPPIO) DIARGED DELLE ROLE NO.L.

	Sample.		Longth	Sample	45 8W Y	
s		ren From		Length	s v308 Equ iv elent	
					Not marked	
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	6209		ЬÓ	W		
	6209 6213 6220	40	45	#		
	9220	45	50	**	**	
				**	*	
	6222 6223 6224	65	65	#	**	
	6224	65	70	**	***	
				**	**	
		- G		**	0.015 T race	
	6245 6246 6247 6248	36	33	*	0.011	
	6247	90	35	**	0.617	
		. 95		**	Trace	
				**	0.030 Trace	
	6259 6260 6263 6264 6265 6266 6267 6268 6269		116	**		
	6264	113	120	**	**	
	6265	120	125	**	*	•
	5255	125				
			175	**	**	
		130 145 150 153		**	**	
	6270		150	**		
	6271	150	123	#	**	
				*	**	
	6276 6276 6276 6277	166		**		
	6276	īπ	175	**	**	
	62.77	175	180	**		
	62/0	180				
	98 [2	763	120	**	**	
		195		**		
	6282	200	205	**	***	
	6286	205	20	**		
	and the second	220	215			
		<i>3.</i> 7	22U 9 9 K	•	**	٠.
			210	**	**	
	6294	230	235	•		
	629 5	255	240			
	2,09		20			
	22.		13.	**		
		266		**	0.086	
	633	265	270	**	0.063	
7	6314	270	275	***		
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	7.55	565		•		
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	6334			**		
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		710	<i>1</i> 26	**	**	
	24/2	434	TAC	48	**	

Note: "Trace" indicates assays of less then O. MAK a Hano.

Sample No.	Pore La Poe Prom		Sample Length Post	Assay % U308 Squivelent
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9.25		250	**	
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6366	375	38 0	10	**
6367	360	385	**	***
0,368	365		**	
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6.77	116		**	**
6374		Les	**	
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6376	l 12 5	430	鐷	**
	430	435	**	*
	435		**	# #
97/3	##2			**
2780		424	•	**
7.75			**	a
6461	16	166	**	**
634	465		**	**
6385	470	475	**	**
6386	472	482	**	
		i go	10	
6389	150	Ĭ35	**	**
6390	495	500	19	•
6390 6391	500	503	3 ' 0'	10
		les Mysl.	D.D.H.	
A6396 6397	68 6" 70 6"	72'6" 72'6"	2,0"	
A6398	72*6*	711 " 62"	**	0.019
λ56Ω	80° (7'	85°0"	5°0"	Truce.
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5603	9010"	95'O''	**	0.011
550	9510"	97'0"		0.013
5605	97.0	100'0"	3'0" 2'0"	0.017
5606 5607	100'0"	102'0"	5 ****	Crace 0.027
	,			•
5608	248101	250'0"	2'0"	0. O45
5009	<u> </u>	252 0"	•	0.103
	25210"	254 '0" 256 '0"	#	0.025
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5612 5613	258.0.	260'0"	**	0,126 Trace
	260'0"	262101	**	
4615	2691 All	26b*c#	**	0.757
5616	264°0"	266'0"	**	0.073
561.7	206 ' ()"	268*0*	**	0.068
56.4	268'0"	270'0"	H W	0.071
	27010"	27210"	94 19	Trace
5620 5621	272'0" 271'0"	270.10" 276.0"	**	0.013
	276°0"	2778	1'6"	் வுக்

CORCUMBION ALLIA, SAL.

ASSAY RESULTS D.D.W. 20. 2.

ri .		Bore Length Feet		Sengale	Ausky	
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		0			Not sample	
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	6215 6216 6218 6219	7		**	0.021	
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	6219	85	90	***		
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	5 %2 7	_95	100	***	0.043	
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	6242 6249 6250 6251 6252 6253 6253 6256 6261 6262 6261 6262 6263 6289	175 180 185 190 190 205 210 223 230 235 240	215		***	
	6255	215	220	#	#	
	6257	220	225		**	
	6256		230		**	
	9258		422	***		
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Sample No.		b		A esay 4,0 70
	Prom	70	Post	Equivalent
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6 3 CM				
6305	330	335	**	
6306	375	340	**	Trace
2.7	340	242	**	**
		166	**	**
6316	333	360	**	***
6 3 27	350	365	***	3111
	30 2	270		Trace
		212 WA	***	***
6324	386	585	***	***************************************
6325	385	390		***
6326	330	3 95	***	**
		440	49	
i i i i i i i i i i i i i i i i i i i	Los	40.) h1n	**	#
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6332	425	420	•	***
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			75 100	0.0L4 Trace
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6354	165	476	**	**
6355	<u>470</u>	475	**	0.013
6356 6360		140	**	Trace
6361	485	485 490		**
6362	196	165		38
6363	495	4.95 500	**	**
6364	500	907	3'0"	**
A6394	Core San 49'0"	20.0. 20.0.	2. D.D.11. 1'0'	0.025
		1 -	*	
a 403	64.16"	65'6" 66'6"	24	0.07
	65.6.	66°6" 67'6"	***	0.060 Trace
1,05	67'6"	64.60		
407	CAT GH	68'6" 69'6"	**	Truce
408	69161	70'6"	**	
409	70'6"	7.5	**	0.050
410		/ E U	,,,,,	Truce
426	89161	90*6*	•	蝶
427	90'6"	41'6"	**	**
428	91'6"	92'6"	**	
	92'6" 93'6"	53.6° 94.6° 95.6° 96.6°	17	Truce
	Oh T fill	Z16#	***	**
412	OK * 6.9	96°6°	**	0.023
423	96.0	07*6*	**	0.232
144	97'6"	GE 1611	1	0.019
	200,6. 200,6.	100'6"	**	1:race 0.025
	101'6"	109168	樓	0.032
418	102'6"	103'6"	糠	0.013
419	103'6"		•	6. 287
420 421		700,64 102,64	**	0.127
		107.6	**	0.020
		- T		

Saule To-	Bore Length Foot From To	Sample Length Feet	Assay % U yO S B qu iv alent	
423 424 425 46392 6393	107'6" 108'6 108'6" 111'6" 110'6" 111'6" 110'6" 111'6"		0.119 0.075 0.023 0.094 0.065	
6395	145°0" 146°0	1'0"	Treco	
A 453	MM, e. Me, e	2'0"	0.055	













