

# Recognition, structural significance, and prospectivity of early ( $F_1$ ) folds in the Minerie 1:100 000 Sheet area, Eastern Goldfields, Western Australia

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In the Eastern Goldfields, shear zones and folds near major faults have controlled gold mineralisation, so an understanding of their configuration is essential in the search for ore. It is also needed to resolve the adjacent stratigraphy, geological evolution, and prospectivity. This report documents the recognition of large  $F_1$  folds extending over 1000 km<sup>2</sup> in the southern part of the Minerie Sheet area.  $D_1$  faults associated with one of the folds may be prospective for gold, as in the neighbouring Leonora area.

## Regional geology

The Minerie region is located in Western Australia's Yilgarn Craton (Fig. 3), and was mapped by AGSO from 1989 to 1995 as part of the National Geoscience Mapping Accord. It comprises a sequence of Archaean greenstones — metamorphosed felsic volcanic and volcanoclastic rocks, mafic flows and sills, ultramafic rock, and minor sedimentary rocks — intruded by granite\*, rare syenite, and Proterozoic mafic dykes.

## Structure

In the southern part of the Minerie Sheet area, major folds ( $F_1$  and  $F_2$ ) and foliation trends ( $S_1$  and  $S_2$ ; Fig. 4) are products of  $D_1$  and  $D_2$  deformation. Two key areas are critical to the interpretation.

Firstly, in the east, a large north-northeast to northeast-trending open to close upright syncline is cut at right-angles in its hinge zone and west limb by a steep northwest-striking foliation. The foliation is assigned to  $S_2$  because it is a prominent schistosity that parallels the general direction of  $S_2$  foliation throughout the Eastern Goldfields. Hence, the syncline is  $F_1$ . A steep north-northeast-striking foliation, also in the hinge zone of the syncline, parallels the fold's axial plane, and is assigned to  $S_1$ . The recognition of this upright  $F_1$  fold raises the possibility of others being present, and three north-northeast-trending open folds in the south may be such.

Secondly, the outcrop pattern of the

volcanoclastic conglomerate, sandstone, and tuff of the Welcome Well Complex and its lateral felsic volcanic equivalents to the west is strongly suggestive of an isoclinal  $F_1$  fold subsequently folded by the  $F_2$  Benalla Anticline. The strata face north on the northern limb and south on the southern limb. The  $F_1$  fold is therefore a tight to isoclinal anticline, possibly

originally recumbent. Its limbs form a south-plunging arc, and its hinge zone has been disrupted and stoped by dolerite. Hence, the original attitudes of the  $F_1$  axis and axial plane are indeterminable. Symmetrically unfolding the Benalla Anticline about its north-south axial plane causes the  $F_1$  fold to take up an east-west trend; Figure 5 shows a

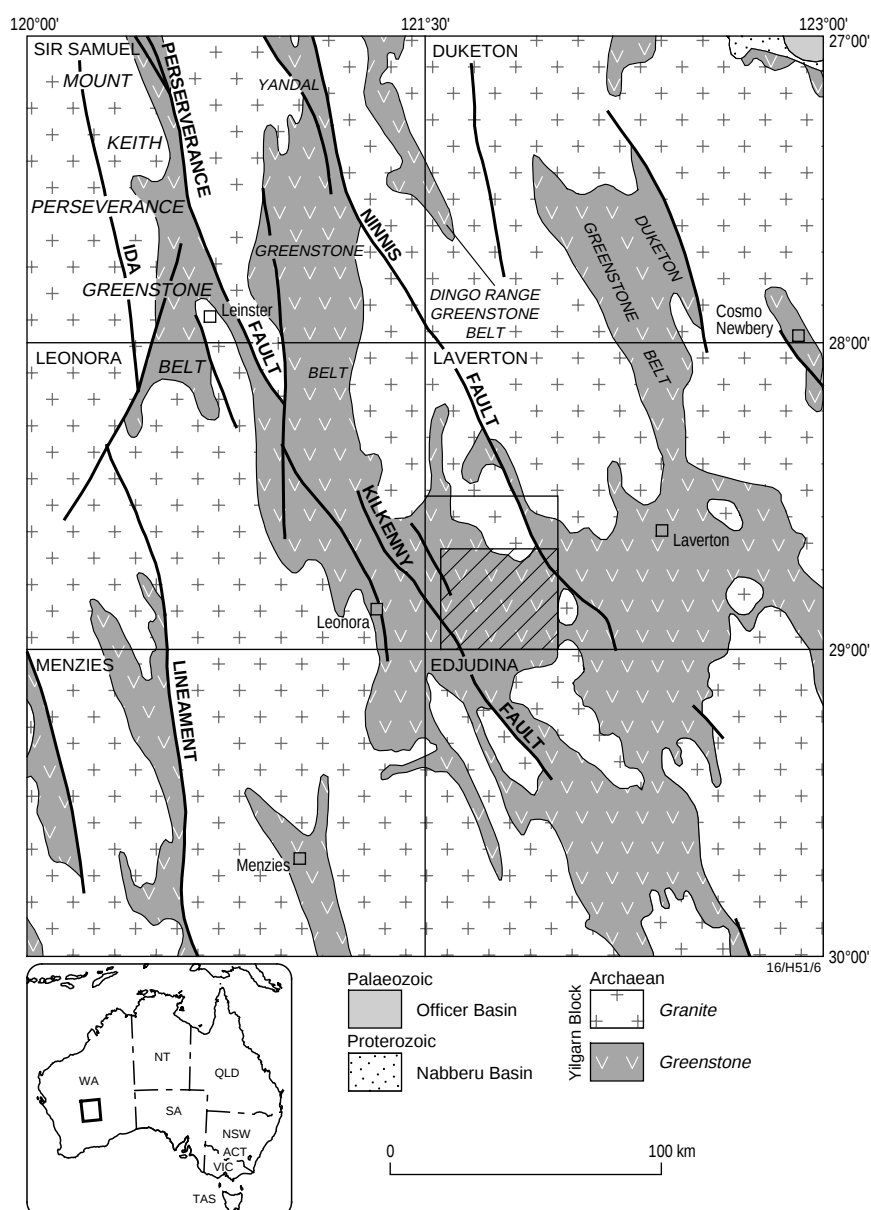


Fig. 3. Regional geology of the central Eastern Goldfields province (after Myers & Hocking, compilers, 1988: Geological map of Western Australia, 1:2 500 000; Geological Survey of Western Australia). The rectangle represents the location of the Minerie 1:100 000 Sheet area; the hatched area, Figure 4.

\* 'Granite' includes syenogranite, monzogranite, granodiorite, and tonalite.

possible sequence of events that led to the present outcrop pattern.

A third foliation, mapped in the country rocks next to the granite body in the north, parallels the granite margin, and may have formed by shear with or without flattening during forceful emplacement.

## Discussion

The north-northeast to northeast trend of the eastern  $F_1$  fold is unusual for the Eastern Goldfields. Large  $D_1$  structures generally trend easterly where they are not reoriented by  $D_2$  (Archibald et al. 1978: *Precambrian Research*, 6, 103–131), or are parallel to  $D_2$  structures where they have been reoriented (Swager & Griffin 1990: *Precambrian Research*, 48, 63–73; other references in Swager 1997: *Precambrian Research*, 83, 18). Witt (1994: Geological Survey of Western Australia, Melita 1:100 000 Explanatory Notes) recognised north-, east-, and northeast-trending large  $F_1$  folds reoriented by  $D_3$  in the northeast of the Melita Sheet area, immediately southwest of Minerie. S. Liu (AGSO, personal communication August 1998) suggests that sinistral movement along the Kilkenny Fault (Fig. 4) could have rotated originally east–west  $F_1$  folds

to their present north-northeast to northeast trend during  $D_2$  and perhaps  $D_3$ .

Chen et al. (1998: Geological Society of Australia, Abstracts 49, 79; in press: Geological Survey of Western Australia, Annual Review 1997–98) and Liu & Chen (1998: *ibid.*, 278) accounted for the north-northeast-trending folds ( $F_1$  of this report) by local compression during east–west  $D_3$  transpression. The folds formed within antidiagonalal jogs defined by sinistral strike-slip on north-northwest-striking faults. The northwest-striking  $S_2$  foliation that cuts the eastern  $F_1$  fold (Fig. 4) contradicts their interpretation.

Two layer-parallel faults on the western limb of the folded  $F_1$  anticline in the west may be  $D_1$  faults. The folded backthrust cutting the core of this fold (Fig. 5) is also  $D_1$ .  $D_1$  shears in the Leonora area carry gold (Williams et al. 1989: *Australian Journal of Earth Sciences*, 36, 383–

403; Williams 1998: AGSO Record 1998/9), and so the interpreted  $D_1$  faults in the Minerie area may be similarly prospective. Several small gold mines are located on or near the eastern layer-parallel fault.

Fieldwork planned in connection with second-edition compilations of the Leonora and Laverton 1:250 000 geological maps will include examination of the various foliations in the Minerie region, their nature and relationships to folds, and shear near the granite contact.

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FIGURES 4 & 5 on opposite page.

Fig. 4. Solid-geology map and cross-section (from mapping by P.R.Williams, M.B.Duggan, and K.L.Currie, 1989–91, and interpretation of aeromagnetic data by the author) of the southern Minerie 1:100 000 Sheet area.

Fig. 5. Possible sequence of events that led to the present outcrop pattern of the Welcome Well Complex and laterally equivalent felsic volcanic units. (a) Isoclinal  $F_1$  anticline. (b) Backthrust in lower limb. (c) Flexure around  $F_2$  Benalla Anticline; dashed lines show subsequent position of strata after dolerite intrusion (which is just as likely to have occurred after stage b and before stage c). (d)  $F_2$  folding of the hinge of the  $F_1$  fold, and intrusion of dolerite.

