

# The Chewings Event: Intracratonic Orogeny in Mesoproterozoic Australia



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# The Chewings Event: Intracratonic Orogeny in Mesoproterozoic Australia



c. 1600-1570 Ma tectonism in Proterozoic Australia

## *Event timing*

Chewings: 1595-1570 Ma

Early Isan: 1600-1580 Ma

NE NAC: 1585-1560 Ma

Olarian: 1610-1590 Ma

Hiltaba: 1595-1575 Ma

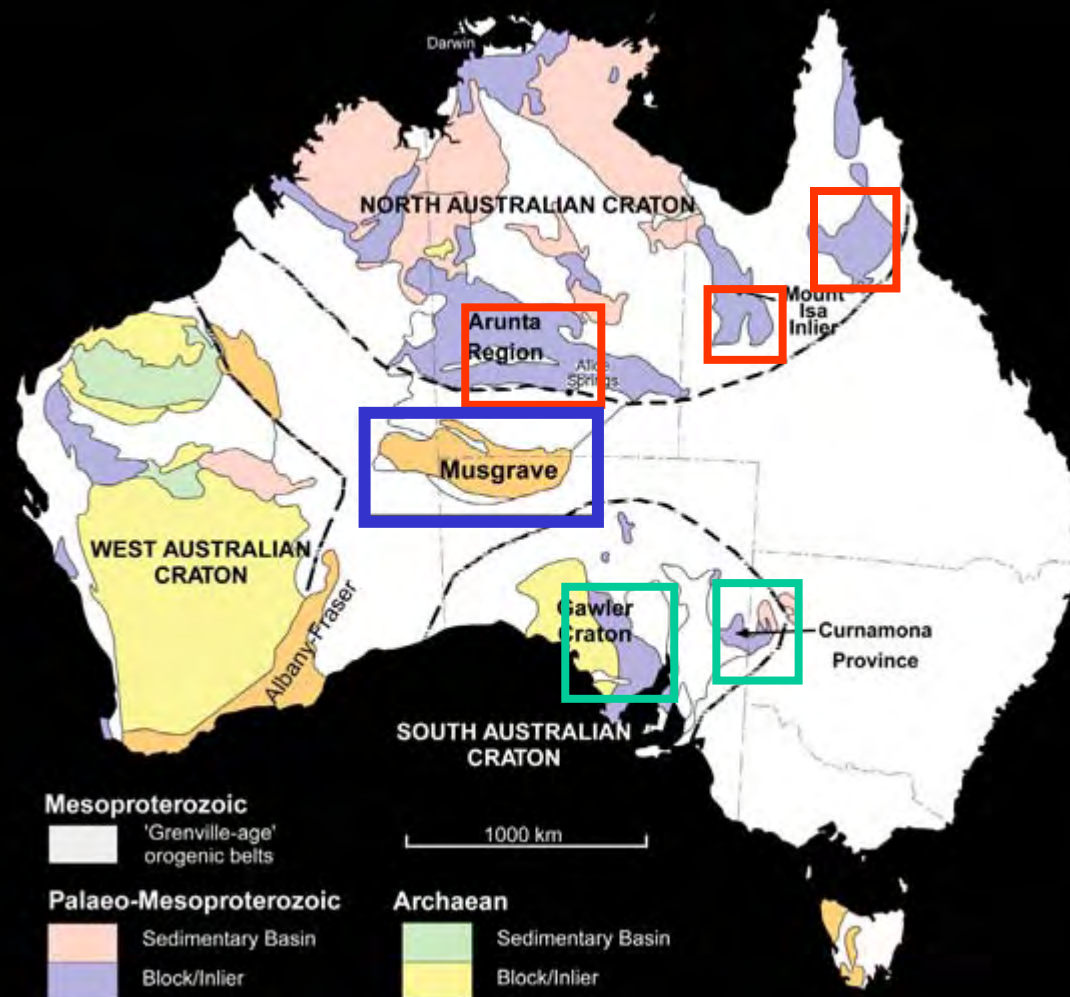
## *Common elements*

Reworked existing crust

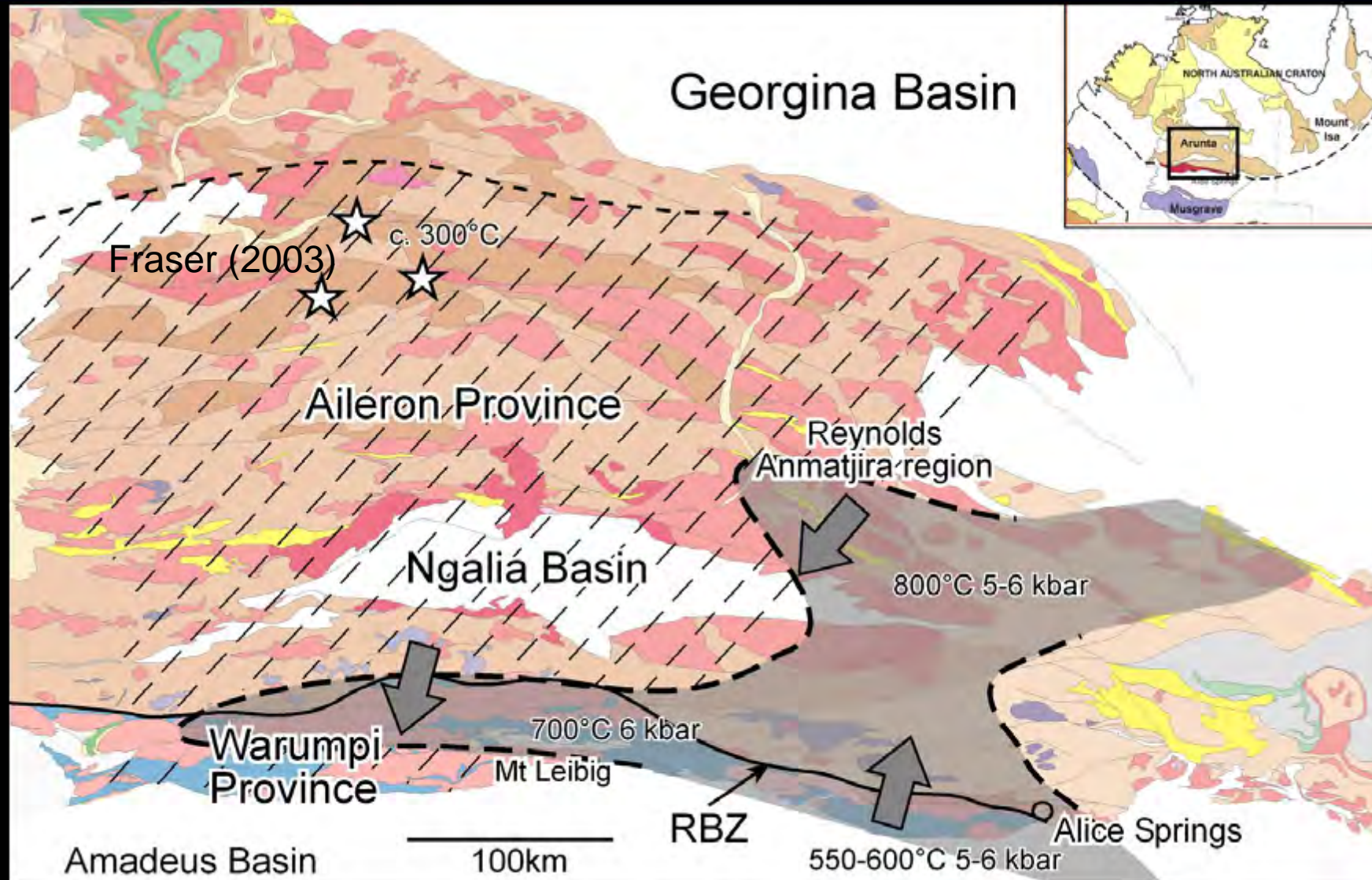
High heat flow: low(ish)-P

Contractional deformation

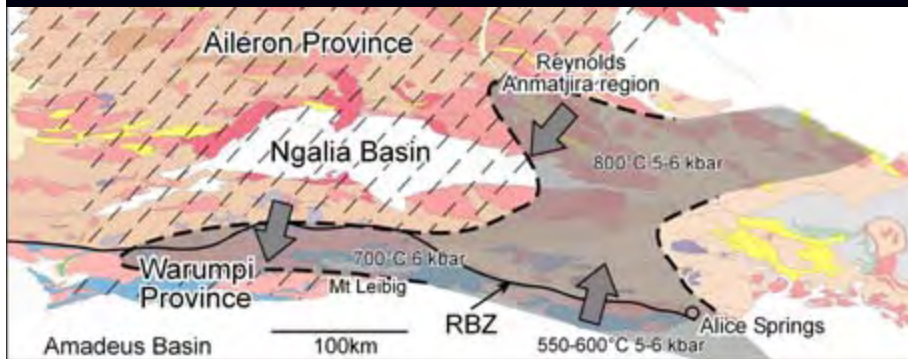
20-30 Ma duration



# Chewings Orogeny: Reworking in the Arunta: 1600-1570 Ma

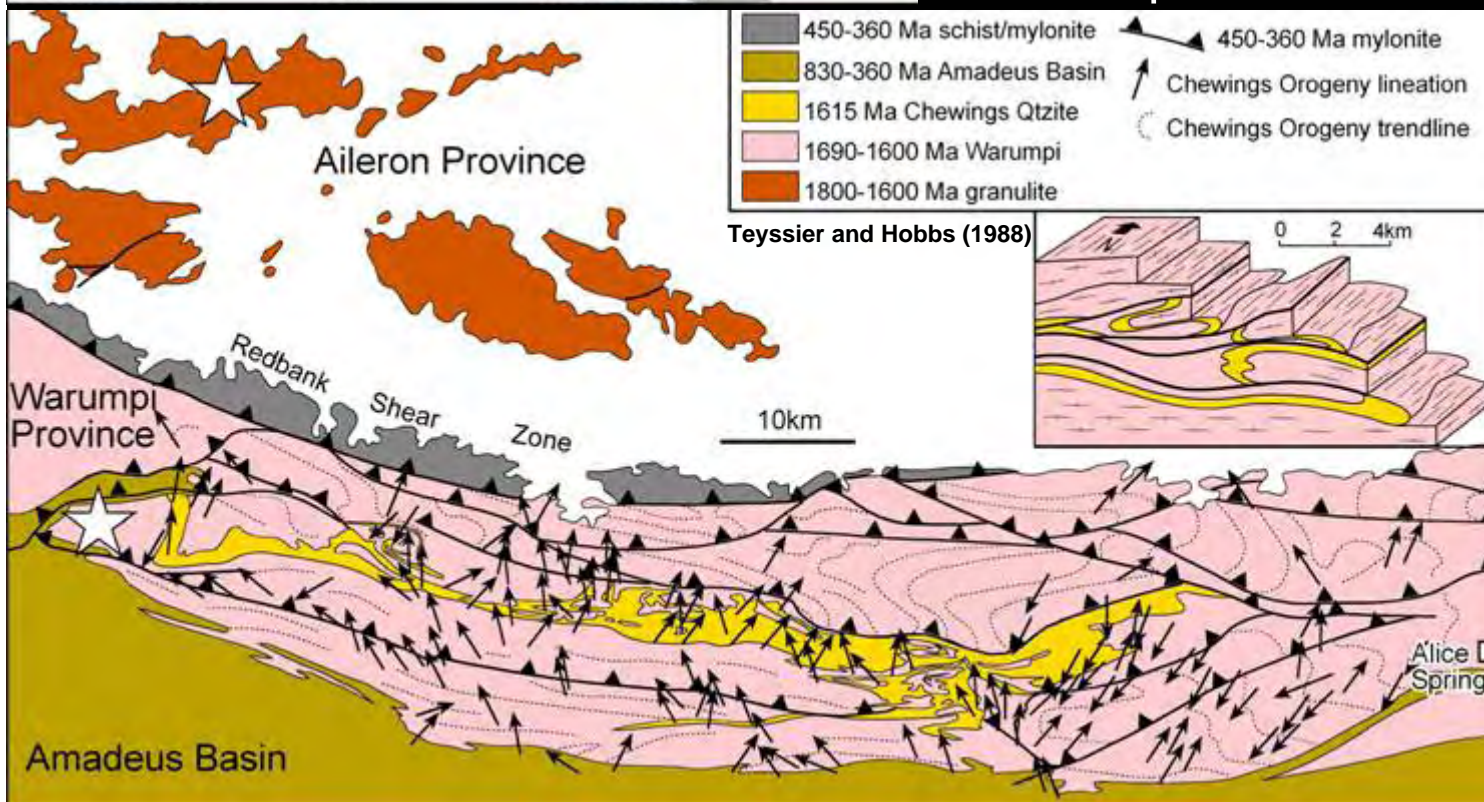


# Chewings Orogeny: Reworking in the Arunta: 1600-1570 Ma



## Chewings Range region:

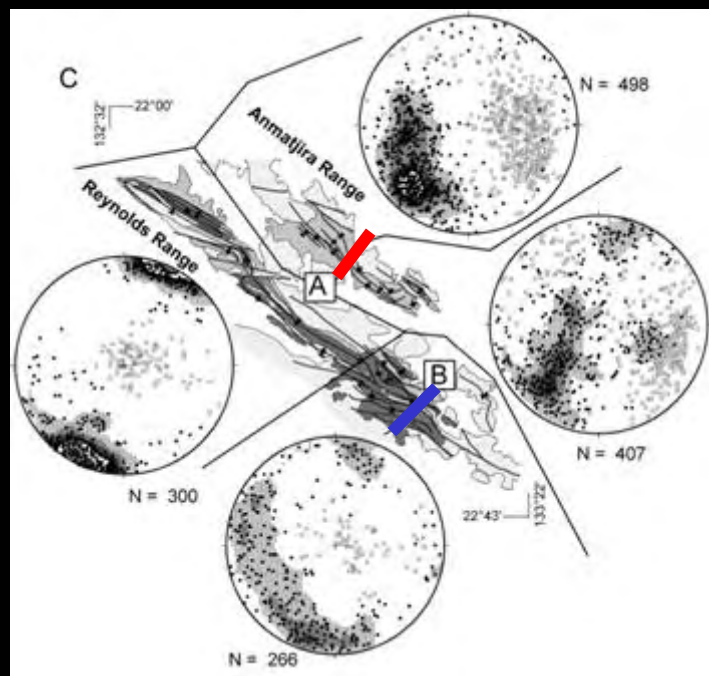
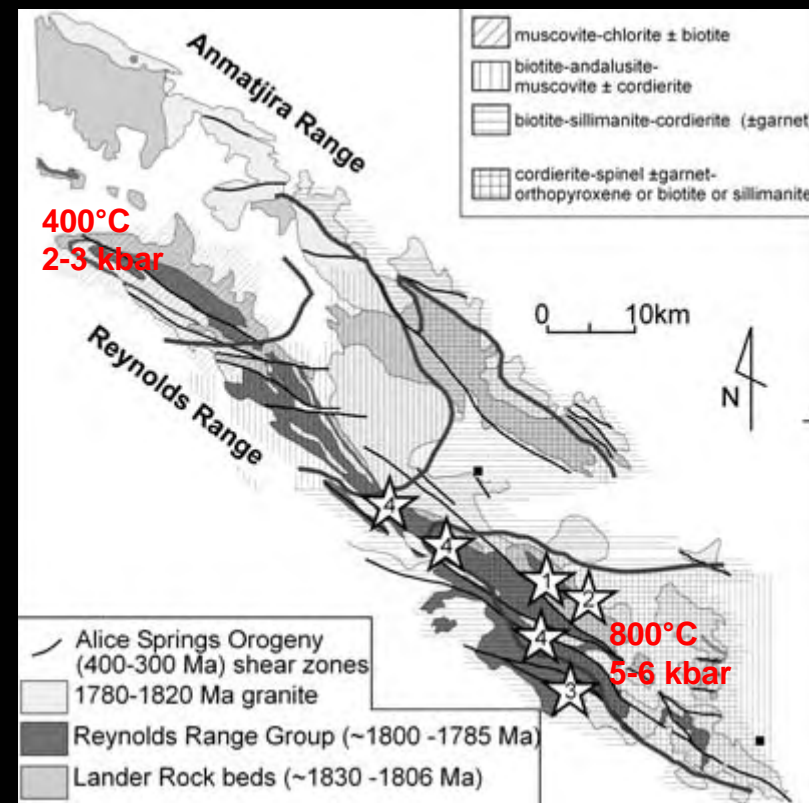
- High-strain large-scale top plate N transport
- Barrovian (600°C, 5-6 kbar) metamorphism.



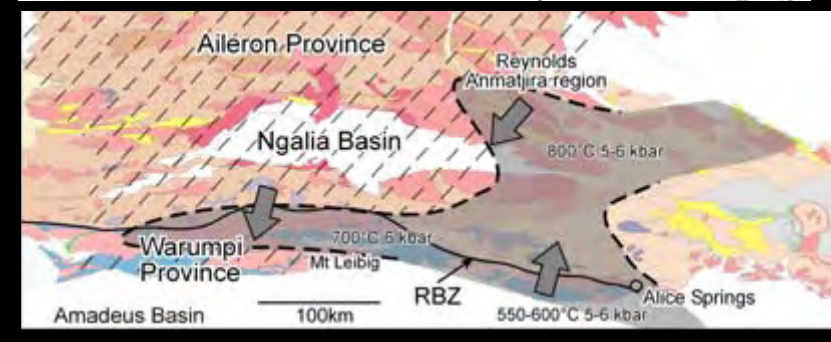
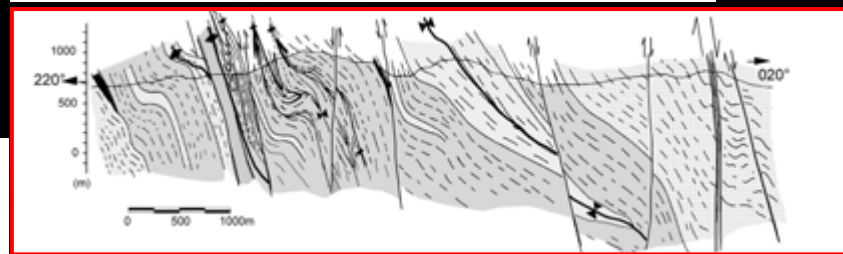
## SW Arunta:

- Top plate south
- Upper amphibolite-grade (700°C, 5-6 kbar) metamorphism.

# Chewings Orogeny: Reworking in the Arunta: 1600-1570 Ma



**Central Arunta Region:**  
 Greenschist to granulite grade HGG metamorphism  
 N-S shortening (~50%), top SW transport



# Chewings Orogeny: Reworking in the Arunta: 1600-1570 Ma

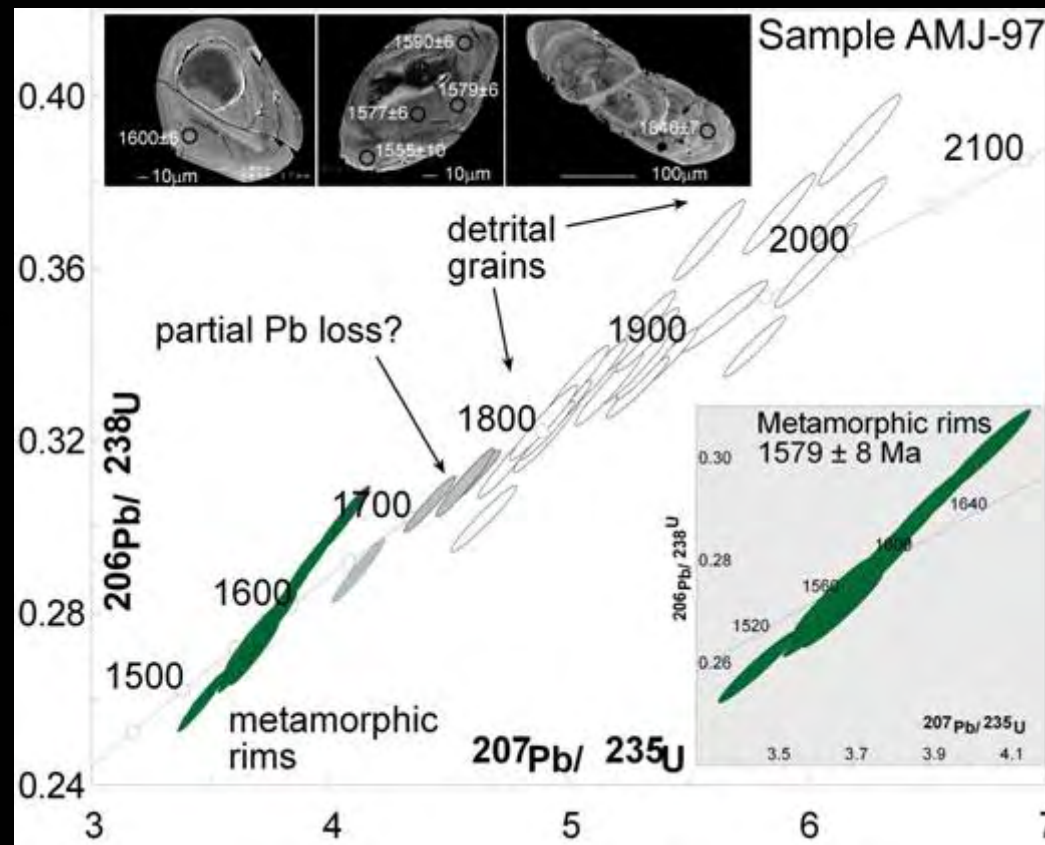
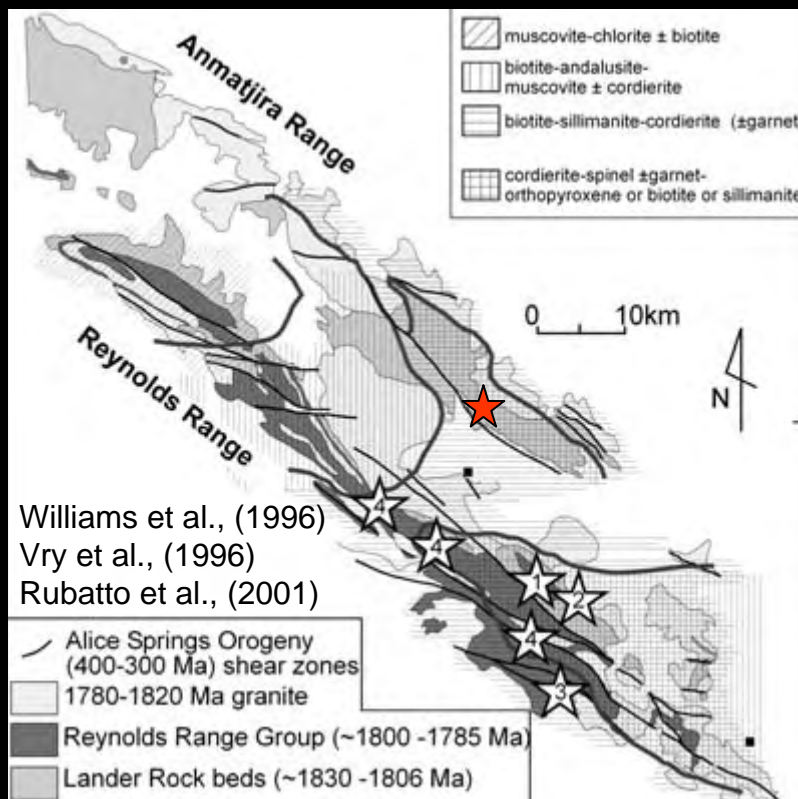


**Central Arunta Region:**  
N-S shortening, top SW transport

Greenschist to granulite grade HHG metamorphism

**Timing of peak metamorphism in the Anmatjira Range**

Garnet-cordierite-spinel-sillimanite-biotite-metapelite

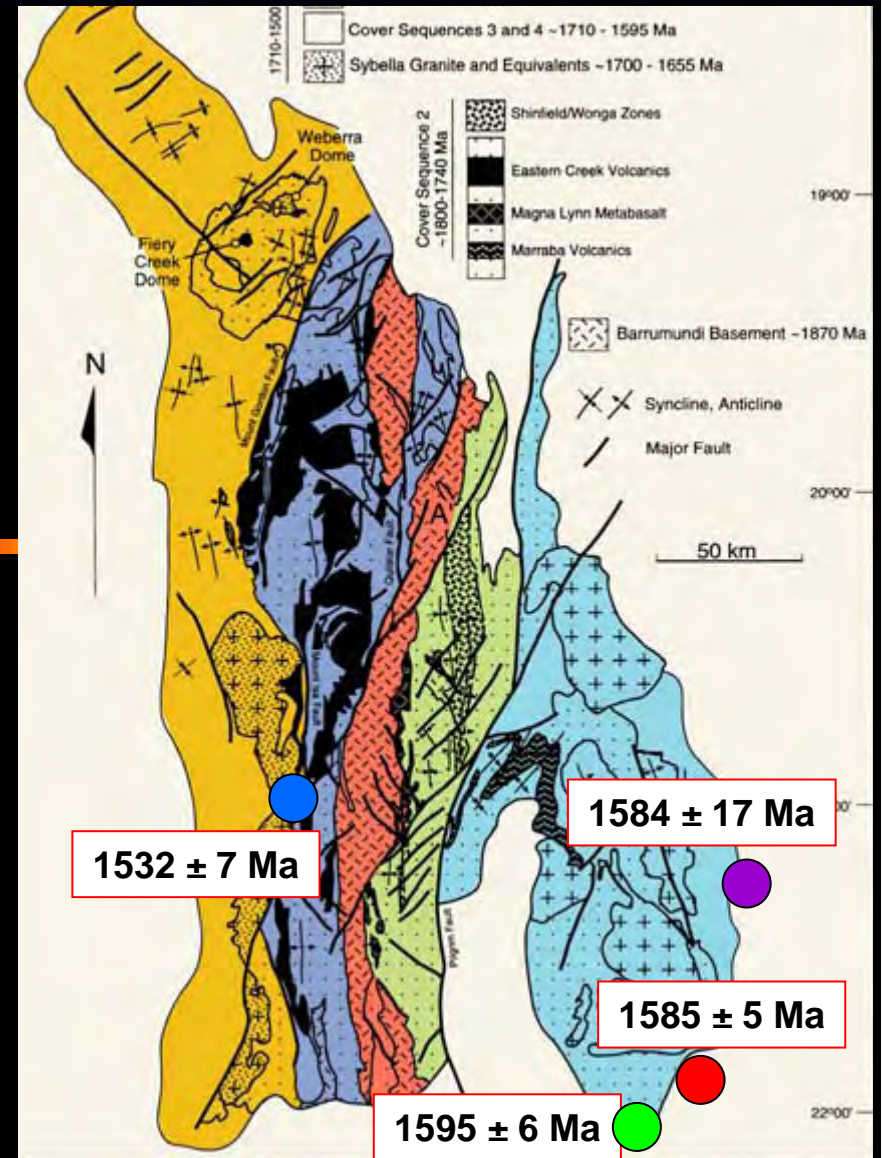


# Early Isan Orogeny: 1600-1570 Ma

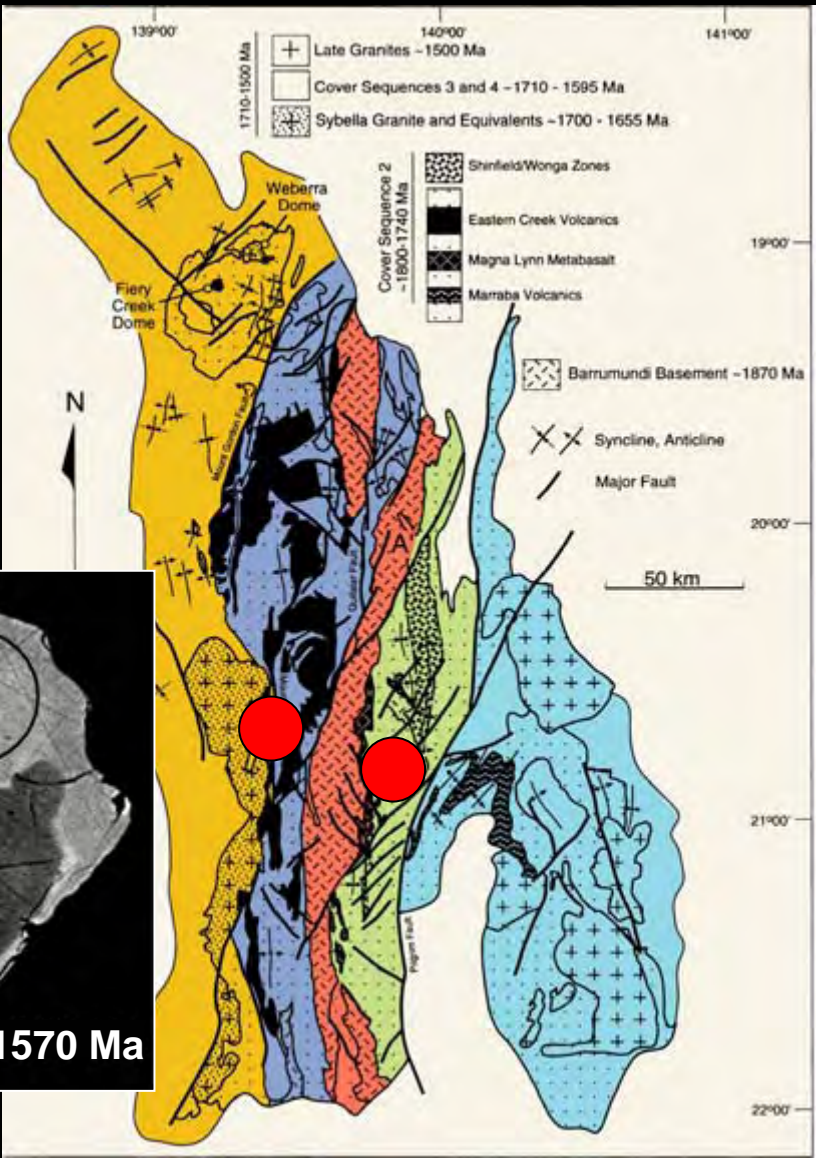
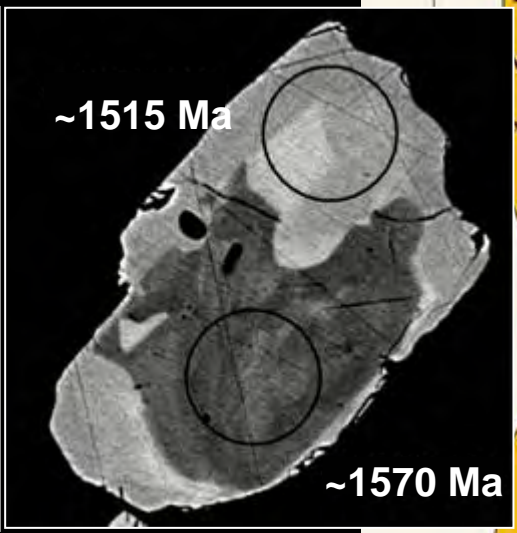
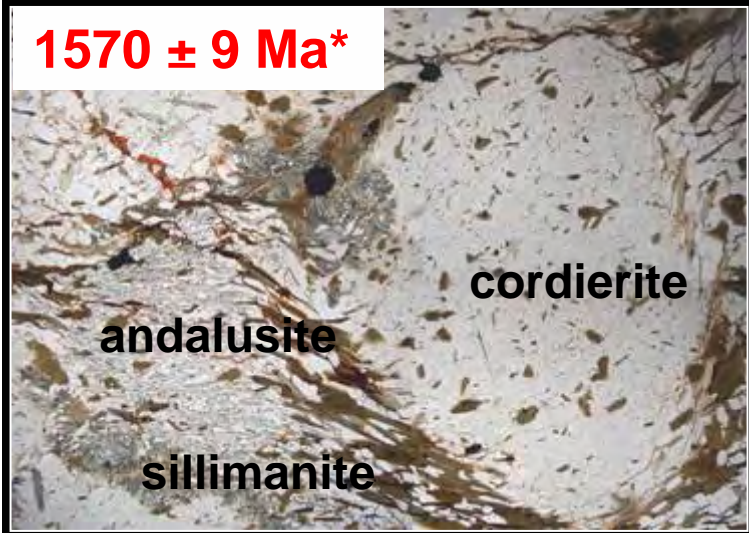
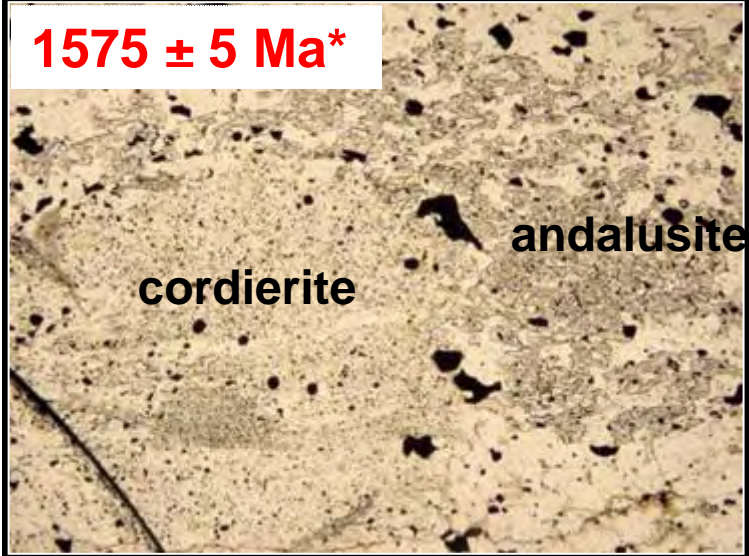
Inverted multi-stage rift complex  
 3 phases of basin development  
 between c. 1760-1595 Ma

## Isan Orogeny (c. 1600-1500)

- Conner & Page, (1995)
- Page & Sun, (1998)
- Rubenach et al., (2001)
- Giles & Nutman, (2002)

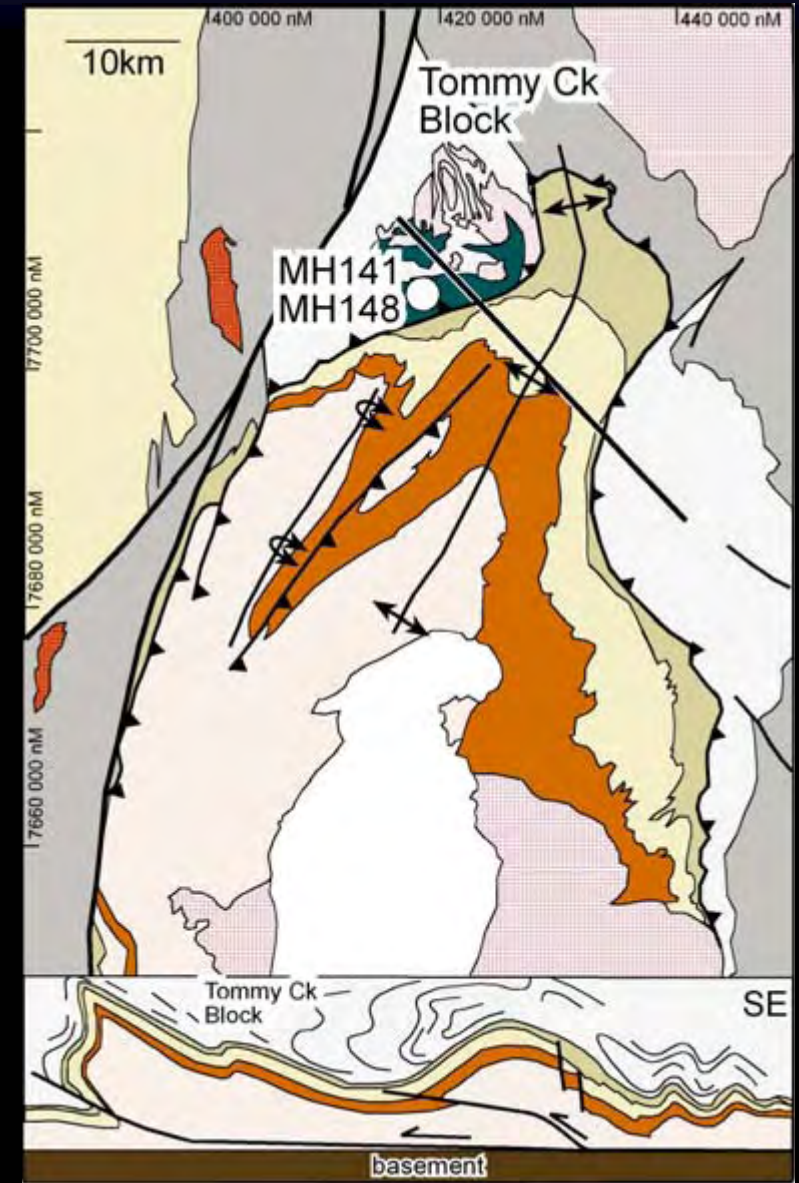
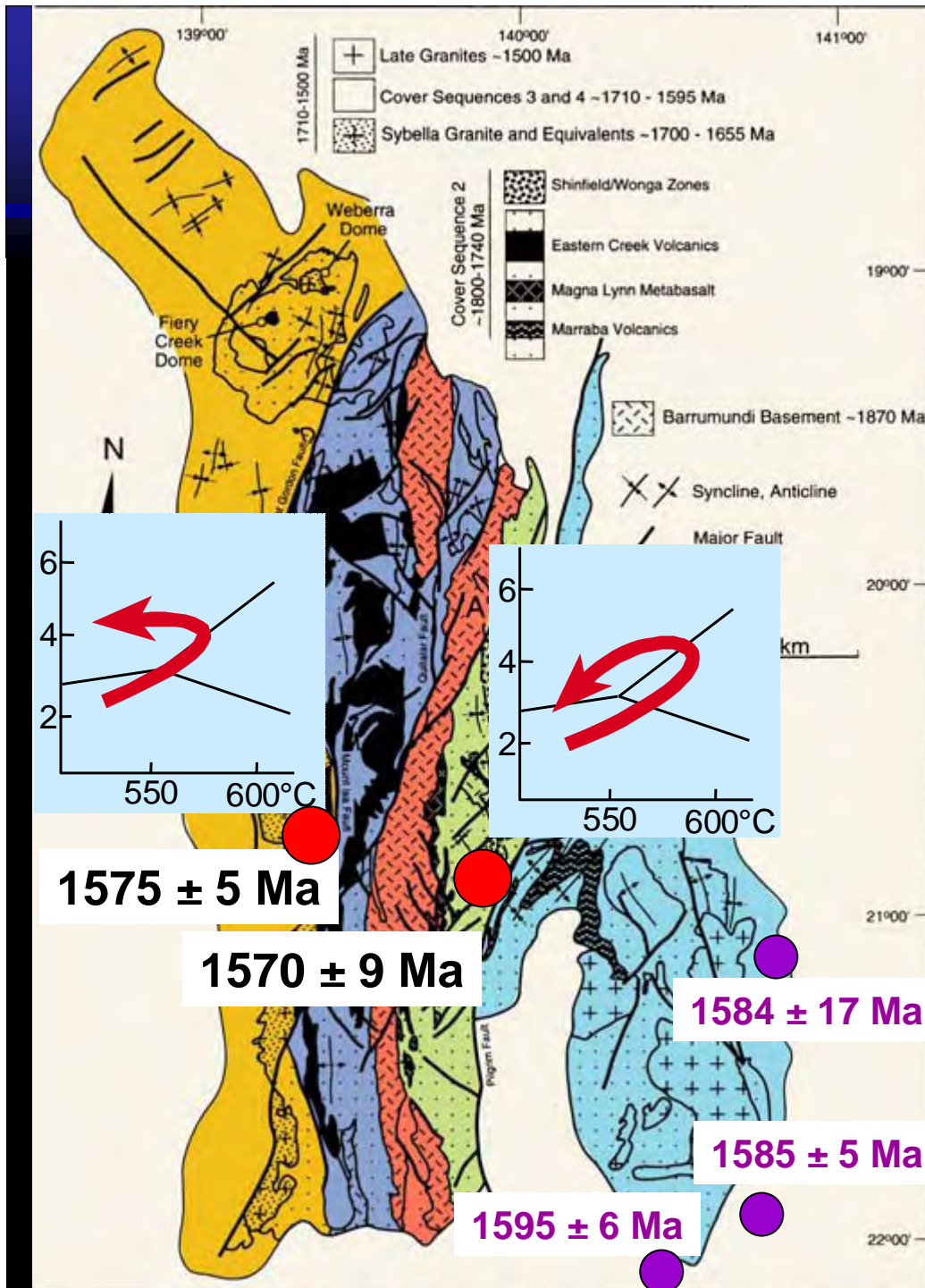


# Early Isan Orogeny: 1600-1570 Ma



\*Hand & Rubatto, (2002)

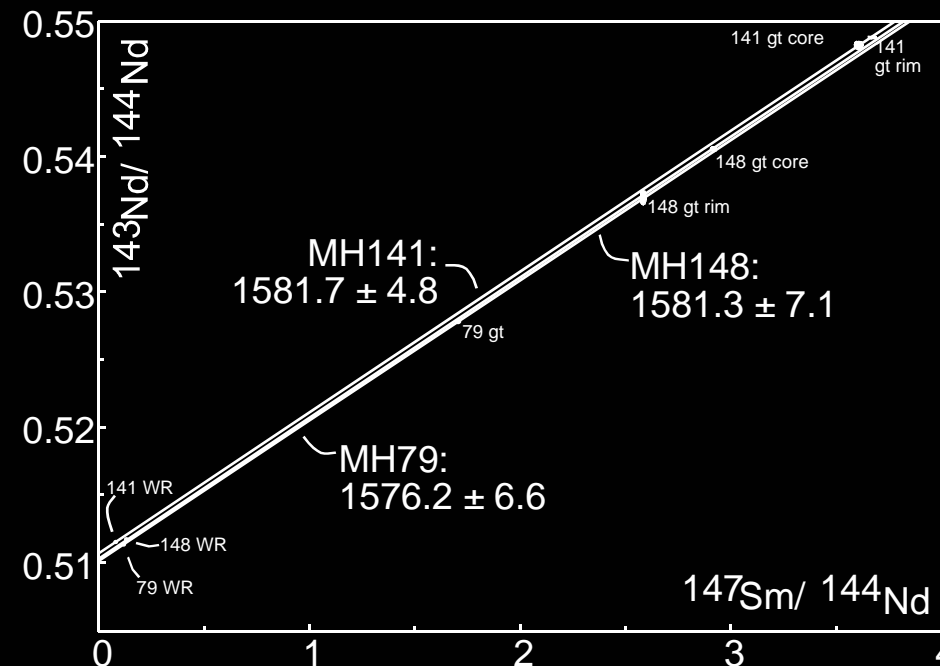
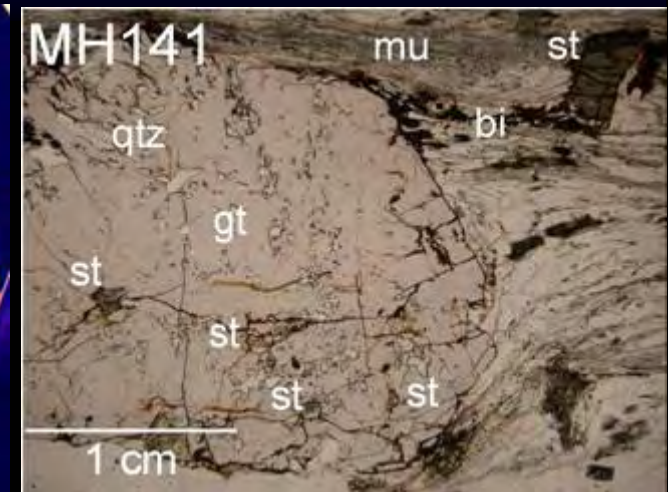
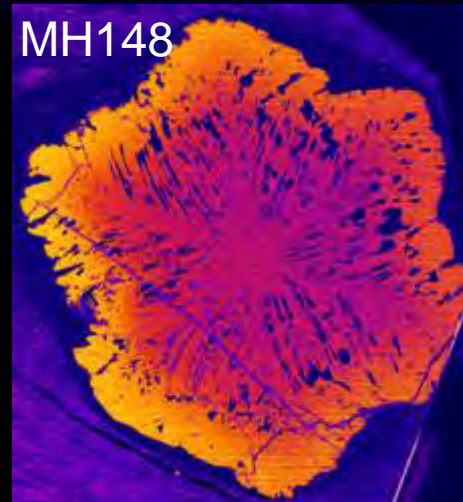
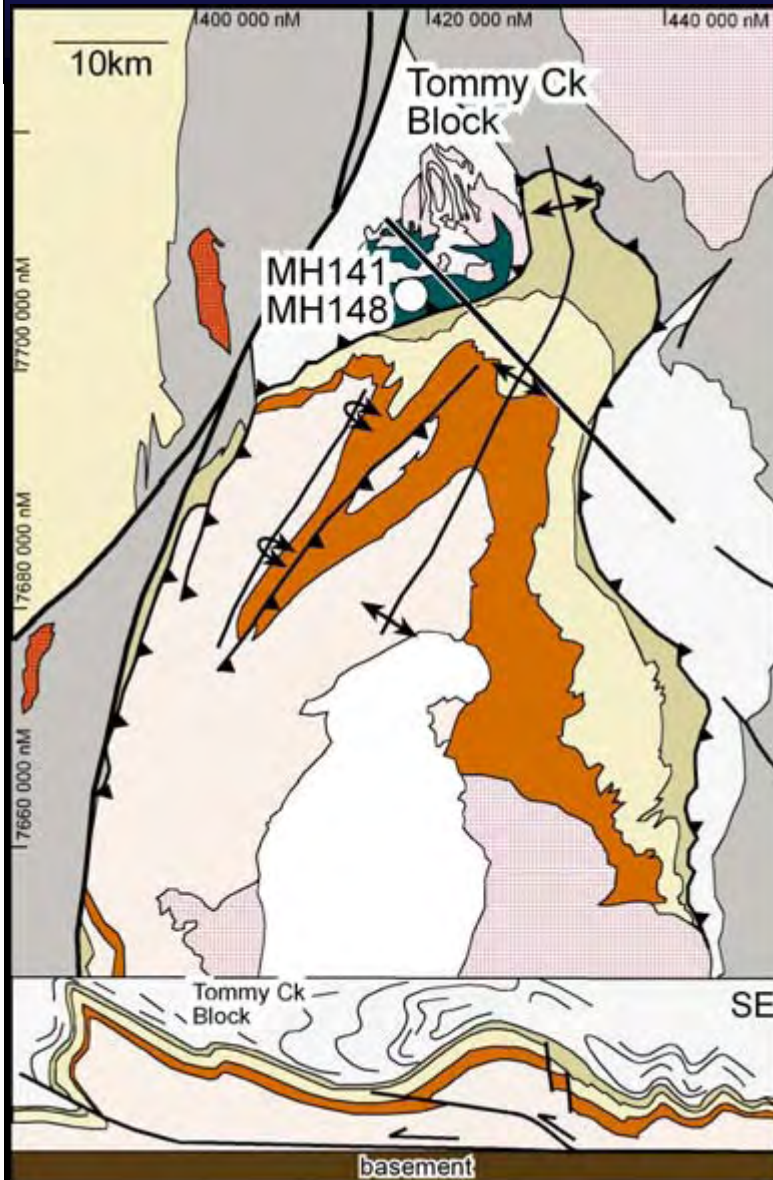
# Early Isan Orogeny: 1600-1570 Ma



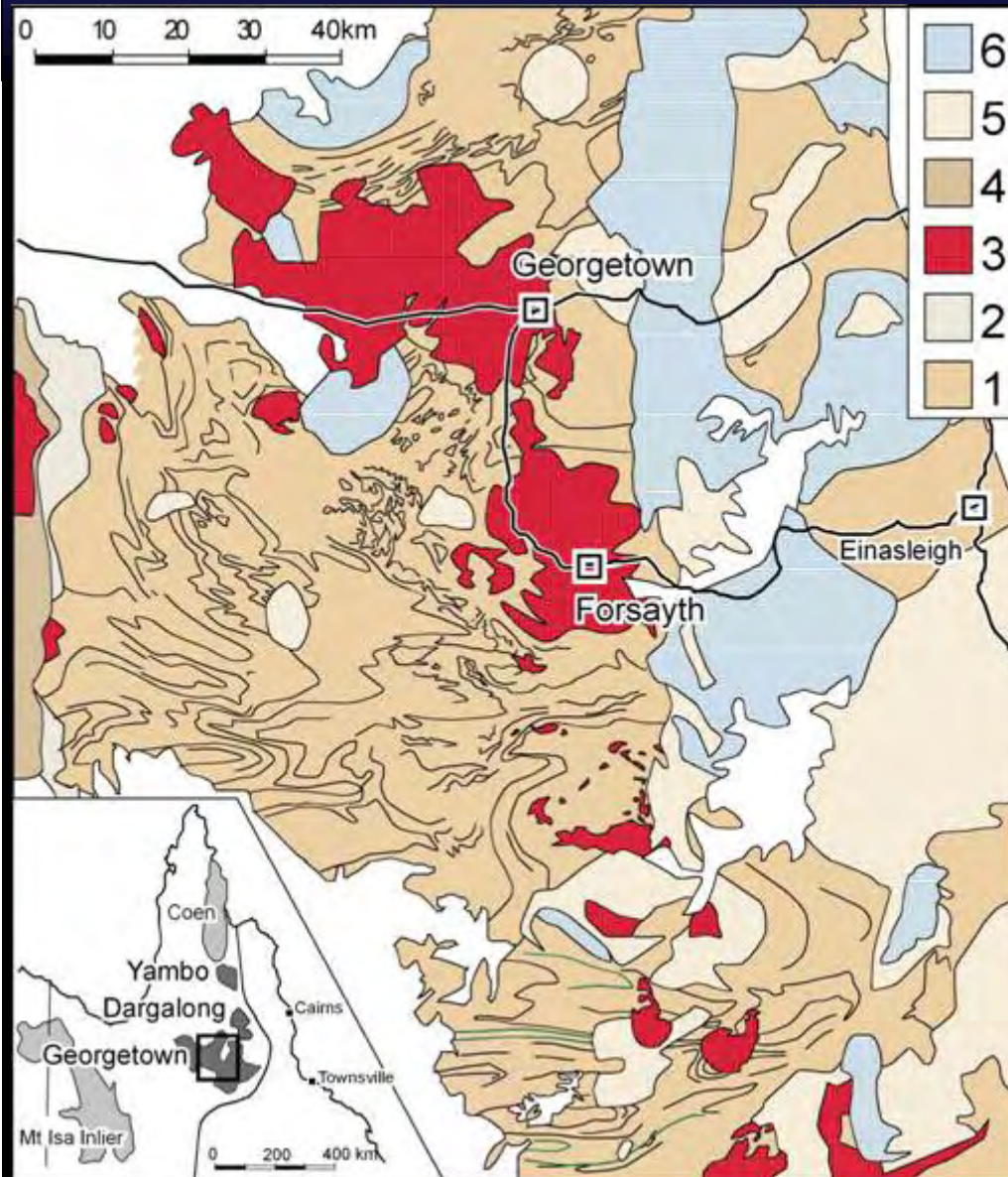
# Early Isan Orogeny: 1600-1570 Ma



Tommy Creek Block: prograde zoned up-P garnets



# Georgetown Inlier et al



## Georgetown Inlier

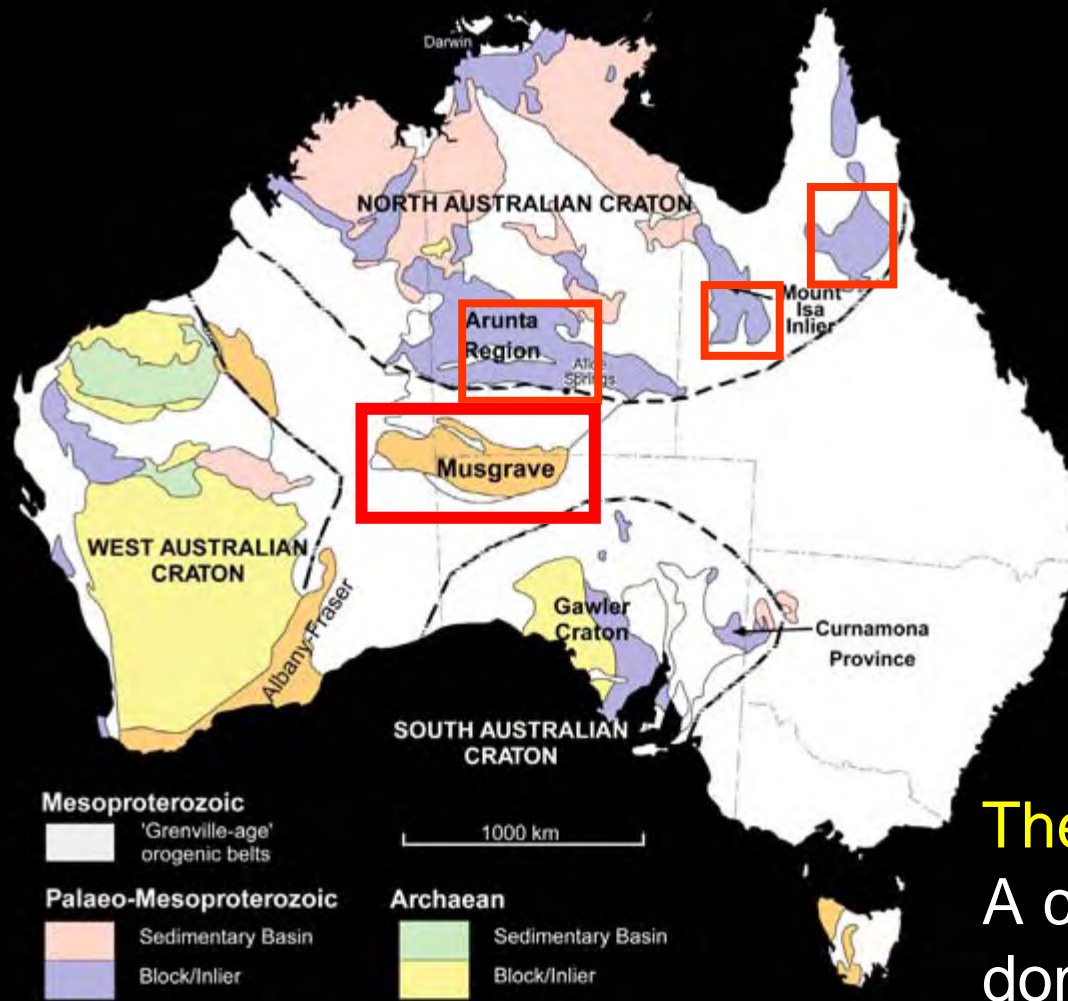
Etheridge Group (~ 1700-1600): equivalents to Isan SS 2 & 3?

Early N-S shortening & Up-P Barrovian amphibolite-grade metamorphism at and prior to  $1562 \pm 4$  Ma (Boger & Hansen, 2004; Black et al., 2005).

## Dargalong and Yambo Inliers

Granulite-grade metamorphism  $1585 \pm 5$  Ma -  $1576 \pm 5$  Ma. (Blewett et al., 1998).

# Intracratonic Orogeny in Mesoproterozoic Australia



## Summary of NAC ICO's

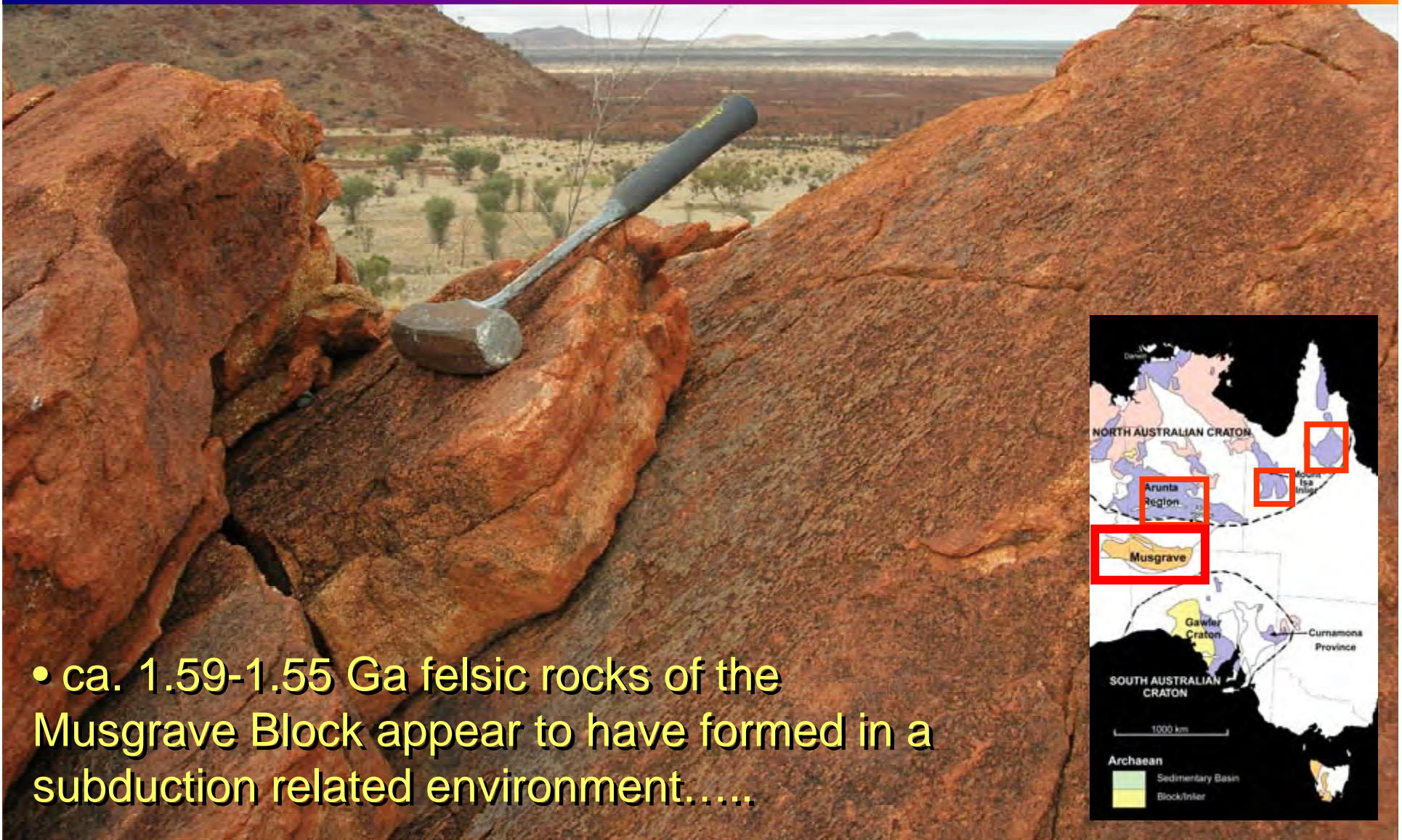
- ❖  $t = 1600-1570$ (ish) Ma
- ❖ High geothermal gradient's
- ❖ Contractional deformation

What else was happening in early Mesoproterozoic Australia?

## The Musgrave Block

A comparatively juvenile felsic-dominated volume separating the Gawler (SAP) from the NAC.

# The Musgrave margin? A driver for c. 1600-1570 Ma NAC ICO's?

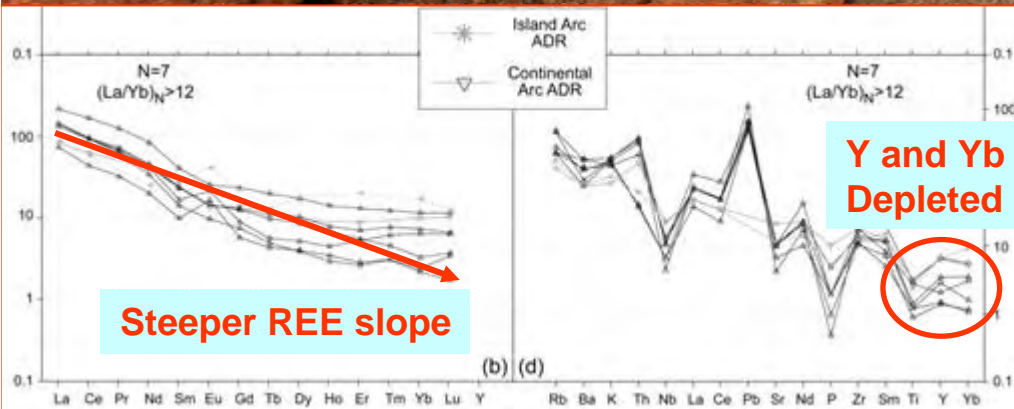
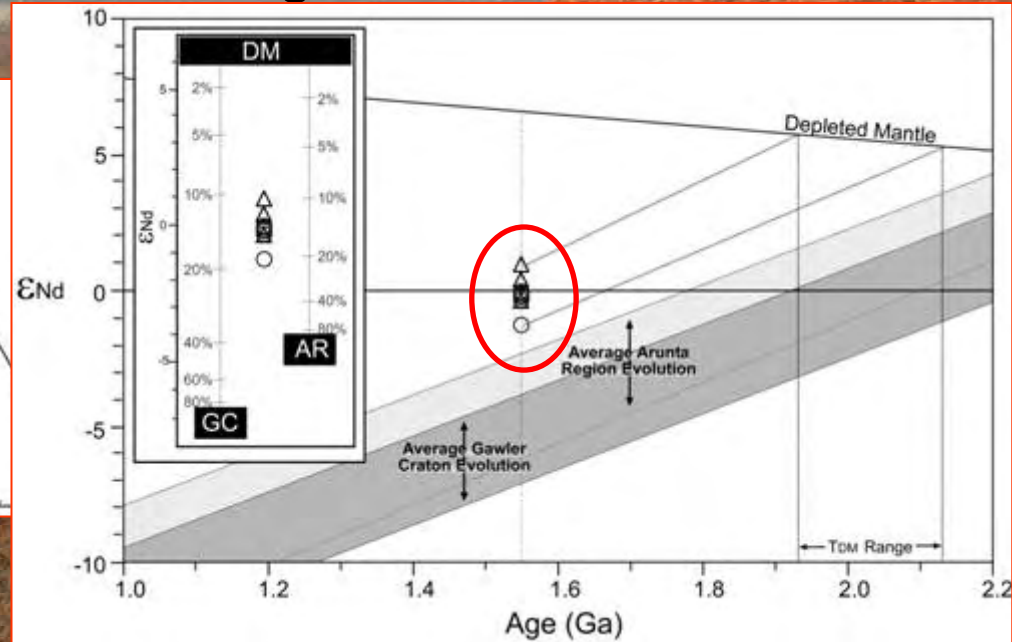
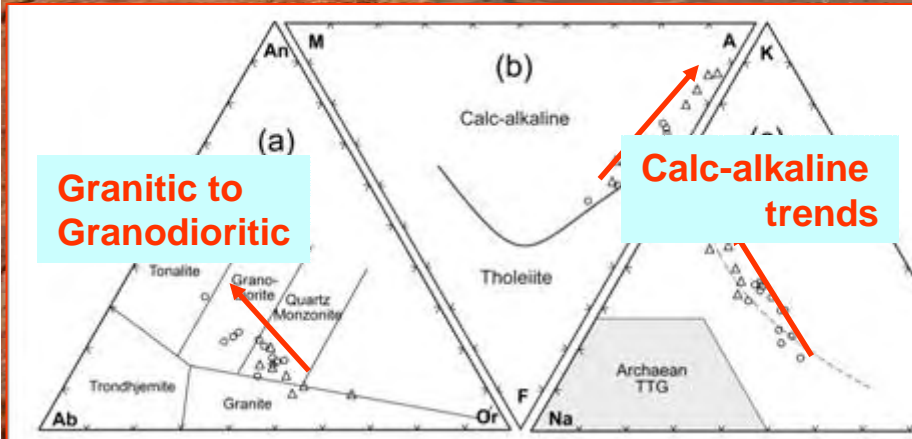


- ca. 1.59-1.55 Ga felsic rocks of the Musgrave Block appear to have formed in a subduction related environment.....

# The Musgrave margin? A driver for c. 1600-1570 Ma NAC ICO's?



- ca. 1.59-1.55 Ga felsic rocks in the Musgrave Block



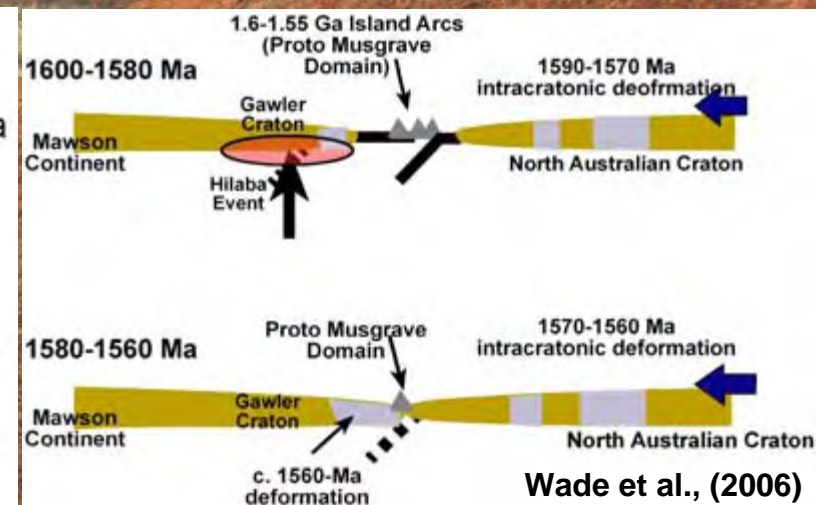
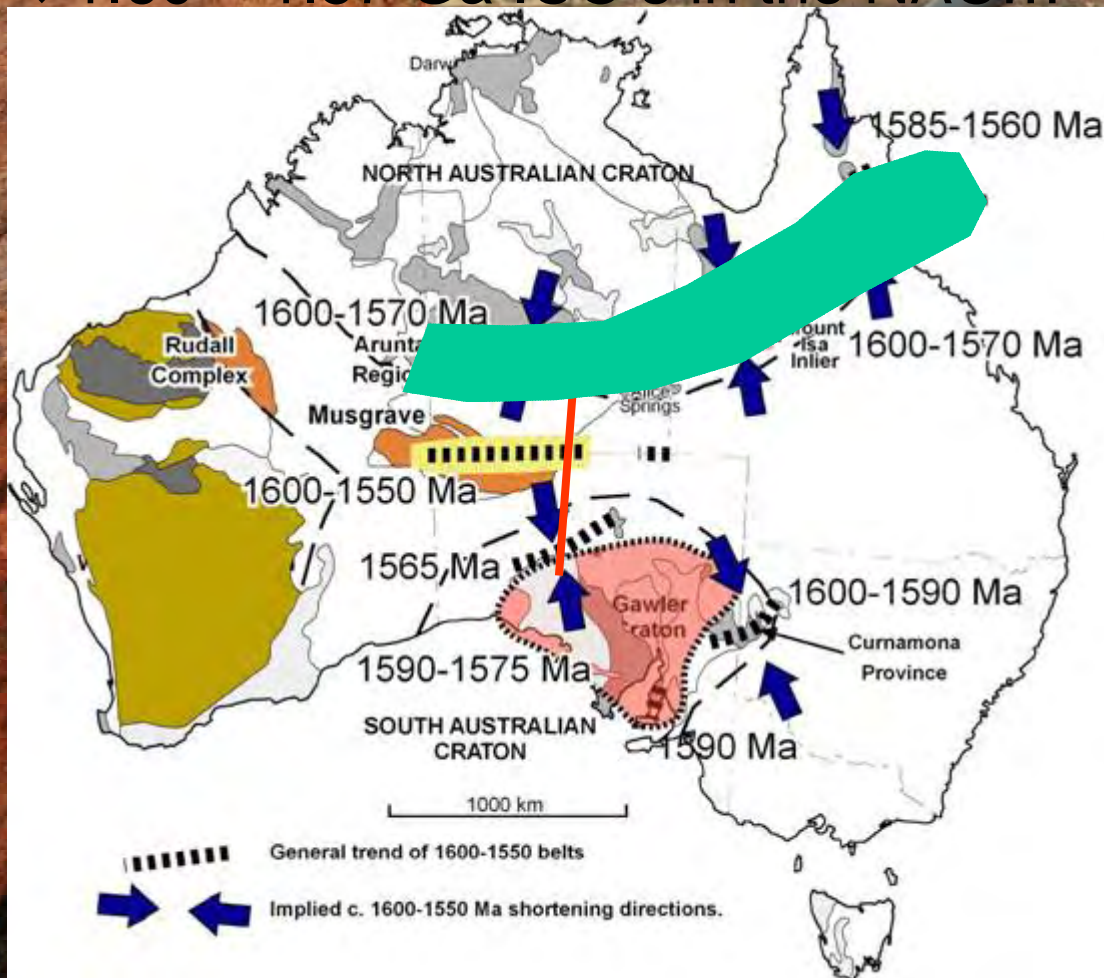
- Felsic rocks are dominantly juvenile with respect to surrounding older crustal domains, have calc alkaline trends and HEE depletion.

Wade et al., (2006)

# The Musgrave margin? A driver for c. 1600-1570 Ma NAC ICO's?



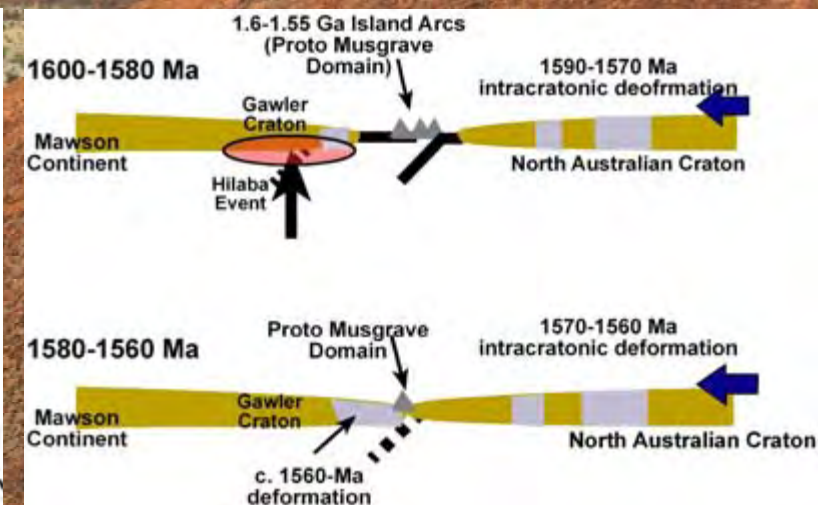
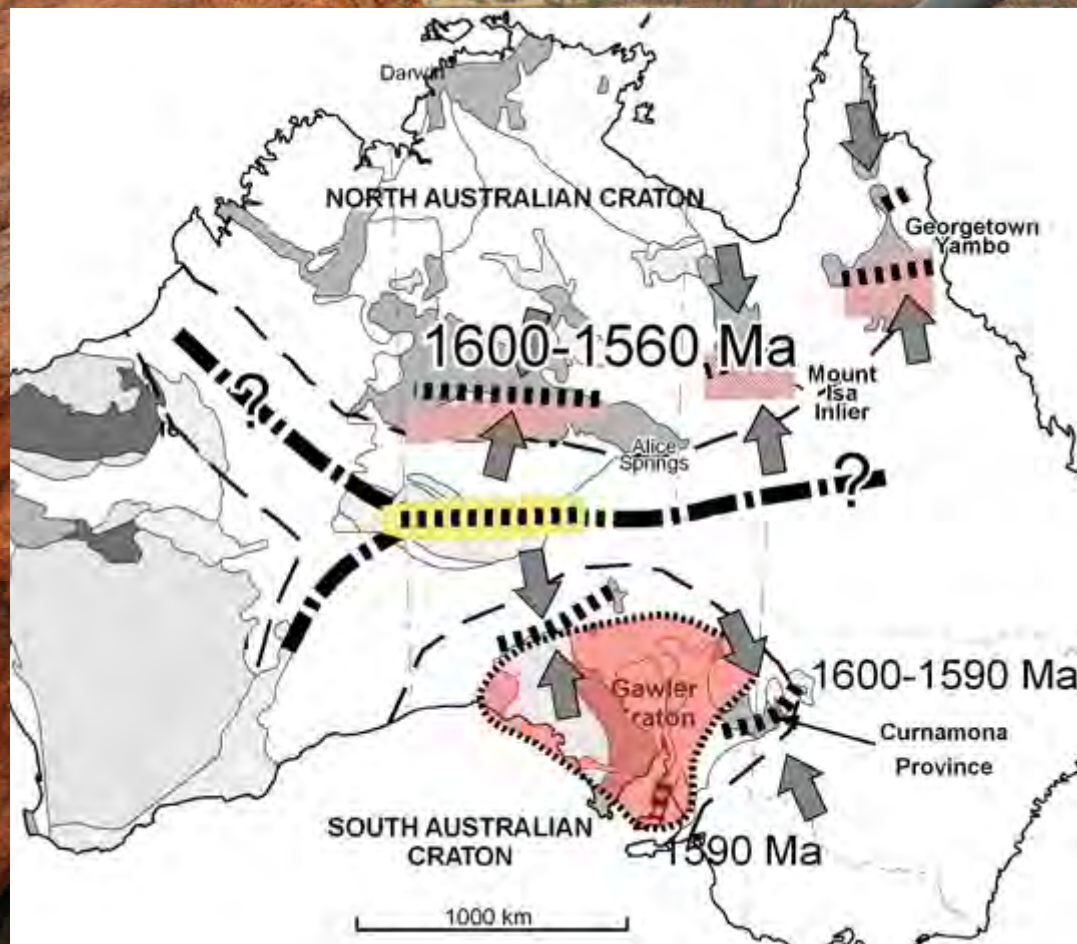
- ❖ ca. 1.59-1.55 Ga arc-related felsic rocks in the Musgrave Block?
- ❖ 1.60 – 1.57 Ga ICO's in the NAC...



- ❖ Is there an early Mesoproterozoic IC orogenic belt in the NAC that mirrors the Musgrave margin?

# The Musgrave margin? A driver for c. 1600-1570 Ma NAC ICO's?

If there was a margin – where might it go?



Is there an early Mesoproterozoic suture beneath the Canning Basin that forms part of the Musgrave margin?

# What controlled the IC responsiveness of the NAC?



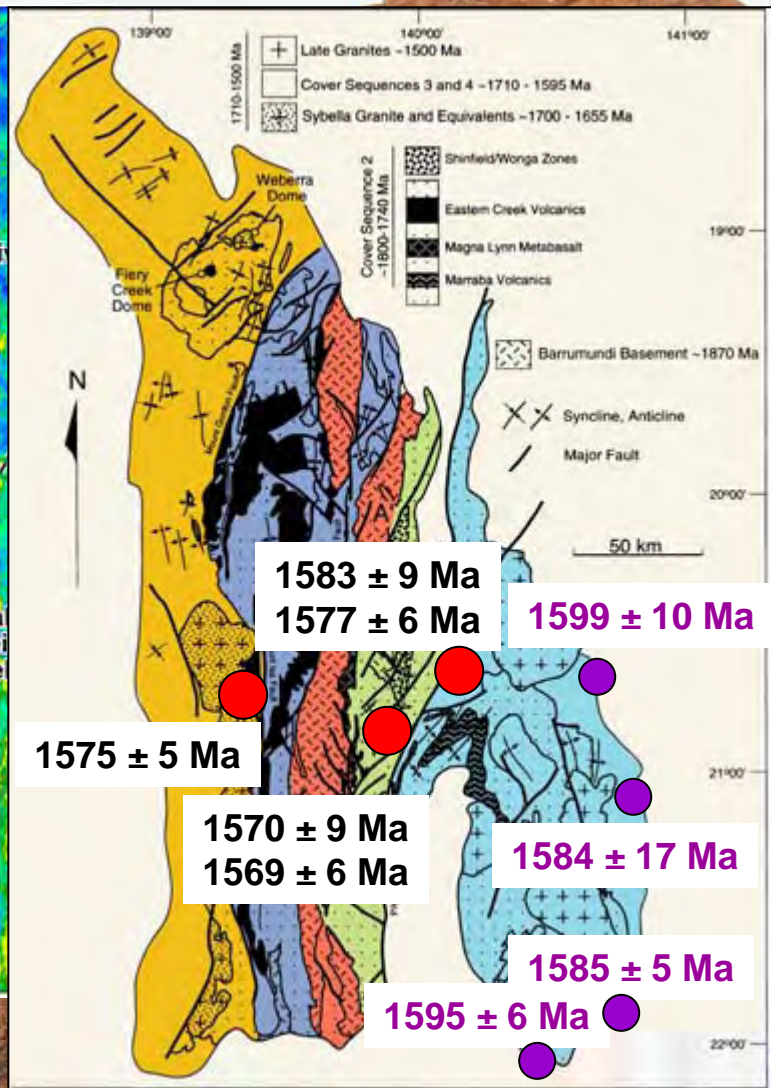
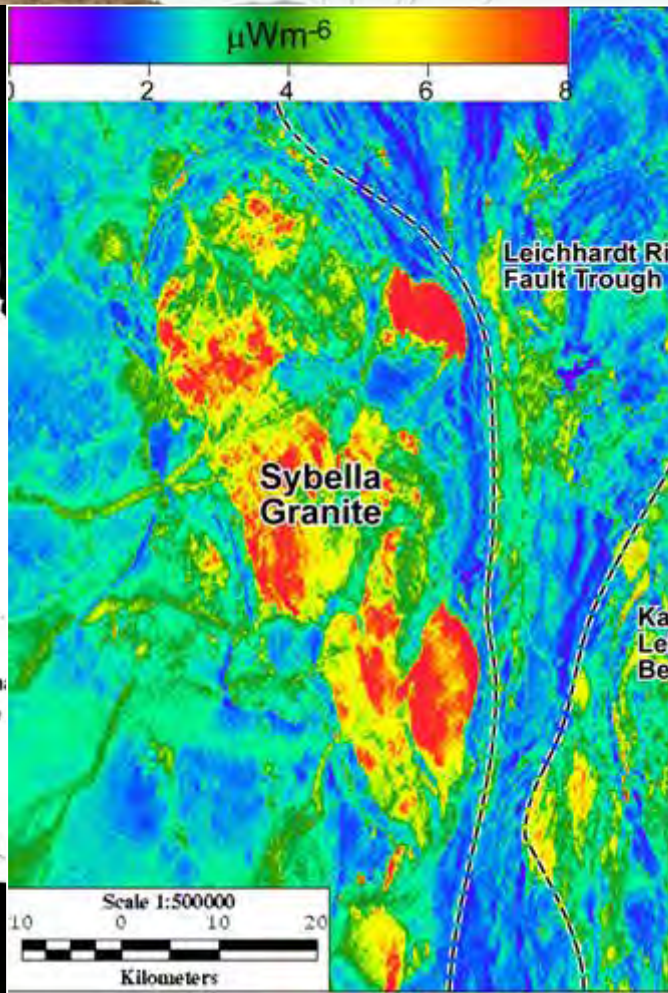
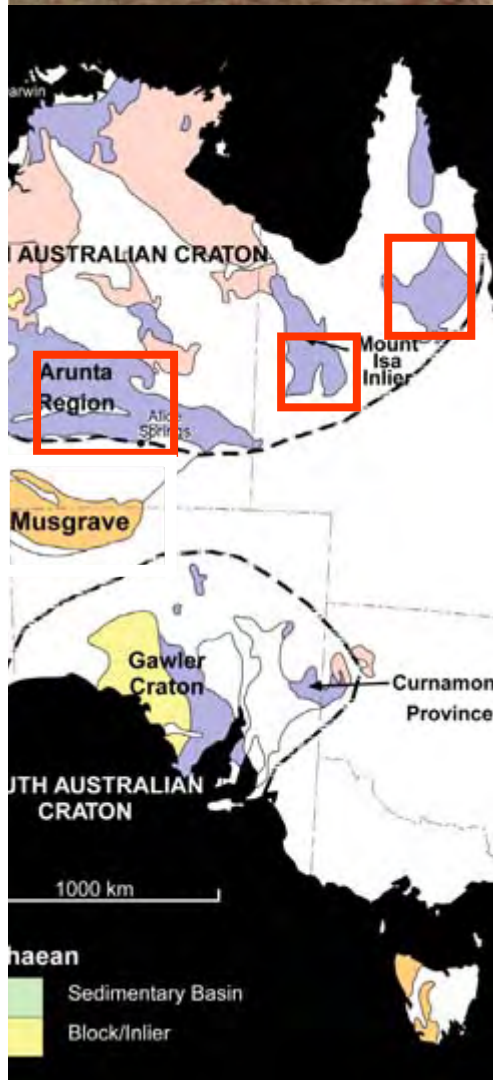
Chewings, Early Isan and Georgetown-Yambo systems had GTG  $\sim 40-60^{\circ}\text{C km}^{-1}$ .

Early Isan Orogeny overlapped with basin development.

Were these IC orogens localised by hot pre-orogenic basins?

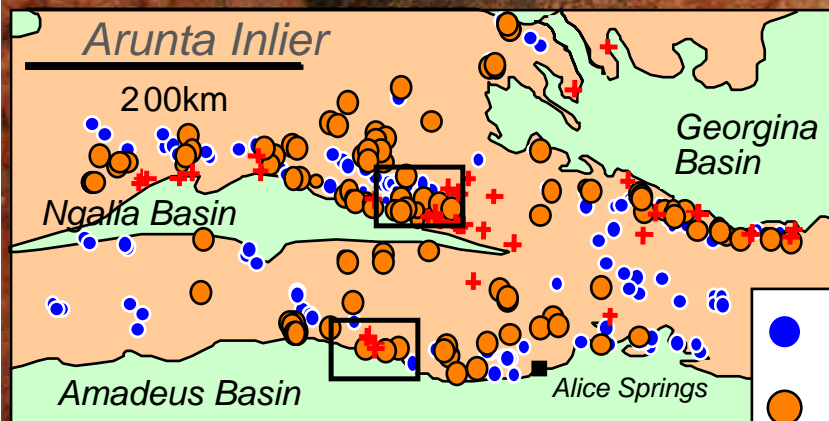
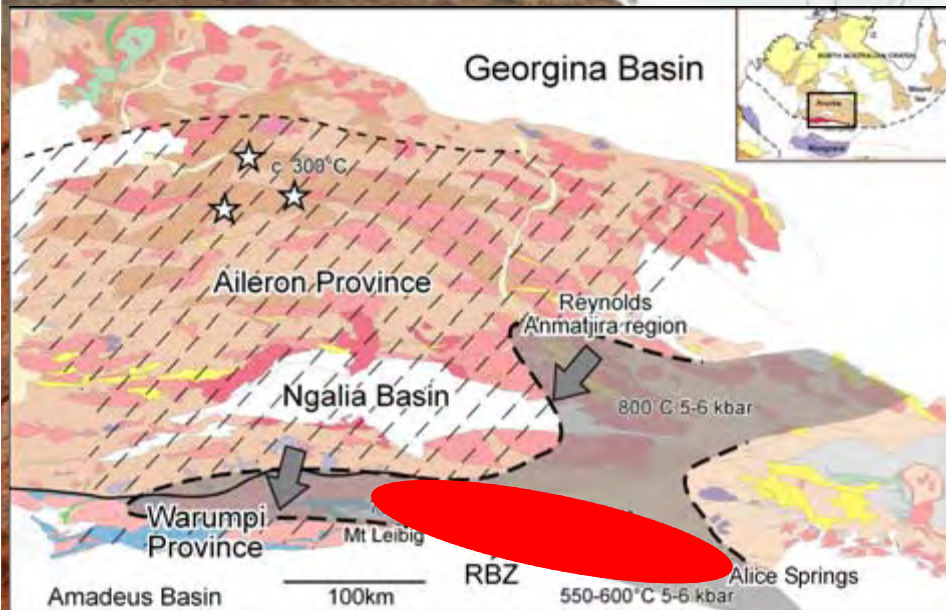


# What controlled the IC responsiveness of the NAC?

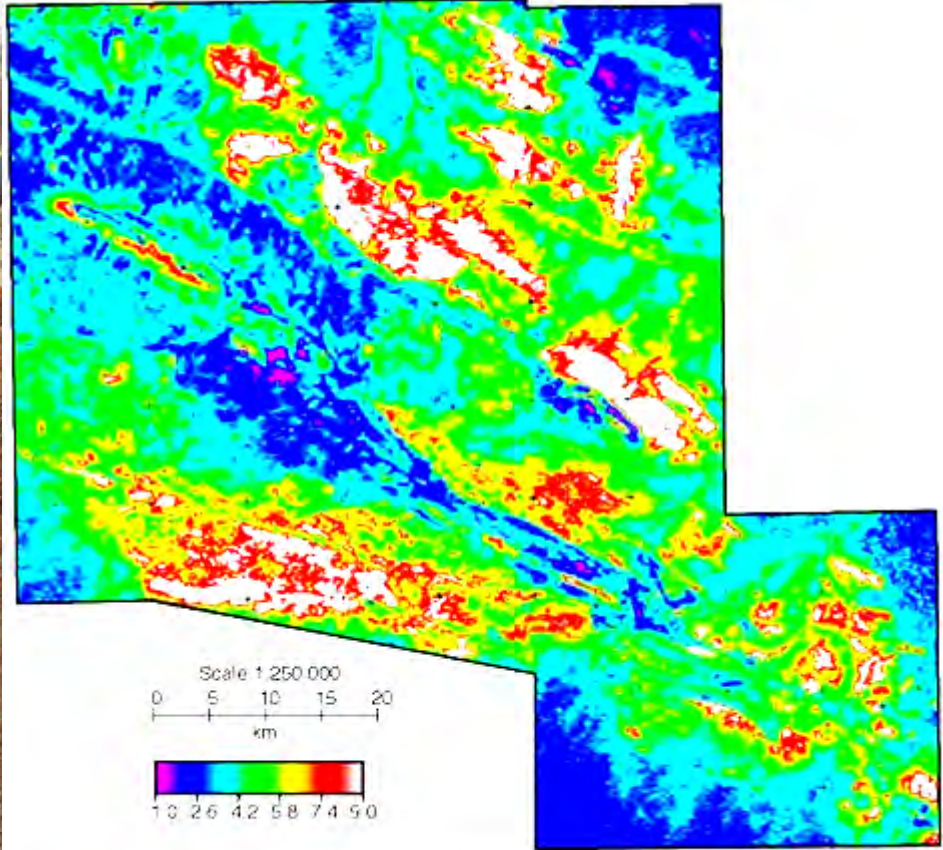


e.g. McLaren et al., (1999)

# What controlled the IC responsiveness of the NAC?



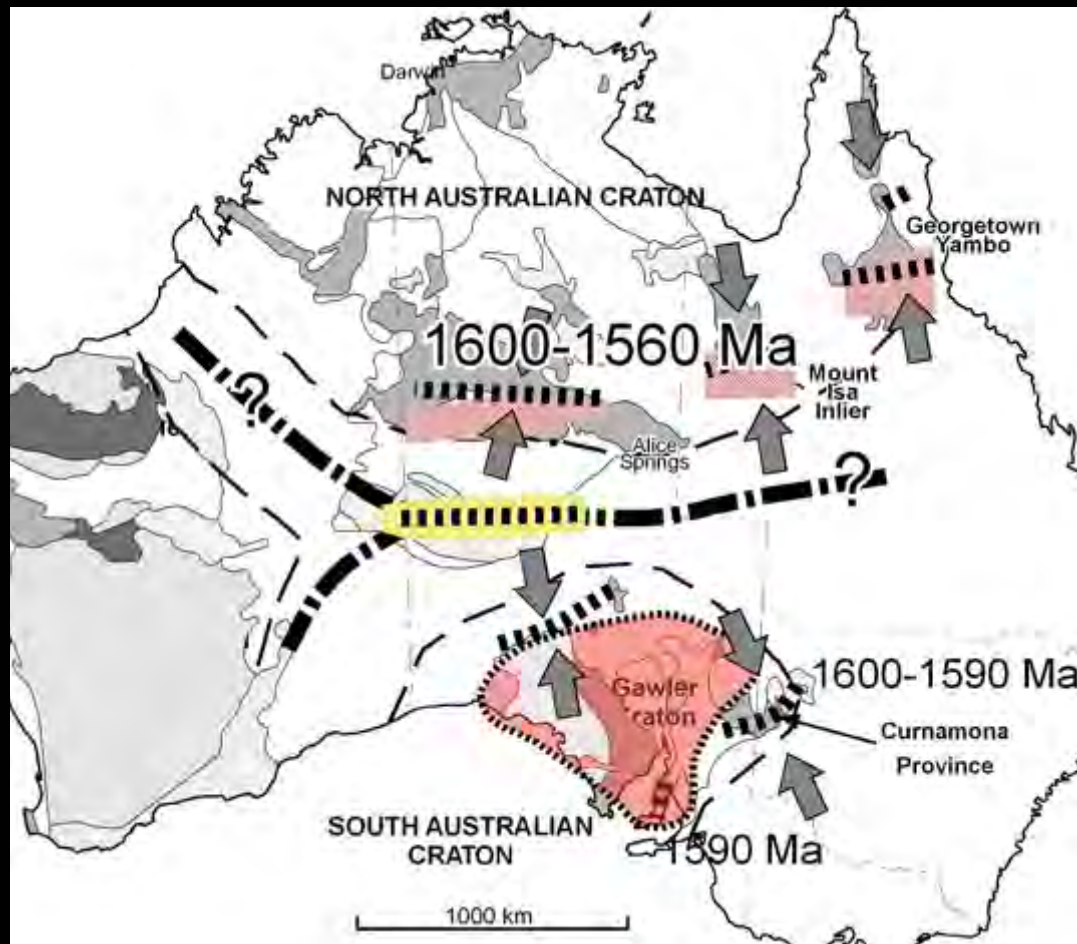
- $< 5 \mu\text{Wm}^{-3}$
- $5-10 \mu\text{Wm}^{-3}$
- +



Reynolds Range: central Arunta

Hot basin development in the southern Arunta @ c. 1620 Ma

# Intracratonic Orogeny in Mesoproterozoic Australia



ca. 1600-1570 Ma high-T reworking of the NAC was synchronous with the development of a juvenile arc system (Musgrave Province).

N-S IC shortening in the NAC reflects convergence on the margin.

IC responsiveness of NAC inherited from pre-orogenic basins and localized hhp crust.

Is there an early Mesoproterozoic suture beneath the Canning Basin?

# Southern Australian Proterozoic



## Curnamona Province

Inversion of 1715-1640.....Ma Willyama Superbasin

HGG metamorphism @ 1610-1590 Ma

