Airborne Gravity 2004

Abstracts from the ASEG-PESA Airborne Gravity 2004 Workshop

Edited by Richard Lane
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# Contents

"Airborne Gravity 2004 Workshop" Record
R.J.L. Lane, M.H. Dransfield, D. Robson, R.J. Smith and G. Walker......................................................v

Requirements and general principles of airborne gravity gradiometers for mineral exploration
F. van Kann..................................................................................................................................................1

The Air-FTG™ airborne gravity gradiometer system
C.A. Murphy..............................................................................................................................................7

The FALCON® airborne gravity gradiometer systems
M.H. Dransfield and J.B. Lee.......................................................... .............................................................15

A superconducting gravity gradiometer tool for exploration
J.M. Lumley, J.P. White, G. Barnes, D. Huang and H.J. Paik.................................................................21

A high resolution airborne gravimeter and airborne gravity gradiometer
B. Tryggvason, B. Main and B. French.......................................................... ........................................41

The AIRGrav airborne gravity system
S. Sander, M. Argyle, S. Elieff, S. Ferguson, V. Lavoie and L. Sander......................................................49

The GT-1A mobile gravimeter
A. Gabell, H. Tuckett and D. Olson.......................................................... ...............................................55

A synopsis of airborne gravity data presentation, modelling, interpretation and inversion software
G. Walker..................................................................................................................................................63

Acquisition and preliminary impressions of airborne gravity gradient and aeromagnetic data in the Eastern Papuan Basin, Papua New Guinea
A. Nelson................................................................................................................................................65

Evaluation of a full tensor gravity gradiometer for kimberlite exploration
D. Hatch....................................................................................................................................................73

Integrating ground and airborne data into regional gravity compilations
R.J.L. Lane................................................................................................................................................81

Airborne gravity data acquisition and processing: A case study in the Prince Charles Mountains, East Antarctica
M. McLean, D. Damaske, V. Damm and G. Reitmayr........................................................................99

AIRGrav airborne gravity survey in Timmins, Ontario
S. Elieff and S. Sander..............................................................................................................................111

Examples of FALCON™ data from diamond exploration projects in Northern Australian
D. Isles and I. Moody...............................................................................................................................121

A comparison of the Falcon® and Air-FTG™ airborne gravity gradiometer systems at the Kokong Test Block, Botswana
D. Hinks, S. McIntosh and R.J.L. Lane.................................................................................................125

Analysis of errors in gravity derived from the Falcon® Airborne Gravity Gradiometer
D.B. Boggs and M.H. Dransfield..............................................................................................................135
Preface

The "Airborne Gravity 2004 Workshop" was held in Sydney on August 15, in conjunction with ASEG-PESA Sydney 2004 (the ASEG's 17th Geophysical Conference and Exhibition). The aims of the workshop were to provide participants with a review of the current state of the art in airborne gravity instrumentation, to present case histories of the use of these methods in minerals and petroleum applications, and to distribute sample data sets. "Airborne gravity" is used in this context to include both airborne gravimeter and airborne gravity gradiometer methods.

The program was split into 2 sessions. The morning session provided a review of the systems, with presentations covering a number of systems currently in operation as well as some that are still under development. The focus shifted in the afternoon session to case histories, with examples from surveys spanning the globe; from Antarctica to the tropics of Papua New Guinea, from Africa through Australia to Canada.

To capture the essence of the day and to promote the ongoing development of airborne geophysical methods, speakers were invited to submit papers for inclusion in a workshop volume. The papers were reviewed prior to publication in this Geoscience Australia Record. Participants received a copy at the workshop, and additional copies of the Record are available on an ongoing basis from Geoscience Australia (www.ga.gov.au).

Units

Physical quantities should be expressed in SI units. The Bureau International des Poids et Mesures (BIPM) is the custodian of this system. To quote from their website (www.bipm.fr): "Its mandate is to provide the basis for a single, coherent system of measurements throughout the world, traceable to the International System of Units (SI)".

The SI unit for acceleration is “metre per second squared” (m/s²). The signals encountered in gravity surveys for exploration are small, and the prefix “micro” is commonly used (micrometre per second squared, µm/s²). The gal (or Gal), equal to 1 cm/s², is a derived unit for acceleration in the CGS system of units. A prefix of “milli” is commonly used (milligal, mGal). In rare cases in the literature, a “gravity unit” (gu) may be encountered. In this publication, the µm/s² has been the preferred unit for gravity measurements, but mGal has been accepted.

1 µm/s² = 10⁻⁶ m/s²
1 mGal = 10 µm/s²
1 gu = 1 µm/s²

The gravity gradient is a gradient of acceleration and so the appropriate units are acceleration units divided by distance units. Thus, “per second squared” (s⁻²) is appropriate in the SI system. Typical gravity gradients measured in exploration are extremely small, and the prefix “nano” is appropriate in most circumstances (per
nanosecond squared, $\text{ns}^{-2}$). The eotvos unit (Eo), although not recognised in either the SI or CGS systems, is used almost universally in geophysics as the unit for gravity gradient measurements. It is equal to $1 \text{ ns}^{-2}$. In this publication, the $\text{ns}^{-2}$ and Eo have both been accepted as units for gravity gradient measurements.

$$1 \text{ ns}^{-2} = 10^{-9} \text{ s}^{-2}$$
$$1 \text{ Eo} = 1 \text{ ns}^{-2}$$

Acknowledgments

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The AIRGrav airborne gravity system

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Abstract
An overview of the AIRGrav system (Airborne Inertially Referenced Gravimeter) is presented. Several design features and their positive effect on survey operations and system accuracy are discussed. These attributes are demonstrated in a case study of an AIRGrav survey of the Turner Valley gas and oil production field in Alberta, Canada.

Introduction
In 1992, Sander Geophysics Limited undertook a project to design an airborne gravimeter that would be well suited to the dynamic environment of an aircraft. The five year research and development project resulted in the AIRGrav system, which has been used for airborne surveying since 1997. The majority of these surveys have been conducted for the petroleum industry, as well as some applications to mining exploration. At present, four AIRGrav systems are employed in worldwide survey operations.

The AIRGrav system
The AIRGrav system consists of a three-axis gyro stabilized inertial platform with three orthogonal accelerometers. Unlike gravimeters used in traditional airborne surveys, the AIRGrav system does not use any spring-type apparatus. The accelerometer is held within 10 arc seconds (0.0028 degrees) of level by a Schuler tuned inertial platform, monitored through the complex interaction of gyroscopes and accelerometers. This arrangement ensures that the gravimeter remains oriented vertically, independent of the manoeuvres of the aircraft.

The three-axis stabilized platform used in AIRGrav is controlled in a fundamentally different way to the platforms used with the Lacoste and Romberg (L&R) or Bell gravimeters. The L&R and Bell two-axis platforms rely on a control loop which attempts to null the horizontal accelerometers on the platform with a settable time constant. The setting of this time constant is a trade-off between averaging out the disturbing effects of horizontal accelerations and still being able to correct for the drift of the gyros. The control loop used in AIRGrav relies on the very low drift of its high-accuracy gyros, and the platform can be Schuler tuned to the period of 84 minutes so that it does not tilt when accelerated (Lawrence, 1992). This allows draped surveys to be flown and allows surveying in turbulence similar to what is practical for high resolution magnetometer surveys. This is demonstrated in Figure 1, which shows there is little correlation between noise level and turbulence.

Figure 1. A comparison of the standard deviation of the vertical accelerations on the gravimeter to the gravity noise, determined from a comparison to ground gravity, for five-minute periods on the same test lines. The turbulence levels experienced have no noticeable effect on the gravity data. The slope of a line fitted to these points is almost zero, and some of the segments flown under the most turbulent conditions have the least gravity noise. These lines were flown under normal daytime conditions over three weeks. The turbulence ranged from calm to moderate. Figure reproduced from Argyle et al., (2000).
The gravity sensor used in AIRGrav is a very accurate three-axis accelerometer with a wide dynamic range. This avoids the saturation that occurs with spring-type sensors during strong vertical motion of the aircraft. Testing has shown the accuracy of gravity recovery to be almost completely unaffected by aircraft dynamics up to what is considered “moderate” turbulence. The wide dynamic range also allows draped surveys to be flown, something which is nearly impossible with other gravimeters. The AIRGrav sensor has extremely low cross-coupling (i.e., errors introduced by deformation of the sensor and conversion of horizontal acceleration into vertical acceleration) due to the very tight servo control of its proof mass, and needs no corrections for this effect. The sensor is extremely linear and does not need a test range to be calibrated. The design of the platform is such that the sensor can be “tumbled” in the gravity field at a calibrated point to set the scale factor and offset.

**AIRGrav operations**

The AIRGrav system itself is lightweight, with an installed weight of less than 100 kg. This allows it to be installed in survey aircraft in conjunction with other geophysical equipment, including that necessary for magnetic surveys. The gravimeter has been flown in a range of aircraft; twin engine Cessna 404 and Britten-Norman Islanders, single turbine engine Cessna 208B Grand Caravans, as well as in a Eurocopter AS 350 helicopter. The resolution and accuracy of the final gravity data is primarily dependent on flying speed and line spacing rather than the type of aircraft.

There are no additional restrictions on flight parameters beyond those typically used for aeromagnetic surveys in terms of operating altitude, ground clearance, and turbulence. Level flight is not required; a draped surface can be used to follow the terrain. A short lead-in of 3-5 km is typically flown to minimize on-line effects from aircraft turns. In practice, a large portion of this extension data can be included in the final data set because turn effects are minimal. The system has been used successfully offshore, in mountainous terrain, in deserts in Africa and the Middle East, and in the Canadian high Arctic in winter conditions.

During field operations, the accelerometer scale factors and offsets can be calibrated at the survey base to correct for slow time variations in these parameters by rotating the platform through 180 degrees to measure the strength of the gravity of the Earth with positive and negative polarity. A test calibration range is not needed. The value used for gravity during calibration is ideally from a ground instrument or a known local gravity point, although it is possible to determine a local gravity value with sufficient accuracy using only the AIRGrav instrument. The system also automatically aligns and calibrates its gyros on start up before each flight, by determining the gyro drifts and XY bias.

Before and after each flight, the accuracy of the local gravity reading computed by the system is verified and instrument drift monitored by measuring gravity on the ground for 15 minutes or more. This period can also be used as an initialization period when the differential GPS data processing is performed, resulting in better accuracy. Daily before-flight and after-flight average readings are plotted and monitored for signs of instrument drift. Linear drift corrections have proved unnecessary and are not applied. Pre- and post-flight readings typically match within +/- 1 mGal, and variations over the course of a survey are only a few mGal. Gravity data can be tied to the local gravity system by comparing the measured gravity to existing gravity data where available, but this is not necessary. In many remote locations where the AIRGrav system has operated, the self calibrated AIRGrav measurements are potentially more accurate than the pre-existing gravity data.

**Data processing**

Whilst flying a survey, accelerations in an aircraft can reach 1 m/s, equivalent to 100 000 mGal. Data processing must extract gravity data from this very dynamic environment. This is achieved by modelling the movements of the aircraft in flight by extremely accurate GPS measurements. Dual frequency GPS receivers are employed on the aircraft and in ground reference stations used for differential GPS processing. Extensive research at SGL has resulted in high quality differential GPS processing techniques which are critical to achieving high resolution and high accuracy gravity data. In addition to providing the lowest possible noise levels, these techniques allow surveys to be completed successfully despite poor satellite geometry and high ionospheric activity, an essential requirement for surveys in far northern and southern latitudes.

Gravimeter data are recorded at 128 Hz. Accelerations are filtered and decimated to match the GPS, recorded at 10 Hz, using specially designed filters to avoid biasing the data. Gravity is calculated by subtracting the GPS derived accelerations from the measured accelerations. The calculated gravity is corrected for the Eötvös effect and normal gravity, and the sample interval is reduced to 2 Hz.
The following standard corrections are then applied to the gravity data to calculate the Bouguer anomaly.

a) Free air correction, \( g_{fa} \) in mGal given by

\[
g_{fa} = -0.3086 \cdot z
\]

where \( z \) is height of the aircraft in metres above the geoid.

b) Bouguer, \( g_{sb} \) given by

\[
g_{sb} = 2\pi\gamma \rho h
\]

where \( \gamma \) is the Universal Gravity constant, \( \rho \) is a density appropriate for the area, and \( h \) is height of the ground above the geoid.

c) Curvature correction, \( g_{ec} \) in mGal given by

\[
g_{ec} = (1.464e-3)h - (0.3533e-6)h^2 + (0.000045e-9)h^3
\]

where \( h \) is height of the ground above the geoid in metres.

d) Terrain, \( g_t \), computed using 2D FFT methods with a density appropriate for the area. Since the accuracy of the terrain correction is limited by the accuracy of the input terrain model, an on-board laser altimeter or LiDAR is employed which, in conjunction with differential GPS processing, enables the calculation of a digital terrain model in the survey area. As in ground surveys, poorly modelled terrain data could cause a significant error in the processed gravity data.

e) Level Correction, \( g_{lc} \), based on averages of intersection differences between primary and tie lines after low-pass filtering. Individual lines are levelled by applying a constant level shift to the entire line to minimise the differences between line and tie line intersections. In recent airborne gravity surveys only very minor constant per-line levelling corrections have been needed.

Applying these corrections, the complete Bouguer anomaly, \( g_{cba} \) is given by

\[
g_{cba} = g_{En} - g_{sb} - g_{ec} + g_t - g_{lc}
\]

where \( g_{En} \) is the calculated gravity adjusted for Eötvös effect and normal gravity.

The gravity data are filtered to remove noise using a low pass FFT filter prior to free air correction. Different filter settings are used depending on the application. After the standard corrections have been applied, the gravity data are gridded using the same minimum curvature gridding algorithm that is used for aeromagnetic data. A 2D spatial filter is then applied to the gridded data to maximize the resolution of the data by allowing noise cancellation between adjacent lines.

**Turner Valley case study**

In 2001, SGL flew a large AIRGrav survey in Western Canada over the Turner Valley area, an oil and gas producing region south of Calgary, Alberta (Pierce et al., 2002a, b; Sander et al., 2002). The survey area covered the foothills of the Rocky Mountains. The general trend of the geology in the area is north-north-west / south-south-east. A total of 12 500 line km of AIRGrav and magnetometer data were acquired simultaneously using a fixed-wing aircraft. The survey was completed in less than five weeks over very mountainous terrain, ranging from approximately 1000 m ASL to approximately 2000 m ASL. East-west traverse lines were spaced at 250 m and north-south tie lines at 1000 m.

Figure 2 shows the AIRGrav and aeromagnetic data acquired during the Turner Valley Survey. The colour image displays the first vertical derivative of Bouguer gravity, with the warm colours representing vertical gravity gradient highs. The grey shades are the shadow of the first vertical derivative of the aeromagnetic data.
The western side of the survey area is dominated by north-south trending faults associated with the foothills region. The eastern side of the area consists of flat lying sediments. Gas and oil producing areas are outlined by solid and dotted lines respectively. The long north-south trending field in the centre of the area is the Turner Valley Field, discovered in 1914. The Turner Valley and Quirk Creek Fields (in the north-west of the area) are generally associated with gravity highs. East-west trending lineaments, marked with dashed lines, were defined by joining terminations seen in the aeromagnetic data. The same lineaments also mark changes in the gravity signal, and offset the Turner Valley Field.

Existing ground gravity readings were compared with the AIRGrav data. Figure 3 shows the very good agreement between the two data sets. The uniform sampling provided by the airborne survey, however, brings out additional detail in areas that are sparsely sampled by ground data. Creating a grid of airborne gravity using only points where ground gravity exists improves the apparent correlation with the ground data, an indication that differences are due to improved airborne sampling (Pierce, 2002). The ability to provide uniform sampling over difficult terrain that is often inaccessible to ground surveys is one of the advantages of an airborne survey. The AIRGrav system is particularly well suited to this task it has low sensitivity to aircraft motion. The survey area is subject to frequent periods of turbulence caused by winds blowing down from the Rocky Mountains, and the climb into the foothills required a draped flying surface.

Estimate of noise

Noise levels and system resolution require careful specification. Survey parameters, along with the instrument itself, are important factors in the resolution and accuracy of the final data products delivered to a client. A survey flown with tight line spacing, for example, takes advantage of over-sampling to reduce noise levels (Sander et al., 2003). Flying speed plays a role in determining AIRGrav resolution since the primary source of noise is from the GPS signal, which is time dependant. It is possible to increase the resolution to a degree simply by flying slower.

AIRGrav typically delivers gravity data which have been low-pass filtered using a filter of half-length 40 s. Converted to wavelength, this equals a half-width of 2.0 km in a Cessna Grand Caravan or Britten-Norman Islander flown at 100 knots, 2.9 km in a Cessna 404 flown at 140 knots, and 1.0 km in a Eurocopter AS350 helicopter flown at 50 knots. These numbers are based on the repeatability of 40 test lines flown near Ottawa using different aircraft.

Once the objectives of the survey are established, line spacing and flying speed are tailored towards meeting those objectives. In a recent petroleum survey flown in the Middle East, for example, 500 m line spacing was used. The final grid had real anomalies as small as 0.4 mGal and 2.0 km in size that were independently verified in a blind test using detailed ground measurements. Noise levels on the final grid data
were 0.2 mGal, based on a measurement of internal consistency made using independent grids formed from alternate lines.

This method of comparing independent grids of airborne data to estimate the precision of the measurements was applied in the Turner Valley example (Sander et al., 2002). The line spacing of 250 m over-sampled the data given the application of a low-pass filter with 4 km cut-off. Two independent grids were produced by separating the data into two subsets (i.e., odd and even line number grids with 500 m line spacing). The rms value of the differences between these grids indicated a precision of 0.3 mGal for the complete dataset.

A case study for a survey in Timmins (Ontario, Canada) (Elieff, 2003; Elieff et al., 2004) quantified noise levels using a number of internal and external measures. This survey was flown with 500 m line spacing, and the gridded data were low-pass filtered with a cut-off wavelength of 2.85 km. The standard deviation of differences at flight and tie line crossovers indicated a precision of 0.45 mGal for the filtered line data. Differences between odd and even flight line number grids had a standard deviation that indicated a precision of 0.15 mGal in the final grid after spatial filtering. The standard deviation of differences between the airborne data and upward continued ground data was 0.62 mGal. This provides an upper bound estimate of the external accuracy of the system since it also includes errors present in the ground data and differences resulting from application of filtering to the airborne dataset but not to the ground dataset.

Conclusions
AIRGrav has been used to acquire airborne gravity data for both on-shore and off-shore surveys. The system allows drape surveys to be flown, resulting in better signal levels in surveys areas with significant topographic relief. Final grids created with low-pass filters in the 2 – 4 km range have noise standard deviations between 0.15 and 0.3 mGal. This number is dependant upon survey parameters, such as line spacing and aircraft speed, as well as the details of the specific filter used.

References
Figure 3. Complete Bouguer Gravity data from the Turner Valley area. The images in (a) and (c) are based on older ground gravity data. These data have been upward continued 500 m and filtered with a 5300 m low-pass filter to match the airborne parameters as closely as possible. The images in (b) and (d) are from the Turner Valley Airborne Survey. The images in (a) and (b) show complete Bouguer gravity data, whilst (c) and (d) show the first vertical derivative of complete Bouguer gravity data. The regional slope reflects the isostatic effect of the Rocky Mountains, just west of the survey area. Figure reproduced from Pierce et al. (2002b).